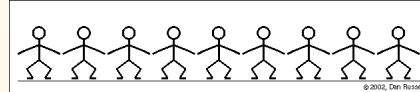


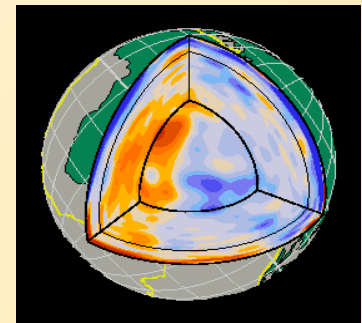
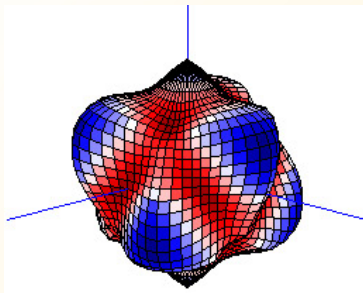
Physical properties in **M**TEX, current status and future developments

Chemnitz **M**TEX Workshop March 2019
Technische Universität Chemnitz, Germany

David Mainprice



(Géosciences Montpellier, Université de Montpellier, France)



CNRS - Université de Montpellier, 34095 Montpellier, France

Plan of this talk, is there a plan ?

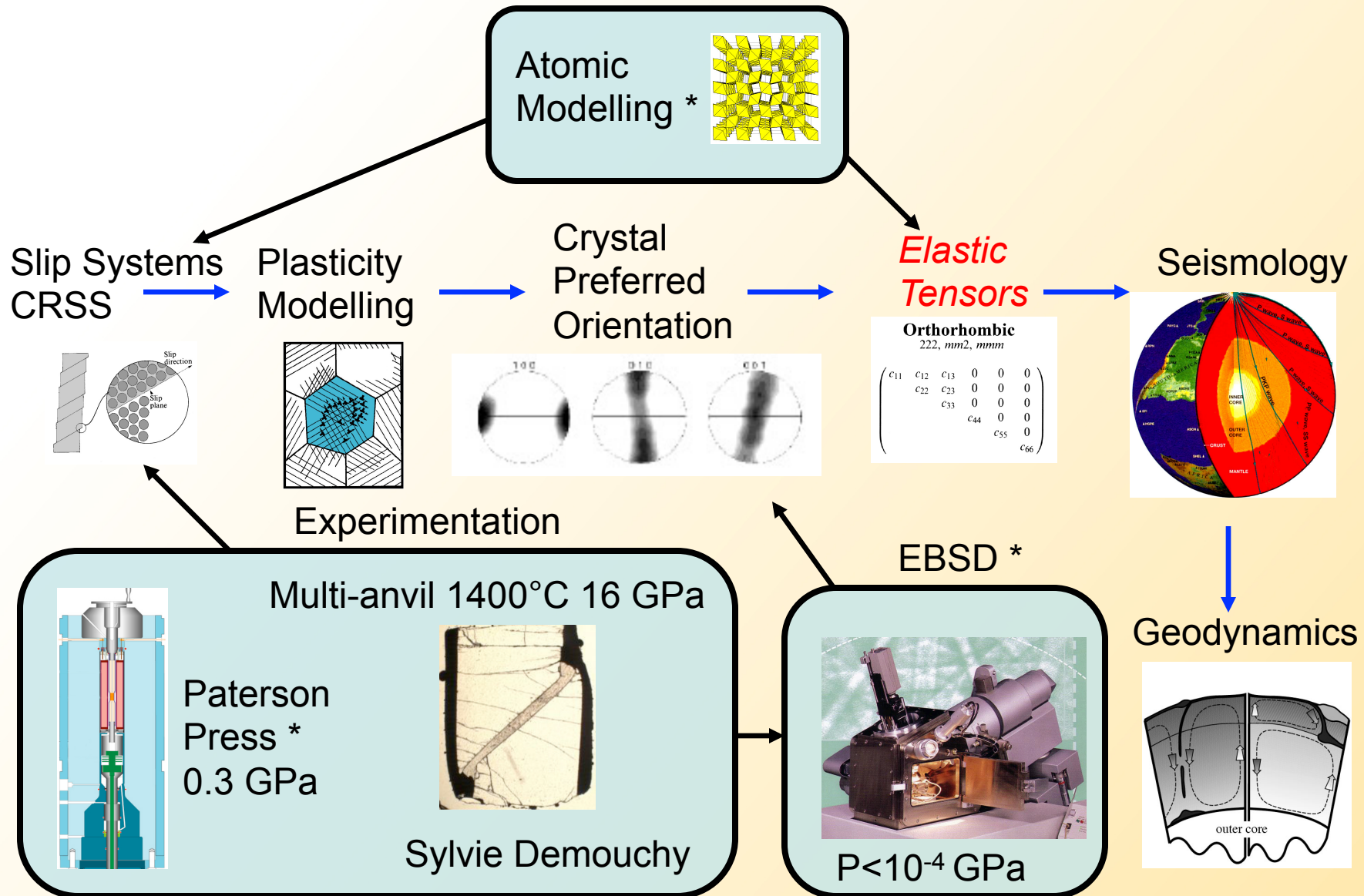
- A brief resumé of what I do ...sometimes
- Single crystal introduction tensor basics – starting with 2nd rank tensors
- Single crystal elasticity tensor basics – Young's modulus to wave propagation
- Effective media : average properties for polycrystalline aggregates, Voigt, Reuss and Hill...
- Future developments 1; Self-consistent (SC) and Differential Effective Media (DEM), FFT methods
- Future developments 2; Importing single crystal tensors from a database, advantages and pitfalls



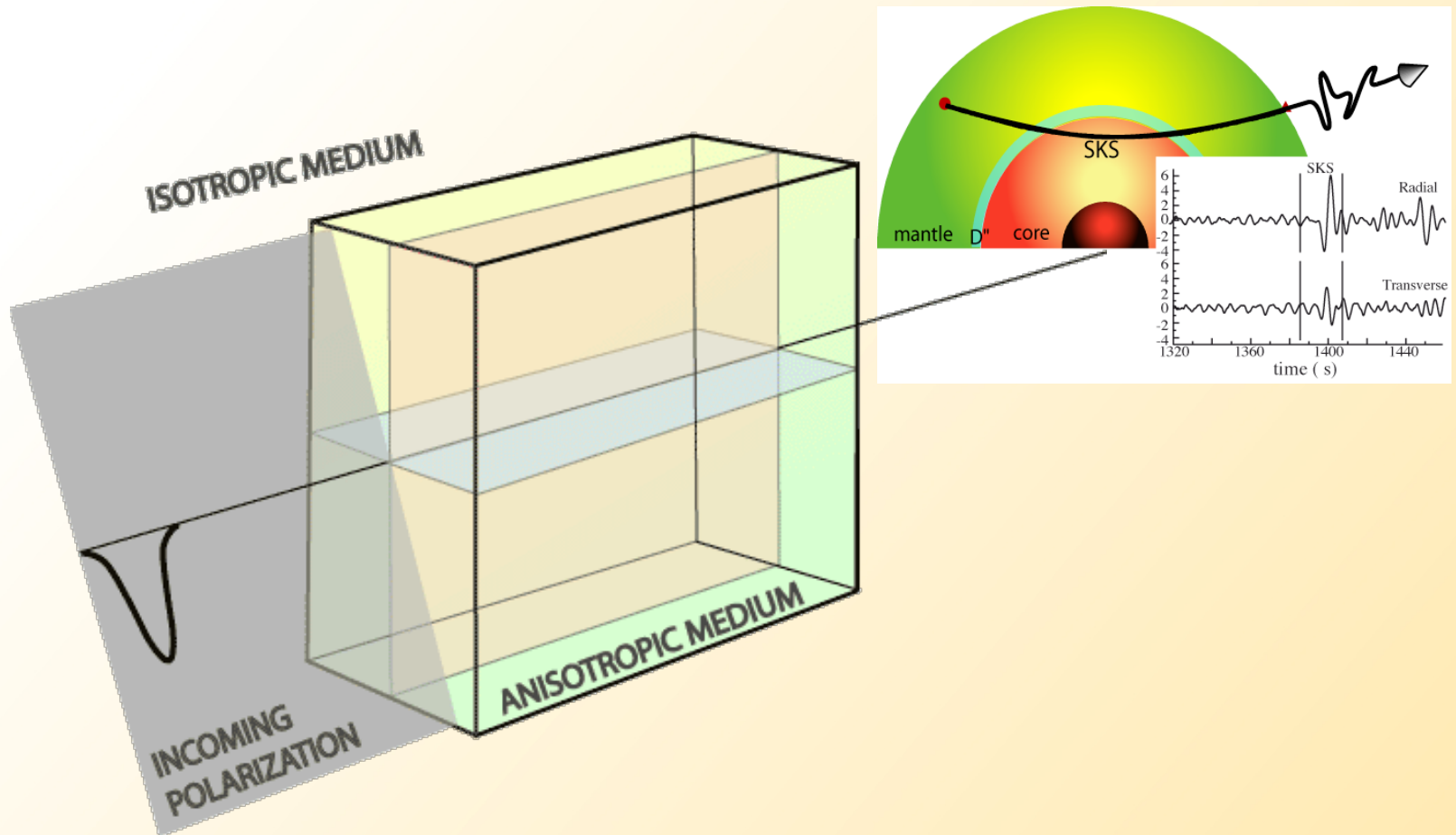
A brief resumé of what I do ...sometimes

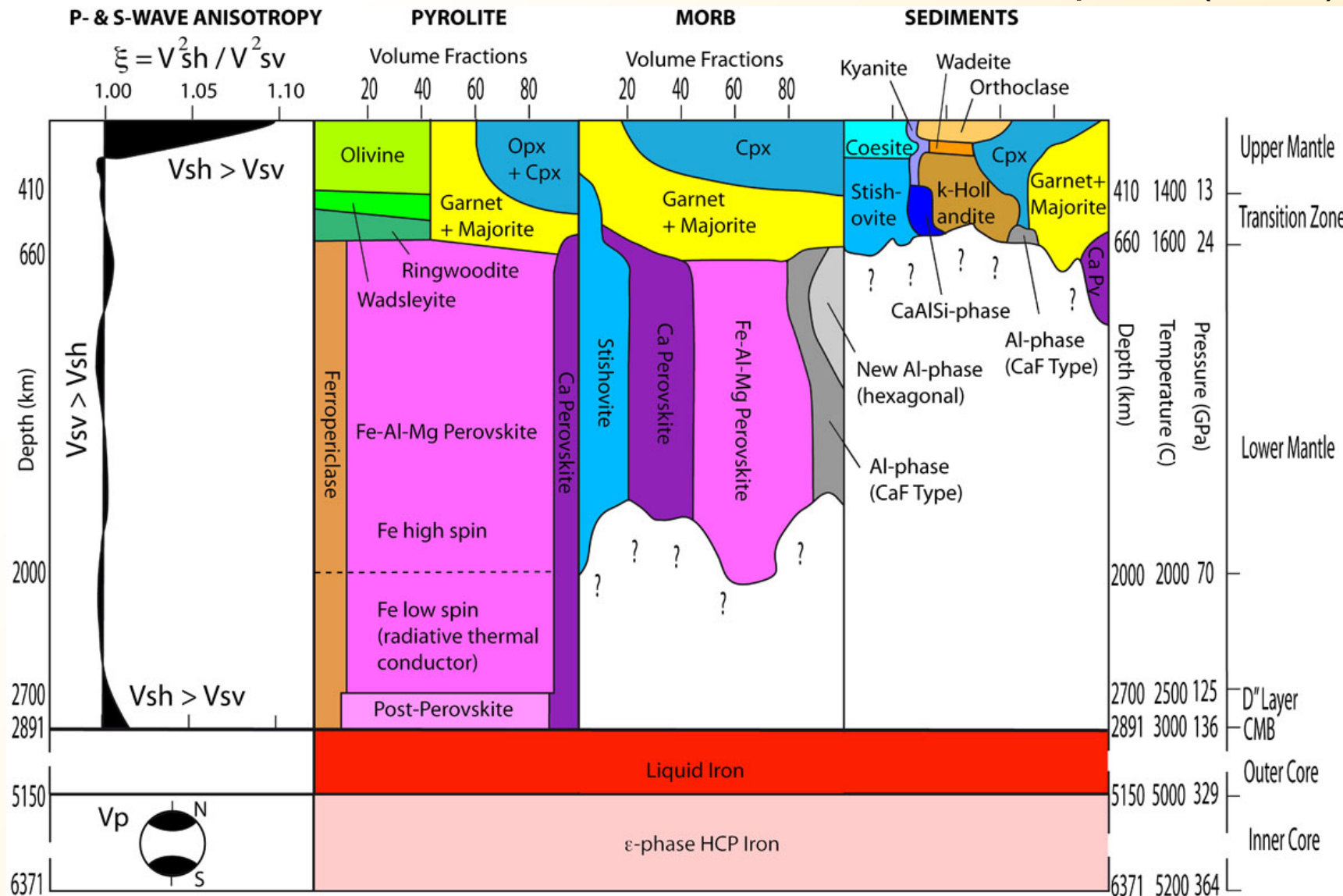
What do I do ?

Atomic Modelling & Experimentation at high P & T



Shear Wave Splitting



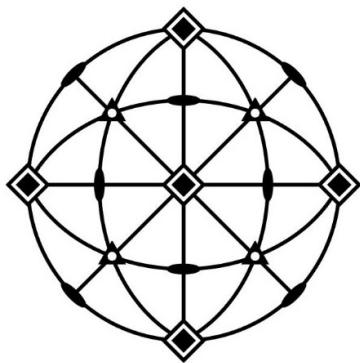


Getting Started

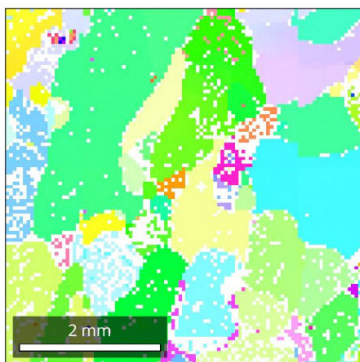
• [Installation Guide](#) • [Configuration](#) • [Frequently Asked Questions](#) • [Release Notes](#) • [Function Overview](#) • [Matlab Basics](#) •

Users Guide

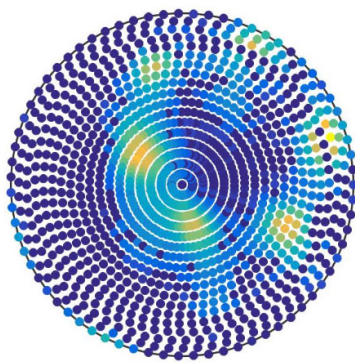
Crystal Geometry



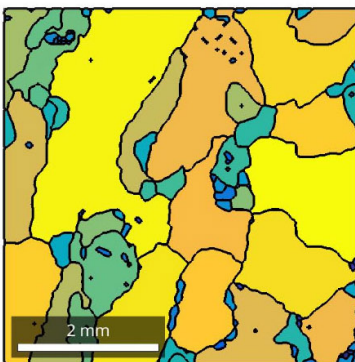
EBSD



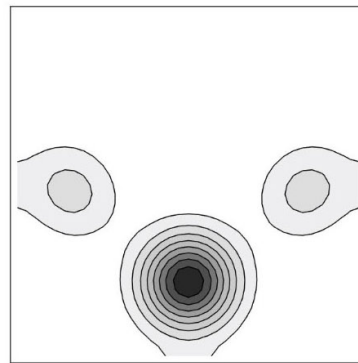
Pole Figures



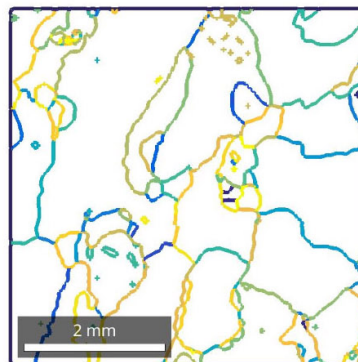
Grains



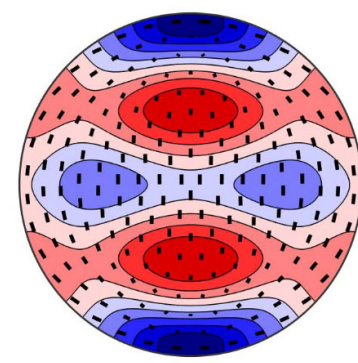
ODFs



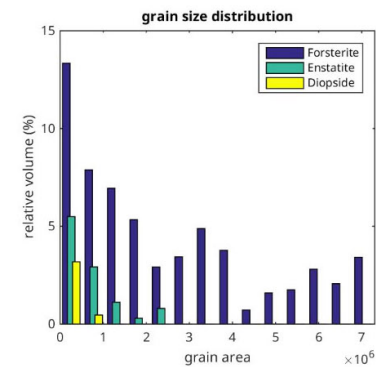
Grain Boundaries



Tensors



Plotting



MTEX Reference papers on physical properties

<http://mtex-toolbox.github.io/publications.html>

MTEX Toolbox

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Reference Publications


If you are using the library for your research, please cite this homepage with the following sample BibTeX entry.

2014

- **Calculating anisotropic piezoelectric properties from texture data using the MTEX open source package**, D. Mainprice, F. Bachmann, R. Hielscher, H. Schaeben, G. E Lloyd: [Geological Society, London, Special Publications, 409, 2014](#), 

2011

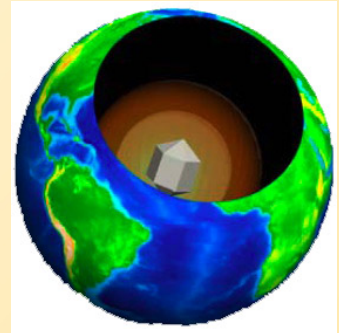
- **Calculating anisotropic physical properties from texture data using the MTEX open source package**, D. Mainprice, R. Hielscher, H. Schaeben: in Prior, D.J., Rutter, E.H., Tatham, D. J. (eds) [Deformation Mechanisms, Rheology and Tectonics: Microstructures, Mechanics and Anisotropy. Geological Society, London, Special Publications, 360, 175-192](#). 



Single crystal introduction tensor
basics – starting with 2nd rank
tensors

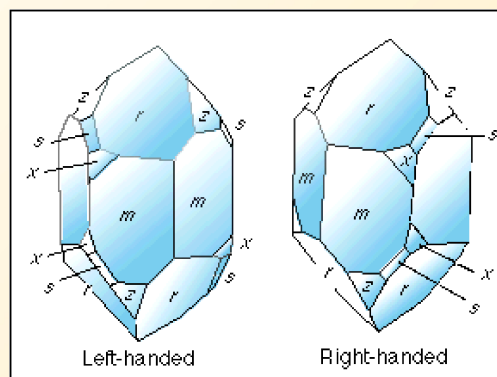
Why are we interested in Single Crystals ?

- To understand the anisotropic physical properties of polycrystalline rocks caused by crystal preferred orientation (CPO) it is important to know the about the simplest case, the single crystal.
- The single orientation (single crystal), has a perfectly defined ODF (orientation distribution function), PFs (pole figure) or IPFs (inverse pole figure).
- To understand the how crystal symmetry, sample symmetry, CPO and single crystal properties combined to produce anisotropic rock properties ...



Physical properties of crystals

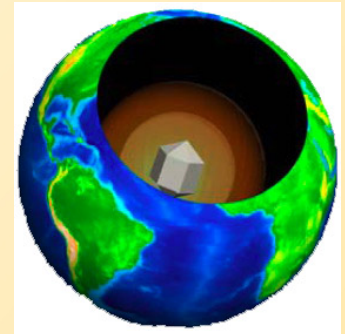
- Thermal conductivity and diffusivity (2th rank tensor) → can be calculated from CPO
- Thermal expansion (2th rank tensor) → can be calculated from CPO
- Electrical conductivity, electrical polarization and dielectric properties) → can be calculated from CPO, BUT may not be relevant if conductivity controlled by high conductivity phases in the grain boundaries (e.g. water or carbon)
- Piezoelectricity (3rd rank tensor) → can be calculated from CPO, if we can determine the CPO of the Left- and Right-handed crystals...



- Elasticity (4th rank tensor) → seismic (elastic) properties, can be calculated from CPO

Why are we interested in tensors of single Crystals ?

- To understand the anisotropic physical properties of polycrystalline materials caused by crystal preferred orientation (**CPO**) or **Texture** it is important to know the about the simplest case, the **single crystal**.
- The single orientation (single crystal), has a perfectly defined **ODF** (orientation distribution function), **PFs** (pole figure) or **IPFs** (inverse pole figure).
- To understand the how **crystal symmetry**, **sample symmetry**, **CPO or Texture** and **single crystal properties** combined to produce anisotropic rock properties ...



Typical choices for tensor reference frames

Crystal Symmetry	x	y	z
Orthorhombic, tetragonal, cubic	a	b	c
Trigonal, hexagonal	a m	m -a	c c
Monoclinic	a* a	b b	c c*
Triclinic	a* y x z y x z	z x x b* b	c* c c*

Cartesian Tensors II

- Note it is a fundamental concept for tensors of physical properties ***that the magnitude physical property does not change*** (i.e. it is invariant) with the rotation of the co-ordinate frame, which is attached to the crystal.
- The crystal may rotate in the sample (external) reference frame, but the tensor reference frame is related to the crystal structure and rotates with the crystal, like a rigid body, in other words it is a co-ordinate transform.

Table 1. Linear orthogonal transformation laws for Cartesian tensors

Name	Rank	New orientation in terms of old	Old orientation in terms of new
Scalar	0	$\mathbf{S}' = \mathbf{S}$	$\mathbf{S} = \mathbf{S}'$
Vector	1	$\mathbf{V}'_i = \mathbf{R}_{ij} \mathbf{V}_j$	$\mathbf{V}_i = \mathbf{R}_{ji} \mathbf{V}'_j$
2 nd rank Tensor	2	$\mathbf{T}'_{ij} = \mathbf{R}_{ik} \mathbf{R}_{jl} \mathbf{T}_{kl}$	$\mathbf{T}_{ij} = \mathbf{R}_{ki} \mathbf{R}_{lj} \mathbf{T}'_{kl}$
3 rd rank Tensor	3	$\mathbf{T}'_{ijk} = \mathbf{R}_{il} \mathbf{R}_{jm} \mathbf{R}_{kn} \mathbf{T}_{lmn}$	$\mathbf{T}_{ijk} = \mathbf{R}_{li} \mathbf{R}_{mj} \mathbf{R}_{nk} \mathbf{T}'_{lmn}$
4 nd rank Tensor	4	$\mathbf{T}'_{ijkl} = \mathbf{R}_{im} \mathbf{R}_{jn} \mathbf{R}_{ko} \mathbf{R}_{lp} \mathbf{T}_{mnop}$	$\mathbf{T}_{ijkl} = \mathbf{R}_{mi} \mathbf{R}_{nj} \mathbf{R}_{ok} \mathbf{R}_{pi} \mathbf{T}'_{mnop}$

S = scalar, V=vector, T=tensor. \mathbf{T}'_{ij} is the **new orientation** of the tensor and \mathbf{T}_{ij} the **old orientation**.

The rotation or transformation matrix \mathbf{R}_{ij} used to calculate the tensor in new orientation is replaced by its transpose \mathbf{R}_{ji} to calculate the old orientation.

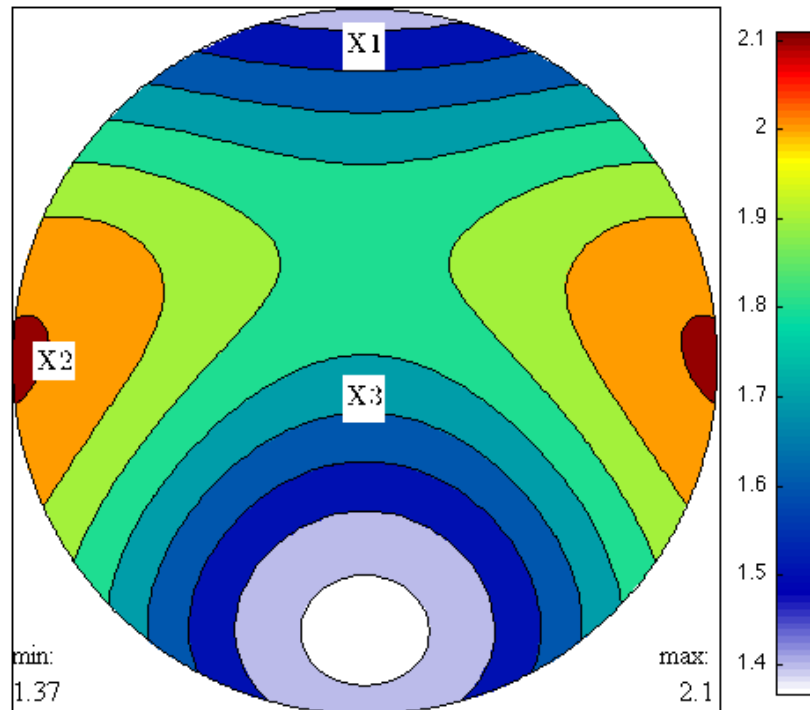
Cartesian tensors have orthogonal reference axes (labelled **x**, **y** and **z** in **MTEX**)

Representing single crystal properties

```

%%
%*****
% Plot thermal conductivity tensor (W/m/K units)
%*****
%
plot(k, 'complete');
colorbar
%
% annotate with tensor orthogonal axes X1,X2 and X3
hold on
plot([xvector,yvector,zvector], 'data', {'X1', 'X2', 'X3'}, 'backgroundcolor', 'w');
hold off
%
% save plot as *.png file
savefigure('Plot_Orthoclase_Single_Crystal_TC_Quadric.png');
%

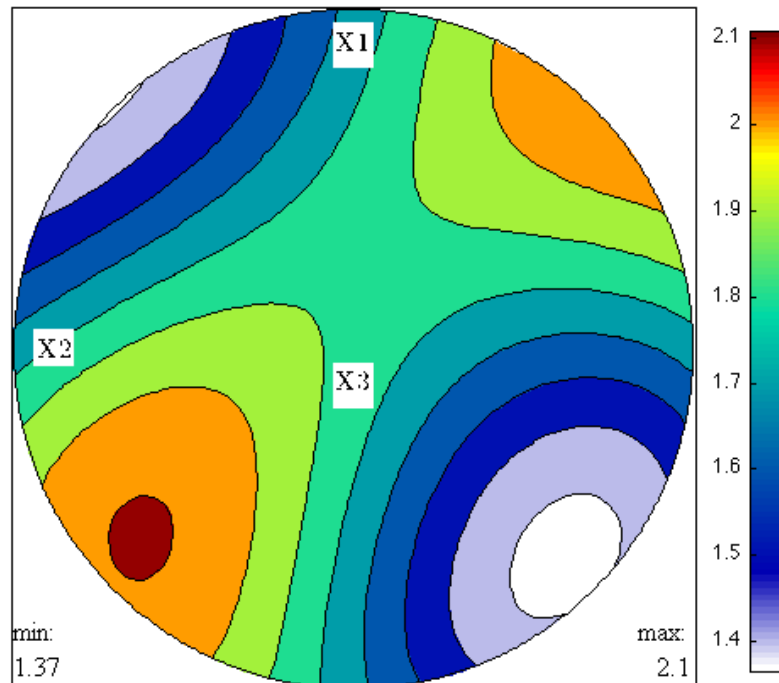
```




```

%%
%*****
% Plot rotated thermal conductivity tensor (W/m/K units)
%*****
% define rotation
g = orientation('Euler',10*degree,20*degree,30*degree,cs_tensor);
% rotate tensor
k_rot = rotate(k,g);
% plot
plot(k_rot,'complete');
colorbar
%
% annotate with tensor orthogonal axes X1,X2 and X3
hold on
plot([xvector,yvector,zvector],'data',{'X1','X2','X3'},'backgroundcolor','w');
hold off
%
% save plot as *.png file
savefigure('Plot_Orthoclase_Single_Crystal_TC_Quadric_rot.png');
%

```



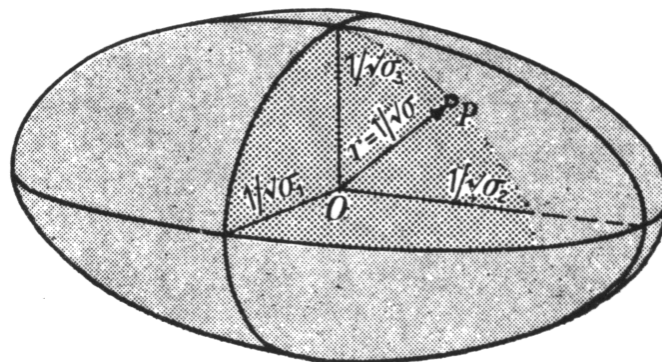
Difference between a transformation matrix (\mathbf{R}_{ij}) and a 2nd rank tensor (\mathbf{T}_{ij})

\mathbf{R}_{ij} is an 3 by 3 matrix relating two (right-handed) reference frames [orthogonal matrix $\mathbf{R}^t \cdot \mathbf{R} = \mathbf{R} \cdot \mathbf{R}^t = \mathbf{I}$, $\mathbf{R}^t = \mathbf{R}^{-1}$ and $\text{Det}(\mathbf{R}) = +1$ where the rows and columns are orthogonal (orthonormal) unit vectors.] (e.g. rotation or orientation matrix) N.B. when \mathbf{R}_{ij} is relating right-handed to left-handed reference frames $\text{Det}(\mathbf{R}) = -1$.

\mathbf{T}_{ij} is a physical quantity (e.g. 2nd rank tensor) that for a given set of reference axes is represented by 9 numbers (a 3 by 3 table).

The representation **quadric** for 2nd Rank Tensors

Geometrical representation of symmetrical second-rank tensors as a second-degree surface (called a **quadric**). The quadric may be an ellipsoid or a hyperboloid. Most common second-rank tensors are symmetric ($T_{ij} = T_{ji}$) and when the 3 principal coefficients are all positive then the property is represented by an ellipsoid with axes $1/\sqrt{T_1}$, $1/\sqrt{T_2}$ and $1/\sqrt{T_3}$, which is the case for electric polarization, electrical and thermal conductivity and optical properties.

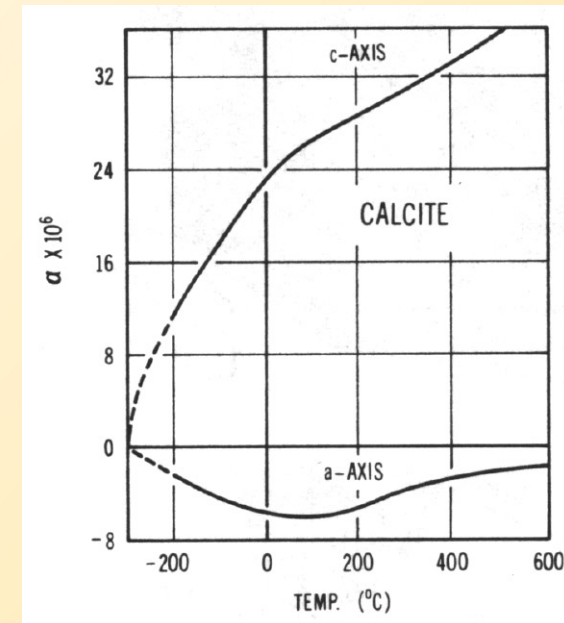
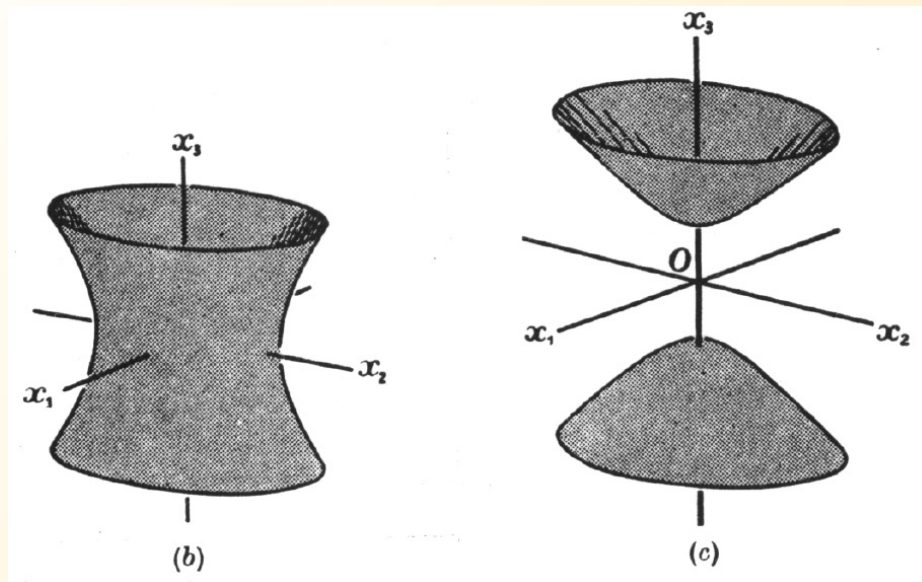


Illustrating the properties of the representation ellipsoid for conductivity:

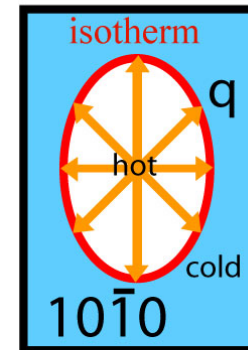
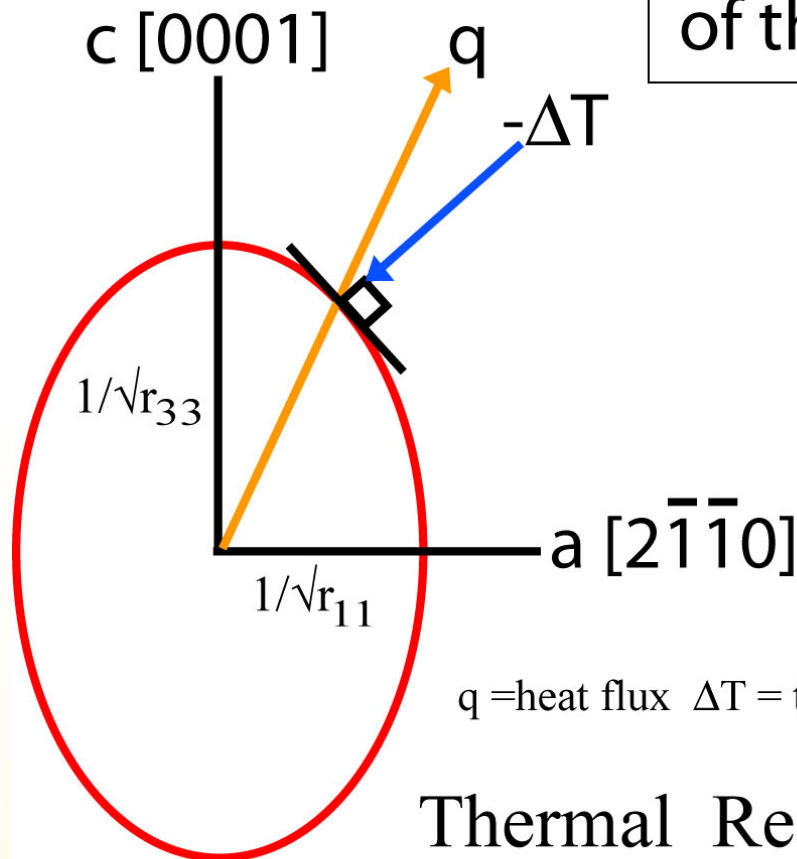
$$\sigma_{ij} x_i x_j = 1.$$

The representation **hyperboloid** for 2nd Rank Tensors

If **one** of the principal coefficients is **negative** then surface is a hyperboloid of one sheet (e.g. thermal expansion of plagioclase feldspar). If **two** of the principal coefficients are **negative** then surface is a hyperboloid of two sheets or caps, this the case for the thermal expansion of calcite with contraction the basal plane. If all three of the principal coefficients are negative then surface is an imaginary ellipsoid, this is the case for many susceptibilities of paramagnetic and diamagnetic minerals, such as quartz, calcite and aragonite.



Radius-normal property of the representation quadric



Radial heat flux (q)
(controlled by crystal
properties not boundary
conditions)

q = heat flux ΔT = thermal gradient

Thermal Resistivity (r)

$$\Delta T_i = - r_{ij} q_j$$

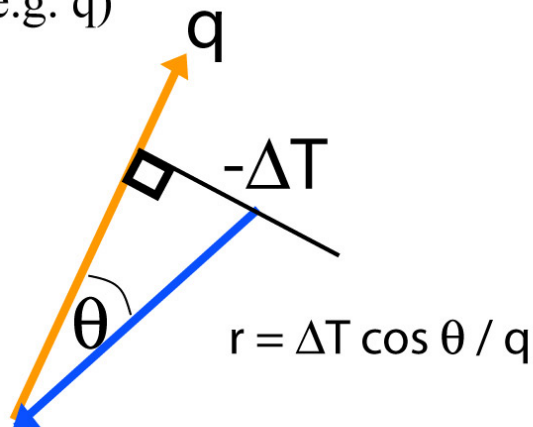
where $r_{11} = 1/k_{11}$ etc

Thermal Conductivity (k)

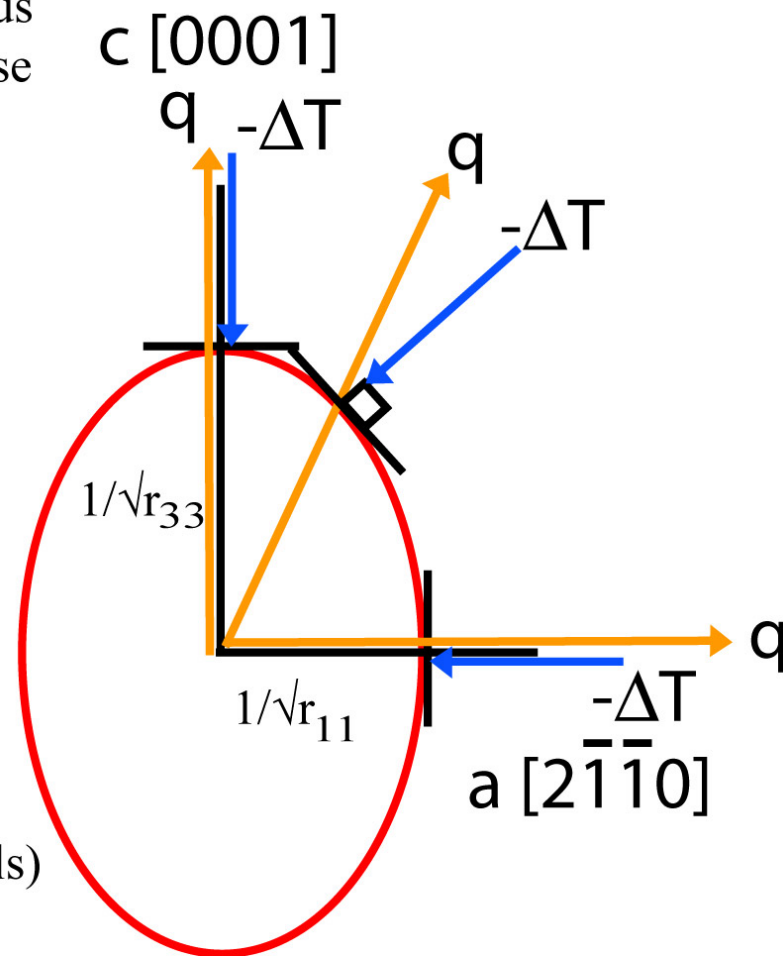
$$q_i = - k_{ij} \Delta T_j$$

The stimulus vector and response vectors for 2nd rank tensors

In general in anisotropic crystals the stimulus vector (e.g. ΔT) is not parallel to the response vector (e.g. q)



In certain directions (e.g. principal axes) the stimulus vectors and the response vectors may be parallel in anisotropic crystals. It is always the case for isotropic crystals (e.g. 2nd rank tensors of cubic crystals)

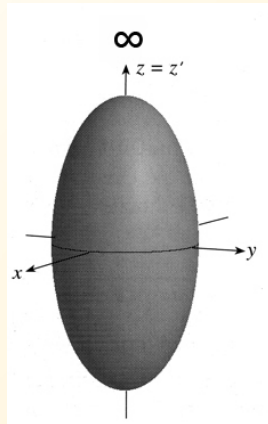


Fundamental concept for tensors

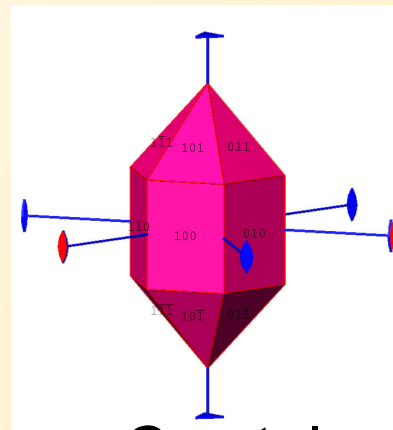
- Physical properties are anisotropic when the applied tensor is not aligned with resultant tensor
- Physical properties are isotropic when the applied tensor is aligned with resultant tensor for all directions
- In specific directions in an anisotropic crystal the applied and resultant tensor may be aligned giving rise to local isotropic or pseudo-isotropic behaviour

Effect of symmetry on Physical Properties : Neumann's Principle

F.E. Neumann's principle (1885) states that "symmetry elements of any physical property of a crystal must include ALL the symmetry elements of the point group of the crystal". This implies that that a given physical property may possess a **higher** symmetry than that possessed by the crystal and it cannot be of a **lower** symmetry than that of the crystal. Some physical properties are inherently centrosymmetric (all symmetric second order tensors and elasticity) which **will add a center of symmetry** in many minerals (e.g. quartz) and result in a **higher** symmetry than the possessed by the crystal.



Physical property
higher symmetry



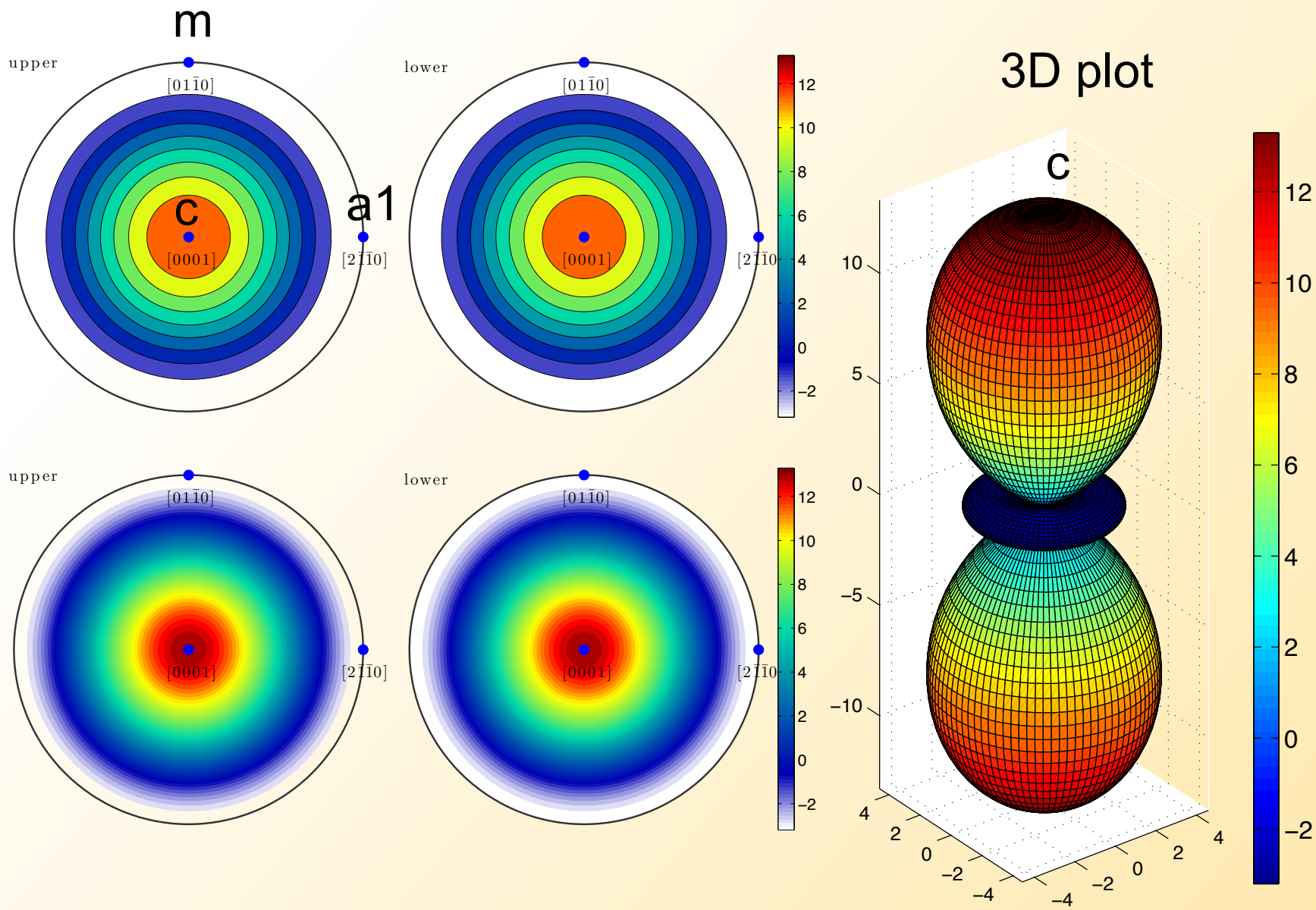
Crystal
lower symmetry

Symmetric 2nd rank tensors for all crystal symmetries

Crystal Symmetry (Independent constants)	Symmetry and tensor axes	Tensor						
Cubic (1)	Isotropic $T_{11} = T_{22} = T_{33}$	$\begin{bmatrix} T_{11} & 0 & 0 \\ 0 & T_{11} & 0 \\ 0 & 0 & T_{11} \end{bmatrix}$						
Hexagonal (2) Trigonal on hexagonal axes (2) Tetragonal (2)	6-fold \parallel z 3-fold \parallel z 4-fold \parallel z $T_{11} = T_{22} \neq T_{33}$	$\begin{bmatrix} T_{11} & 0 & 0 \\ 0 & T_{11} & 0 \\ 0 & 0 & T_{33} \end{bmatrix}$						
Orthorhombic (3)	2-fold \parallel x 2-fold \parallel y 2-fold \parallel z $T_{11} \neq T_{22} \neq T_{33}$	$\begin{bmatrix} T_{11} & 0 & 0 \\ 0 & T_{22} & 0 \\ 0 & 0 & T_{33} \end{bmatrix}$						
Monoclinic (4)	2-fold \parallel y $T_{13} = T_{31}$ 2-fold \parallel z $T_{12} = T_{21}$ 2-fold \parallel x $T_{23} = T_{32}$ & $T_{11} \neq T_{22} \neq T_{33}$	<table style="width: 100%; border-collapse: collapse;"> <tr> <td style="text-align: center; padding: 5px;">2-fold \parallel y</td> <td style="text-align: center; padding: 5px;">2-fold \parallel z</td> <td style="text-align: center; padding: 5px;">2-fold \parallel x</td> </tr> <tr> <td style="text-align: center; padding: 5px;">$\begin{bmatrix} T_{11} & 0 & T_{13} \\ 0 & T_{22} & 0 \\ T_{13} & 0 & T_{33} \end{bmatrix}$</td> <td style="text-align: center; padding: 5px;">$\begin{bmatrix} T_{11} & T_{12} & 0 \\ T_{12} & T_{22} & 0 \\ 0 & 0 & T_{33} \end{bmatrix}$</td> <td style="text-align: center; padding: 5px;">$\begin{bmatrix} T_{11} & 0 & 0 \\ 0 & T_{22} & T_{23} \\ 0 & T_{23} & T_{33} \end{bmatrix}$</td> </tr> </table>	2-fold \parallel y	2-fold \parallel z	2-fold \parallel x	$\begin{bmatrix} T_{11} & 0 & T_{13} \\ 0 & T_{22} & 0 \\ T_{13} & 0 & T_{33} \end{bmatrix}$	$\begin{bmatrix} T_{11} & T_{12} & 0 \\ T_{12} & T_{22} & 0 \\ 0 & 0 & T_{33} \end{bmatrix}$	$\begin{bmatrix} T_{11} & 0 & 0 \\ 0 & T_{22} & T_{23} \\ 0 & T_{23} & T_{33} \end{bmatrix}$
2-fold \parallel y	2-fold \parallel z	2-fold \parallel x						
$\begin{bmatrix} T_{11} & 0 & T_{13} \\ 0 & T_{22} & 0 \\ T_{13} & 0 & T_{33} \end{bmatrix}$	$\begin{bmatrix} T_{11} & T_{12} & 0 \\ T_{12} & T_{22} & 0 \\ 0 & 0 & T_{33} \end{bmatrix}$	$\begin{bmatrix} T_{11} & 0 & 0 \\ 0 & T_{22} & T_{23} \\ 0 & T_{23} & T_{33} \end{bmatrix}$						
Triclinic (6)	x, y & z in arbitrary orientations 6 independent values $T_{11} \neq T_{22} \neq T_{33} \neq T_{12} \neq T_{13} \neq T_{23}$	$\begin{bmatrix} T_{11} & T_{12} & T_{13} \\ T_{21} & T_{22} & T_{23} \\ T_{13} & T_{32} & T_{33} \end{bmatrix}$						

x, y and **z** are the orthogonal reference axes for Cartesian tensors

Thermal expansion of calcite



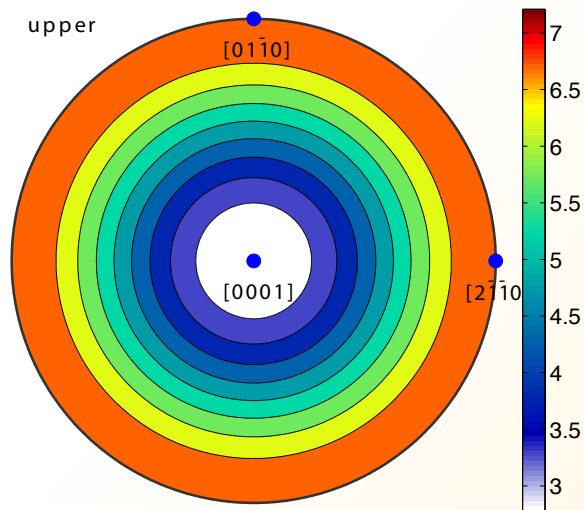
```

%*****
% Step 1 : Define the tensor reference frame for Calcite
%*****
cs_tensor = crystalSymmetry('-3m1',[4.9896, 4.9896, 17.0610],...
    [90, 90, 120]*degree,'mineral','calcite','x||a','z||c');
%*****
% Step 2 : Import 2nd rank tensor as 3 by 3 matrix M
%*****
% Thermal expansion Tensor for Calcite given by Fei (1995) AGU Bookshelf
% Series Mineral Physics vol.2 Chapter 6 Thermal expansion
% in x 10-6 1/K units
%
% a11=-3.20 a12=0.00 a13=0.00
% a21=0.00 a22=-3.20 a23=0.00
% a33=0.00 a32=0.00 a33=13.3
%
% Enter tensor as matrix,M, line by line.
%
M = [[ -3.20  0.00  0.00 ]; ...
     [  0.00 -3.20  0.00 ]; ...
     [  0.00  0.00 13.30 ]];
% Define tensor object with MTEX command tensor
% for the alpha(ij) thermal expansion with units x10-6 1/K
%
alpha_calcite = tensor(M,'name','thermal expansion','unit','x10-6
1/K',cs_tensor)
%*****
% Step 3 : Plot thermal expansion tensor of calcite
%*****
% Create list of crystallographic directions
List = [Miller(2,-1,-1,0,cs_tensor,'UVTW'),...
        Miller(0,1,-1,0,cs_tensor,'UVTW'),...
        Miller(0,0,0,1,cs_tensor,'UVTW')]
% plot 2nd rank tensor
plot(alpha_calcite,'complete')
colorbar
% annotate with crystal directions
hold on
plot(List,'labeled')
hold off
% save plot as *.pdf file
saveFigure('Plot_Calcite_Single_Crystal_alpha_tensor.pdf')

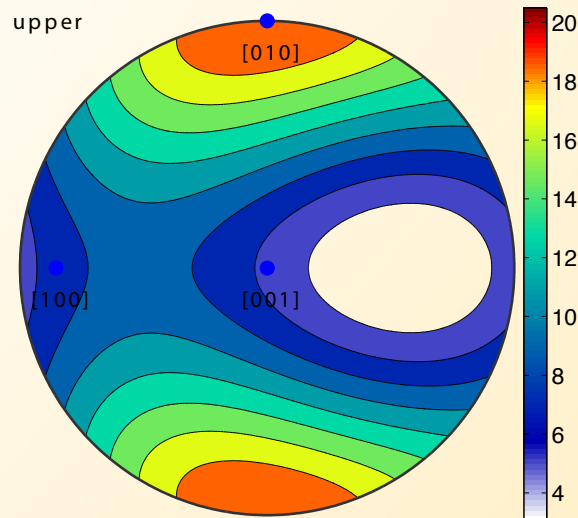
```

Single crystal thermal expansion – effect of symmetry

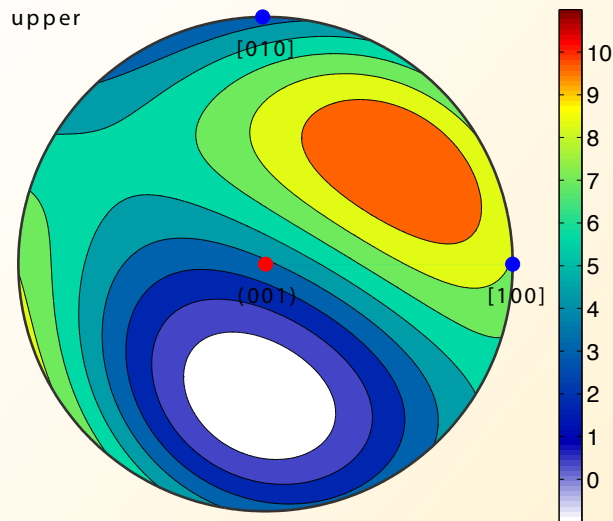
α -Quartz (trigonal -3m)



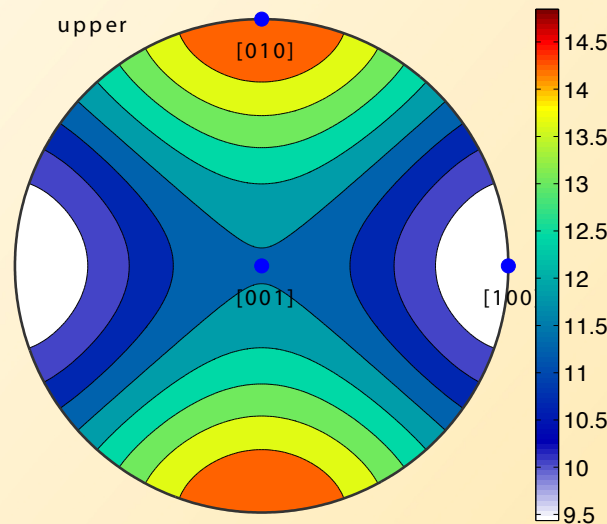
Diopside (monoclinic 12/m1)



Bytownite An76 (triclinic -1)



Forsterite (Orthorhombic mmm)



Optical properties - olivine

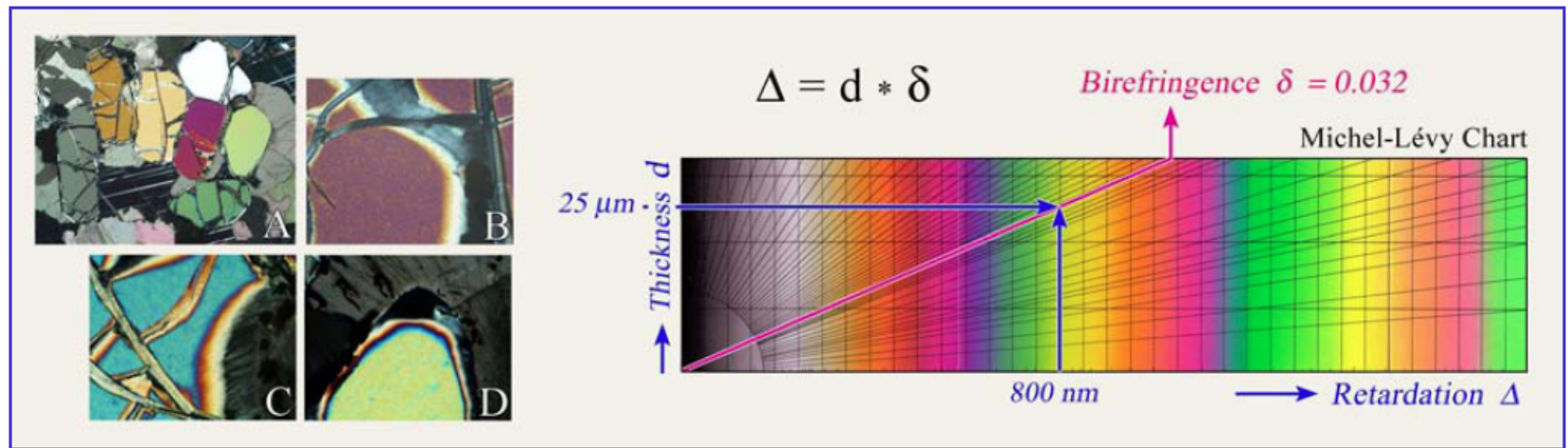


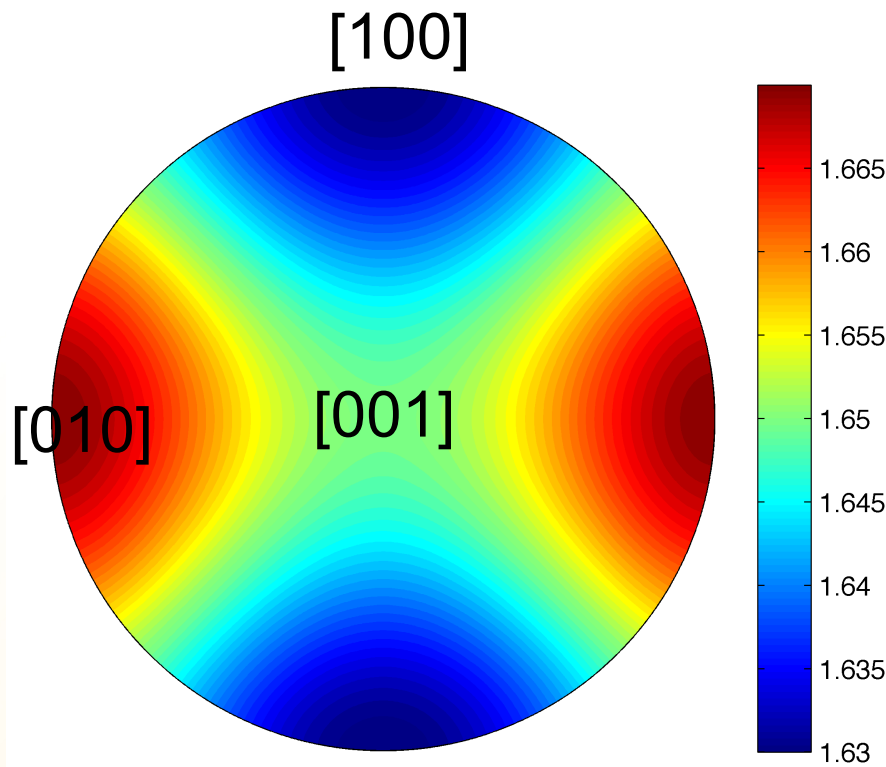
Figure 4.2.3-8: Left: (A) Variation of interference colour in differently oriented forsterite grains; (B) to (D) Decreasing interference colours at wedging-out edges of sections of forsterite grains that have different orientation in thin section. Right: Determination of birefringence using crystal thickness and interference colour; example of forsterite grain (D).

M.M. Raith , P. Raase, J. Reinhardt (2011)

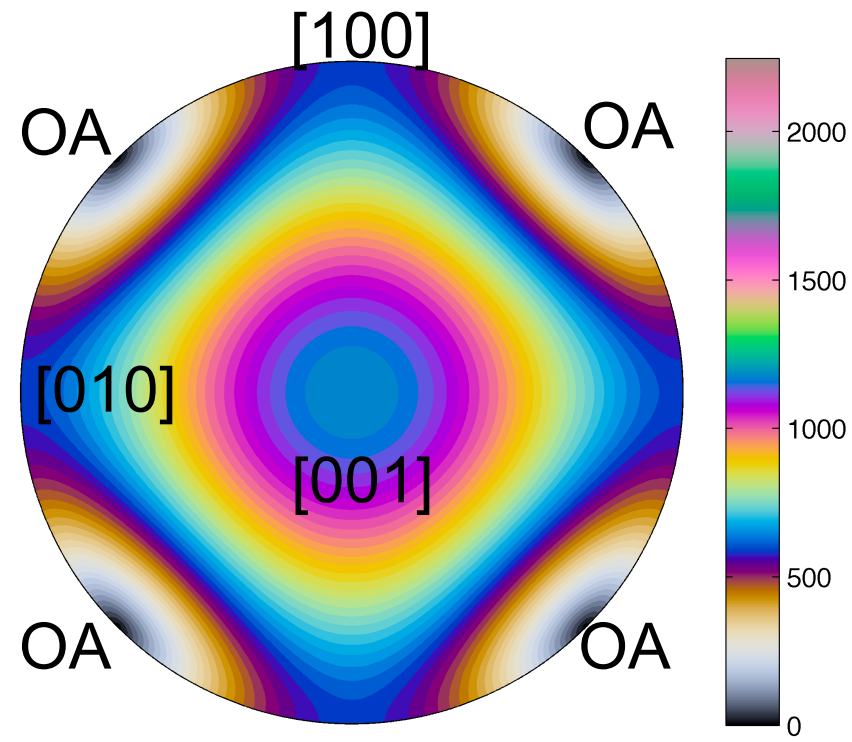
Guide To Thin Section Microscopy.

e-book ISBN 978-3-00-033606-5 (PDF) – it is free !


Olivine – refractive index and retardation in MTEX



Refractive index



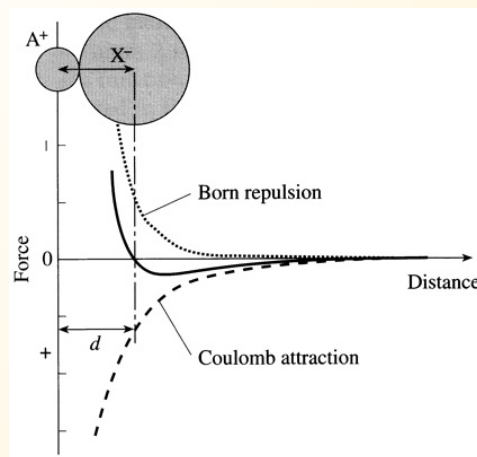
Retardation (nm): Michel Levy colours



Single crystal elasticity tensor basics –
Young's modulus to wave propagation

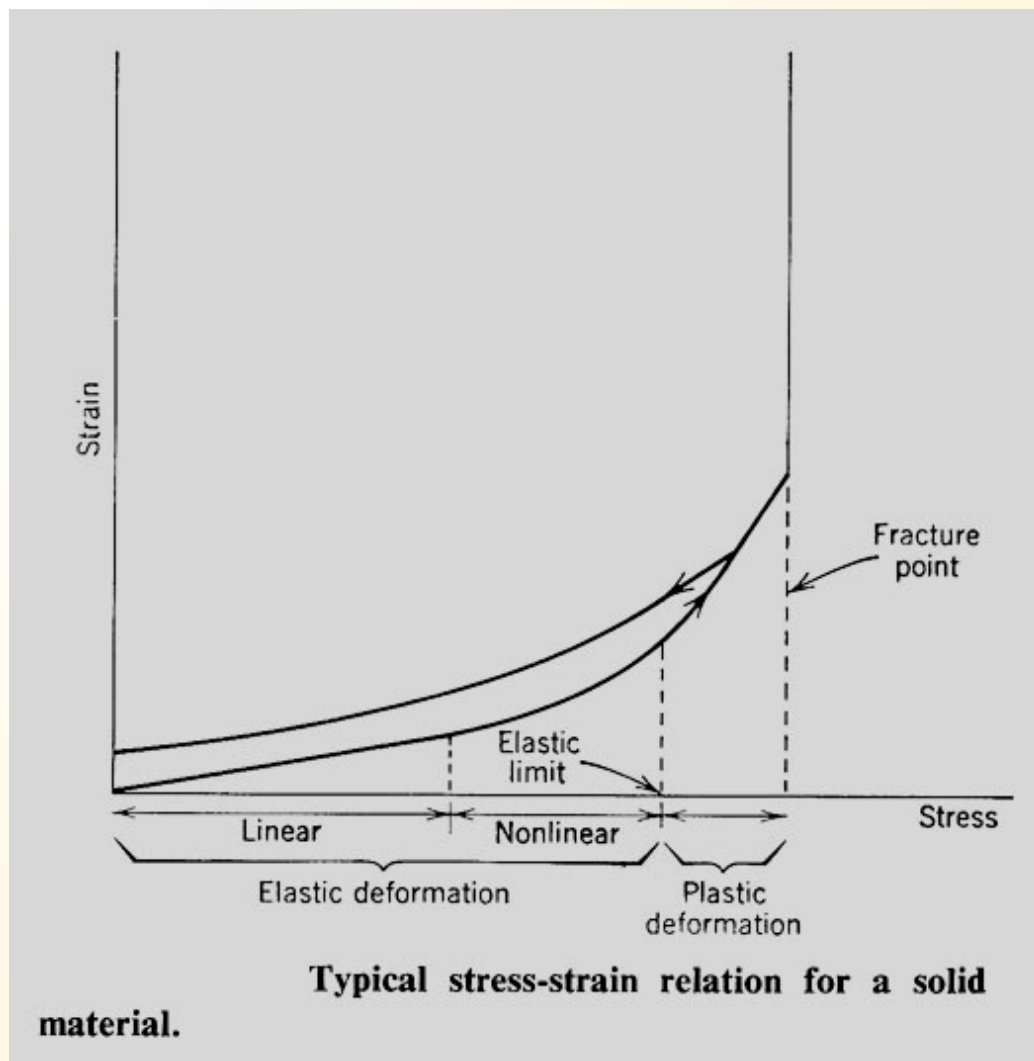
4th Rank Tensors - important for geophysics

- In any crystalline material there is balance between **Coulomb attractive** forces between oppositely charge ions and **Born repulsive** forces due to the overlap of electron shells. At any given thermodynamic state the crystal will tend toward an equilibrium structure.



- For a change in hydrostatic or non-hydrostatic stress, the crystal structure will adjust at the atomic level to the new **thermodynamic state**.
- The fundamental nature of atomic forces in the determination of elastic properties has been illustrated by the emergence of **first principles atomic modeling** to predict single crystal elastic tensors of geophysical importance at **very high PT conditions** of interior of planets.

Elastic (small linear) Strains



Hooke's Law

$$\boldsymbol{\sigma} = \mathbf{c} \boldsymbol{\varepsilon} \text{ and } \boldsymbol{\varepsilon} = \mathbf{s} \boldsymbol{\sigma}$$

where \mathbf{c} = stiffness coefficients (dimensions of stress)

\mathbf{s} = compliance coefficients (dimensions of 1/stress)

$\boldsymbol{\sigma}$ = stress tensor (second order symmetric tensor)

$\boldsymbol{\varepsilon}$ = deformation tensor (second order symmetric tensor)

$$\sigma_{ij} = C_{ijkl} \varepsilon_{kl}$$

or

$$\varepsilon_{ij} = S_{ijkl} \sigma_{kl}$$

i, j, k, l can have the values 1, 2 or 3

so $3 \times 3 \times 3 \times 3 = 3^4 = 81$ coefficients.

But due to the symmetry of the deformation and stress tensors the 81 coefficients are not independent. In addition thermodynamic considerations of the crystal energy also reduce the number of independent coefficients. In Voigt notation we can write the C_{ijkl} tensor as 6 by 6 symmetric tensor C_{ij} with 21 independent values for a triclinic crystal.

81 elastic constants 9 x 9

- If stress and strain where NOT symmetric
- 9 values of stress and strain

$$\begin{bmatrix} \sigma_{11} \\ \sigma_{22} \\ \sigma_{33} \\ \sigma_{23} \\ \sigma_{31} \\ \sigma_{12} \\ \sigma_{32} \\ \sigma_{13} \\ \sigma_{21} \end{bmatrix} = \begin{bmatrix} c_{11} & c_{12} & c_{13} & c_{14} & c_{15} & c_{16} & c_{14} & c_{15} & c_{16} \\ c_{12} & c_{22} & c_{23} & c_{24} & c_{25} & c_{26} & c_{24} & c_{25} & c_{26} \\ c_{13} & c_{23} & c_{33} & c_{34} & c_{35} & c_{36} & c_{34} & c_{35} & c_{36} \\ c_{14} & c_{24} & c_{34} & c_{44} & c_{45} & c_{46} & c_{44} & c_{45} & c_{46} \\ c_{15} & c_{25} & c_{35} & c_{45} & c_{55} & c_{56} & c_{45} & c_{55} & c_{56} \\ c_{16} & c_{26} & c_{36} & c_{46} & c_{56} & c_{66} & c_{46} & c_{56} & c_{66} \\ c_{14} & c_{24} & c_{34} & c_{44} & c_{45} & c_{46} & c_{44} & c_{45} & c_{46} \\ c_{15} & c_{25} & c_{35} & c_{45} & c_{55} & c_{56} & c_{45} & c_{55} & c_{56} \\ c_{16} & c_{26} & c_{36} & c_{46} & c_{56} & c_{66} & c_{46} & c_{56} & c_{66} \end{bmatrix} \begin{bmatrix} \epsilon_{11} \\ \epsilon_{22} \\ \epsilon_{33} \\ \epsilon_{23} \\ \epsilon_{31} \\ \epsilon_{12} \\ \epsilon_{32} \\ \epsilon_{13} \\ \epsilon_{21} \end{bmatrix}$$

Stress and Strain Tensors

- Both symmetric for small linear elastic strains

$$\begin{bmatrix} \sigma_{11} & \sigma_{12} & \sigma_{31} \\ \sigma_{12} & \sigma_{22} & \sigma_{23} \\ \sigma_{31} & \sigma_{23} & \sigma_{33} \end{bmatrix}$$

$$\begin{bmatrix} \epsilon_{11} & \epsilon_{12} & \epsilon_{31} \\ \epsilon_{12} & \epsilon_{22} & \epsilon_{23} \\ \epsilon_{31} & \epsilon_{23} & \epsilon_{33} \end{bmatrix}$$

- 6 Independent values $ij = ji$
(not 9 because $ij \neq ji$)

$$\sigma_{ij} = C_{ijkl} \epsilon_{kl}$$

- $C_{ijkl} = (\partial \sigma_{ij} / \partial \epsilon_{kl})$ and $S_{ijkl} = (\partial \epsilon_{ij} / \partial \sigma_{kl})$
- Stress tensor $\sigma_{ij} = \sigma_{ji}$ hence $C_{ijkl} = C_{jikl}$
- Strain tensor $\epsilon_{kl} = \epsilon_{lk}$ hence $C_{ijkl} = C_{ijlk}$
- The elastic constants can be expressed as a function of the crystal energy per unit volume of crystal (U) as

$$C_{ijkl} = (\partial^2 U / \partial \epsilon_{ij} \partial \epsilon_{kl}) \text{ where the strains } \partial \epsilon_{ij} \text{ and } \partial \epsilon_{kl} \text{ can be inter-changed, hence } C_{ijkl} = C_{klij}$$

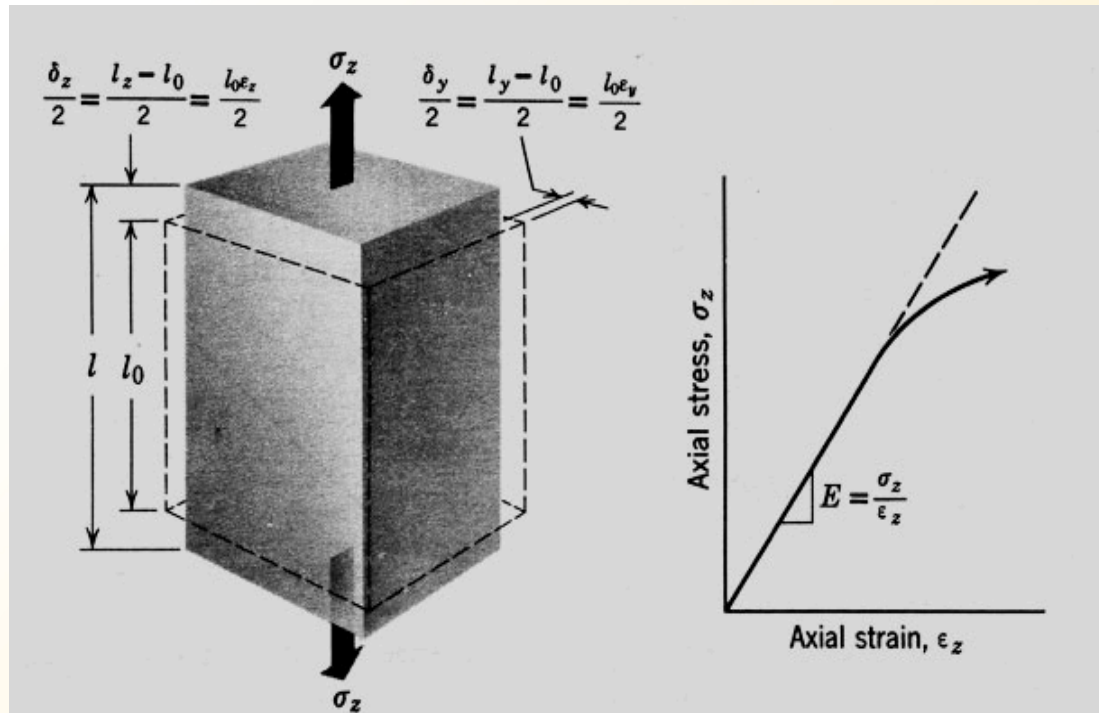
Life was not meant to be easy in elasticity !

21 independent elastic constants 6 x 6 symmetric tensor

- Stress and strain tensor are symmetric
- 6 values of stress and strain
- The strains $\partial \epsilon_{ij}$ and $\partial \epsilon_{kl}$ are inter-changable

$$\begin{bmatrix} \sigma_{11} \\ \sigma_{22} \\ \sigma_{33} \\ \sigma_{23} \\ \sigma_{31} \\ \sigma_{12} \end{bmatrix} = \begin{bmatrix} c_{11} & c_{12} & c_{13} & c_{14} & c_{15} & c_{16} \\ c_{12} & c_{22} & c_{23} & c_{24} & c_{25} & c_{26} \\ c_{13} & c_{23} & c_{33} & c_{34} & c_{35} & c_{36} \\ c_{14} & c_{24} & c_{34} & c_{44} & c_{45} & c_{46} \\ c_{15} & c_{25} & c_{35} & c_{45} & c_{55} & c_{56} \\ c_{16} & c_{26} & c_{36} & c_{46} & c_{56} & c_{66} \end{bmatrix} \begin{bmatrix} \epsilon_{11} \\ \epsilon_{22} \\ \epsilon_{33} \\ \gamma_{23} \\ \gamma_{31} \\ \gamma_{12} \end{bmatrix}$$

Young's modulus & Poisson's ratio



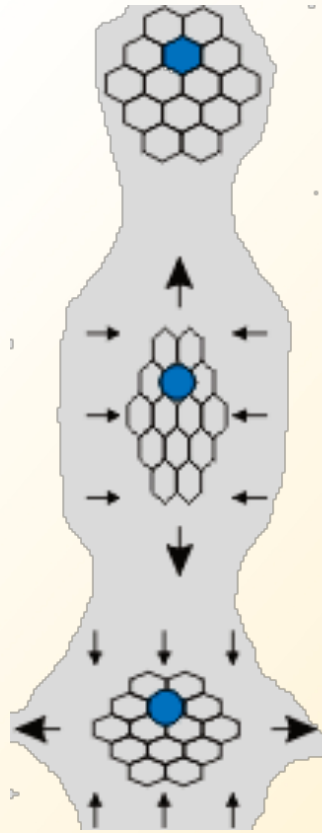
Uniaxial tensile (or compressive) stress. Poisson's ratio, ν , is the ratio of transverse to axial strain. Dashed lines represent initial stress-free shape: a cube of edge length l_0 .

Poisson's ratio, ν , another elastic constant, is the ratio of transverse to axial strain

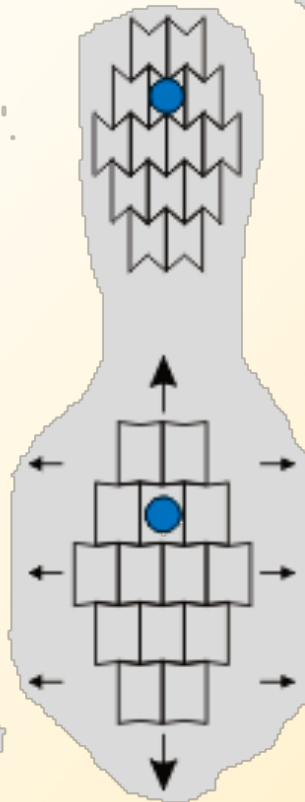
$$\nu = \frac{-\epsilon_y}{\epsilon_z}$$

Auxetic behaviour – negative Poisson's ratio

Poisson's ratio +ve



Poisson's ratio -ve



The case of cork ?



At low strain the Poisson's ratio is 0.2, nearly zero – ideal to pushing into the bottle, or even better drawing it out ! At high strains, beware the Poisson's ratio can become negative -0.01 and the cork is impossible to pull out without braking it !

The Voigt Notation 1

Voigt

$p, q = 1$ to 6

Tensors

$i, j, k, l = 1$ to 3

Matrix (p,q)	1	2	3	4	5	6
Tensor (ij or kl)	11	22	33	23 or 32	31 or 13	12 or 21

$$\sigma_{ij} = \begin{bmatrix} \sigma_{11} & \sigma_{12} & \sigma_{13} \\ \sigma_{12} & \sigma_{22} & \sigma_{23} \\ \sigma_{13} & \sigma_{23} & \sigma_{33} \end{bmatrix} \longrightarrow \sigma_i = \begin{bmatrix} \sigma_1 & \sigma_6 & \sigma_5 \\ \sigma_6 & \sigma_2 & \sigma_4 \\ \sigma_5 & \sigma_4 & \sigma_3 \end{bmatrix}$$

$$\epsilon_{ij} = \begin{bmatrix} \epsilon_{11} & \epsilon_{12} & \epsilon_{13} \\ \epsilon_{21} & \epsilon_{22} & \epsilon_{23} \\ \epsilon_{31} & \epsilon_{32} & \epsilon_{33} \end{bmatrix} \longrightarrow \epsilon_i = \begin{bmatrix} \epsilon_1 & \epsilon_{6/2} & \epsilon_{5/2} \\ \epsilon_{6/2} & \epsilon_2 & \epsilon_{4/2} \\ \epsilon_{5/2} & \epsilon_{4/2} & \epsilon_3 \end{bmatrix}$$

Life was not meant to be easy in elasticity !

The Voigt Notation 2

In addition we need to introduce factors of 2 and 4 into the equations relating compliance in tensor and matrix notations due the factor of 2 due to shear strain tensor $\varepsilon \rightarrow$ shear strain matrix $\varepsilon/2$.

$$S_{ijkl} = S_{pq} \quad \text{for } p=1,2,3 \text{ and } q = 1,2,3$$

$$2S_{ijkl} = S_{pq} \quad \text{for either } p= \text{ or } q = 4,5,6$$

$$4S_{ijkl} = S_{pq} \quad \text{for either } p= \text{ and } q = 4,5,6$$

However

$$C_{ijkl} = C_{pq}$$

for all i,j,k,l and all p and q (for $i,j,k,l = 1,2,3$ and $p,q = 1$ to 6)

Transformation law

$$p = \delta_{ij}.i + (1-\delta_{ij})(9-i-j)$$

$$q = \delta_{kl}.k + (1-\delta_{kl})(9-k-l)$$

with the Kronecker delta , $\delta_{ij} = 0$ when $i \neq j$ and $\delta_{ij} = 1$ when $i=j$

Life was not meant to be easy in elasticity !

Crystal symmetry classes and their Elastic tensors

Cubic (3) 23, $m\bar{3}$, 432, $\bar{4}3m$, $m\bar{3}m$

$$\begin{bmatrix} c_{11} & c_{12} & c_{12} & 0 & 0 & 0 \\ c_{12} & c_{11} & c_{12} & 0 & 0 & 0 \\ c_{12} & c_{12} & c_{11} & 0 & 0 & 0 \\ 0 & 0 & 0 & c_{44} & 0 & 0 \\ 0 & 0 & 0 & 0 & c_{44} & 0 \\ 0 & 0 & 0 & 0 & 0 & c_{44} \end{bmatrix}$$

Hexagonal (5) 6, $\bar{6}$, 6/m, 622, 6mm, $\bar{6}2m$, 6/mmm

$$\begin{bmatrix} c_{11} & c_{12} & c_{13} & 0 & 0 & 0 \\ c_{12} & c_{11} & c_{13} & 0 & 0 & 0 \\ c_{13} & c_{13} & c_{33} & 0 & 0 & 0 \\ 0 & 0 & 0 & c_{44} & 0 & 0 \\ 0 & 0 & 0 & 0 & c_{44} & 0 \\ 0 & 0 & 0 & 0 & 0 & \frac{1}{2}(c_{11}-c_{12}) \end{bmatrix}$$

Orthorhombic (9) 222, $mm2$, mmm

$$\begin{bmatrix} c_{11} & c_{12} & c_{13} & 0 & 0 & 0 \\ c_{12} & c_{22} & c_{23} & 0 & 0 & 0 \\ c_{13} & c_{23} & c_{33} & 0 & 0 & 0 \\ 0 & 0 & 0 & c_{44} & 0 & 0 \\ 0 & 0 & 0 & 0 & c_{55} & 0 \\ 0 & 0 & 0 & 0 & 0 & c_{66} \end{bmatrix}$$

Monoclinic (13) 2, m , $2/m$

$$\begin{bmatrix} c_{11} & c_{12} & c_{13} & 0 & c_{15} & 0 \\ c_{12} & c_{22} & c_{23} & 0 & c_{25} & 0 \\ c_{13} & c_{23} & c_{33} & 0 & c_{35} & 0 \\ 0 & 0 & 0 & c_{44} & 0 & c_{46} \\ c_{14} & c_{25} & c_{35} & 0 & c_{55} & 0 \\ 0 & 0 & 0 & c_{46} & 0 & c_{66} \end{bmatrix}$$

Trigonal (6) 32, $3m$, $\bar{3}m$

$$\begin{bmatrix} c_{11} & c_{12} & c_{13} & c_{14} & 0 & 0 \\ c_{12} & c_{11} & c_{13} & -c_{14} & 0 & 0 \\ c_{13} & c_{13} & c_{33} & 0 & 0 & 0 \\ c_{14} & -c_{14} & 0 & c_{44} & 0 & 0 \\ 0 & 0 & 0 & 0 & c_{44} & c_{14} \\ 0 & 0 & 0 & 0 & c_{14} & \frac{1}{2}(c_{11}-c_{12}) \end{bmatrix}$$

Triclinic (21) 1, $\bar{1}$

$$\begin{bmatrix} c_{11} & c_{12} & c_{13} & c_{14} & c_{15} & c_{16} \\ c_{12} & c_{22} & c_{23} & c_{24} & c_{25} & c_{26} \\ c_{13} & c_{23} & c_{33} & c_{34} & c_{35} & c_{36} \\ c_{14} & c_{24} & c_{34} & c_{44} & c_{45} & c_{46} \\ c_{15} & c_{25} & c_{35} & c_{45} & c_{55} & c_{56} \\ c_{16} & c_{26} & c_{36} & c_{46} & c_{56} & c_{66} \end{bmatrix}$$

Trigonal (7) 3, $\bar{3}$

$$\begin{bmatrix} c_{11} & c_{12} & c_{13} & c_{14} & -c_{25} & 0 \\ c_{12} & c_{11} & c_{13} & -c_{14} & c_{25} & 0 \\ c_{13} & c_{13} & c_{33} & 0 & 0 & 0 \\ c_{14} & -c_{14} & 0 & c_{44} & 0 & c_{25} \\ -c_{25} & c_{25} & 0 & 0 & c_{44} & c_{14} \\ 0 & 0 & 0 & c_{25} & c_{14} & \frac{1}{2}(c_{11}-c_{12}) \end{bmatrix}$$

Tetragonal (6) 422, $4mm$, $\bar{4}2m$, $4/mmm$

$$\begin{bmatrix} c_{11} & c_{12} & c_{13} & 0 & 0 & 0 \\ c_{12} & c_{11} & c_{13} & 0 & 0 & 0 \\ c_{13} & c_{13} & c_{33} & 0 & 0 & 0 \\ 0 & 0 & 0 & c_{44} & 0 & 0 \\ 0 & 0 & 0 & 0 & c_{44} & 0 \\ 0 & 0 & 0 & 0 & 0 & c_{66} \end{bmatrix}$$

Tetragonal (7), $4, \bar{4}$ 4/m

$$\begin{bmatrix} c_{11} & c_{12} & c_{13} & 0 & 0 & c_{16} \\ c_{12} & c_{11} & c_{13} & 0 & 0 & -c_{16} \\ c_{13} & c_{13} & c_{33} & 0 & 0 & 0 \\ 0 & 0 & 0 & c_{44} & 0 & 0 \\ 0 & 0 & 0 & 0 & c_{44} & 0 \\ c_{16} & -c_{16} & 0 & 0 & 0 & c_{66} \end{bmatrix}$$

Orthorhombic Crystals

Elastic Stiffness Cij

J.D. Bass (1995) (Mg,Fe)₂SiO₄

Table 1. Elastic Constants of Orthorhombic Crystals at Room P & T

Material	ρ Mg/m ³	Subscript <i>ij</i> in modulus <i>c_{ij}</i> (GPa)									<i>K_s</i> GPa	<i>G</i> GPa	References
		11	22	33	44	55	66	12	13	23			
<i>Perovskites</i>													
MgSiO ₃	4.108	515	525	435	179	202	175	117	117	139	146.4	184.2	149
NaMgF ₃	3.058	125.7	147.3	142.5	46.7	44.8	50.4	49.5	45.1	43.1	75.7	46.7	155
<i>Pyroxenes</i>													
Enstatite (En ₁₀₀), MgSiO ₃	3.198	224.7	177.9	213.6	77.6	75.9	81.6	72.4	54.1	52.7	107.8	75.7	142
Ferrosilite (Fs ₁₀₀), FeSiO ₃	4.002	198	136	175	59	58	49	84	72	55	101	52	9
En ₉₄ Fs ₆	3.272	229.3	167.0	193.9	79.7	76.1	77.1	73.6	49.8	46.6	102.3	73.9	31
En _{84.5} Fs _{15.2}	3.335	229.9	165.4	205.7	83.1	76.4	78.5	70.1	57.3	49.6	105.0	75.5	64
En ₈₀ Fs ₂₀	3.354	228.6	160.5	210.4	81.8	75.5	77.7	71.0	54.8	46.0	103.5	74.9	31
	3.373	231.0	169.8	215.7	82.8	76.5	78.1	78.9	61.4	49.1	109.4	75.2	137
Protoenstatite, MgSiO ₃	3.052	213	152	246	81	44	67	76	59	70	112	63	123
<i>Olivines</i>													
Forsterite (Fo ₁₀₀), Mg ₂ SiO ₄	3.221	328	200	235	66.7	81.3	80.9	69	69	73	129.5	81.1	46
Fayalite (Fa ₁₀₀), Fe ₂ SiO ₄	3.818	266	168	232	32.3	46.5	57	94	92	92	134	50.7	55
Fo ₉₁ Fa ₉	3.25	320.2	195.9	233.8	63.5	76.9	78.1	67.9	70.5	78.5	129.5	77.6	136
Fo ₉₃ Fa ₇	3.311	323.7	197.6	235.1	64.6	78.1	79.0	66.4	71.6	75.6	129.4	79.1	65
Fo _{91.3} Fa _{8.1}	3.316	324	196	232	63.9	77.9	78.8	71.5	71.5	68.8	128.1	78.7	82
Fo ₉₂ Fa ₈	3.299	319	192	238	63.8	78.3	79.7	59	76	72	126.7	79.0	82
Mn ₂ SiO ₄ ,	4.129	258.4	165.6	206.8	45.3	55.6	57.8	87	95	92	128	54	117
Monticellite, CaMgSiO ₄	3.116	216	150	184	50.6	56.5	59.2	59	71	77	106	55.2	92
Ni ₂ SiO ₄	4.933	340	238	253	71	87	78	109	110	113	165	80	13
Co ₂ SiO ₄	4.706	307.8	194.7	234.2	46.7	63.9	64.8	102	105	103	148	62	117
Mg ₂ GeO ₄	4.029	312	187	217	57.2	66.1	71	60	65	66	120	72	140

Fo93

AGU Bookshelf Reference Books (free !)

<http://www.agu.org/reference/minphys.html>



Defining a single crystal elastic tensor in MTEX

1. cs_Tensor with frame 'x||a' & 'z||c'
2. Tensor in Voigt matrix form M
3. C = tensor(M,cs_tensor...)

Editor Window

```

*****
% Define elastic stiffness tensor (GPa)
*****

% Reference Elastic constants
% Kinazawa, M., Anderson, O.L., 1969. Elastic moduli,
% pressure derivatives, and temperature derivatives of
% single-crystal olivine and single-crystal forsterite.
% J. Geophys. Res., 74, 5961-5972.

% Reference Crystal Structure
% San Carlos OLIVINE Fo93 room pressure d=3.311 g/cm3

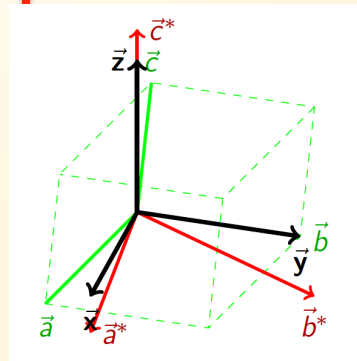
% Define density (g/cm3)
rho = 3.311;

% Define Cartesian tensor crystal symmetry
% tensor frame X and Z
cs_Tensor = symmetry('mmm',[ 4.7646 10.2296 5.9942],...
[ 90.0000 90.0000 90.0000]*degree,'X||a','Z||c',...
'mineral','Olivine Fo93');

% Elastic Cij stiffness tensor (GPa) as matrix M
M = ...
[[ 323.70 66.40 71.60 0.00 0.00 0.00];...
[ 66.40 197.60 75.60 0.00 0.00 0.00];...
[ 71.60 75.60 235.10 0.00 0.00 0.00];...
[ 0.00 0.00 0.00 64.60 0.00 0.00];...
[ 0.00 0.00 0.00 0.00 78.10 0.00];...
[ 0.00 0.00 0.00 0.00 0.00 79.00]];

% M as stiffness tensor C with MTEX tensor command
C = tensor(M,cs_Tensor,'rank',4,'propertyname',...
'Olivine Fo93 Stiffness tensor 1969','unit','GPa')

```



Cartesian Tensor frame **x,y,z**

Command Window

```

C = tensor (show methods, plot)
propertyname: Olivine Fo93 Stiffness tensor 1969
unit        : GPa
rank        : 4 (3 x 3 x 3 x 3)
mineral     : Olivine Fo93 (mmm)

tensor in Voigt matrix representation:
323.7  66.4  71.6   0   0   0
66.4  197.6  75.6   0   0   0
71.6  75.6  235.1   0   0   0
0      0      0  64.6   0   0
0      0      0   0  78.1   0
0      0      0   0   0  79

```

Single crystal velocity calculation

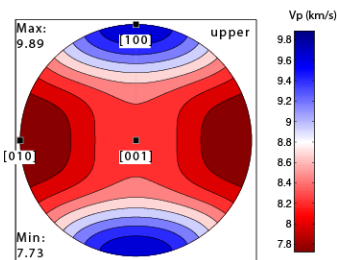
Solve Christoffel symmetric tensor (3 by 3)
at each grid point n

$$T_{ik}(n) = C_{ijkl}n_jn_l$$

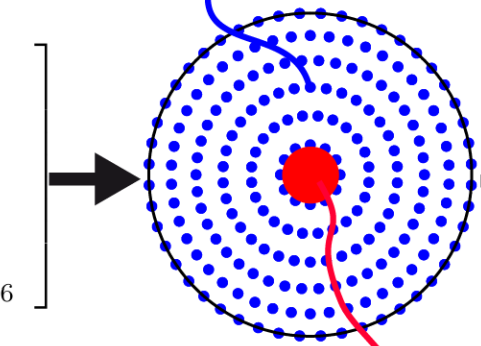
$$V_p = \sqrt{\frac{\lambda_1}{\rho}}, \quad V_{s1} = \sqrt{\frac{\lambda_2}{\rho}}, \quad V_{s2} = \sqrt{\frac{\lambda_3}{\rho}}$$

Single Crystal Cij for Olivine
(orthorhombic)

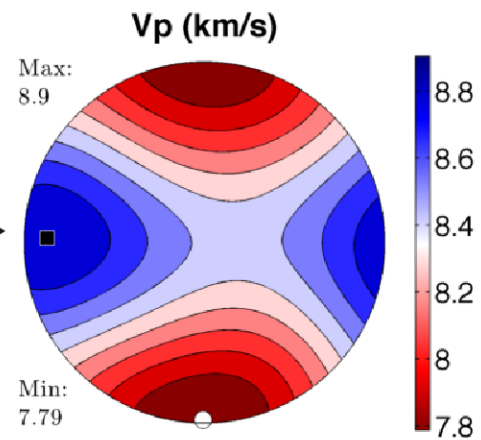
C_{11}	C_{12}	C_{13}	0	0	0
C_{12}	C_{22}	C_{23}	0	0	0
C_{13}	C_{23}	C_{33}	0	0	0
0	0	0	C_{44}	0	0
0	0	0	0	C_{55}	0
0	0	0	0	0	C_{66}



crystal coordinates



pcrystal coordinates



```
plot(C_Crystal,'density',rho_Crystal,
'PlotType','velocity','vp','complete','contour')
```

Plot and rotate a tensor of any rank

```

%% Plot elastic tensor property Vp in km/s with density rho
plot(C,'density',rho,'PlotType','velocity','vp','complete','contourf'
')
annotate([xvector,yvector,zvector],'label',{'[100] ','[010] ','[001]
'}...
,'BackgroundColor','w');
colorbar
savefigure('/MatLab_Programs/Plot_Olivine_Single_Crystal_Vp_Label.pdf')
%
%% rotate and plot
r = rotation('Euler',45*degree,30*degree,60*degree)
% rotate elastic tensor
C_rotated = rotate(C,r)
% rotate specimen frame vectors
x_rotated = rotate(xvector,r)
y_rotated = rotate(yvector,r)
z_rotated = rotate(zvector,r)
% plot tensor
plot(C_rotated,'density',rho,'PlotType','velocity','vp',...
'complete','contourf')
annotate([x_rotated,y_rotated,z_rotated],'label',...
{'[100] ','[010] ','[001] '},'BackgroundColor','w');
colorbar
savefigure(...
'/MatLab_Programs/Plot_Olivine_Single_Crystal_Vp_rotated.pdf');

```

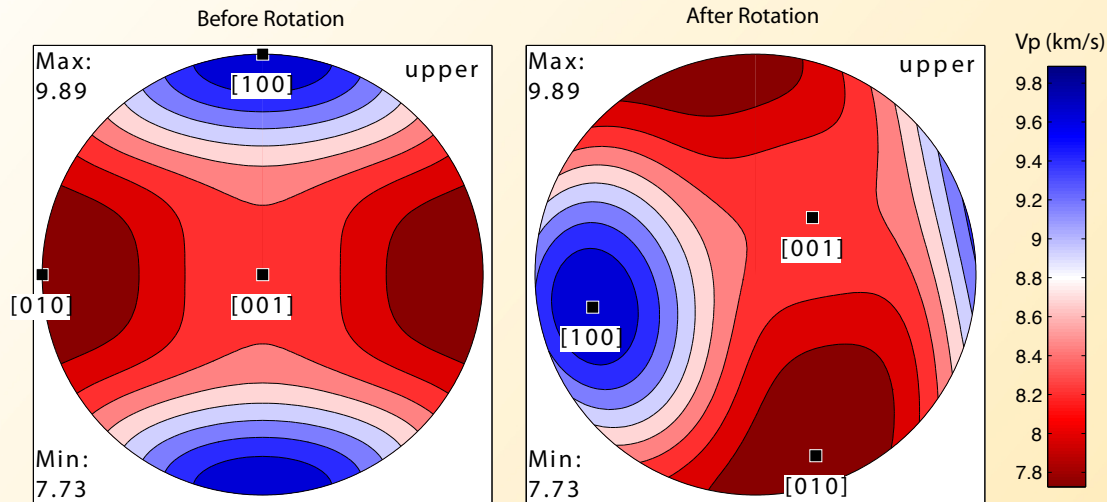
Rotate a 4th rank tensor

$$C'_{ijkl} = r_{im}r_{jn}r_{ko}r_{lp} C_{mnop}$$

```

Plot(C,'density',rho,'PlotType','velocity','vp','complete','contourf')
r=rotation('Euler',45*degree,30*degree,60*degree)
C_rotated = rotate(C,r)      x_rotated = rotate(xvector,r)

```



In **M**TEX the reference frames are defined

The command to define the crystal reference of the **Euler frame** is

```
CS = crystalSymmetry('triclinic',[5.2957,9.1810,9.4228]...  
,[90.372,98.880,90.110]*degree,'xlla*','zllc','mineral','Talc')
```

The command to define the crystal reference of the **Tensor frame** is

```
CS_tensor = crystalSymmetry('triclinic',[5.2957,9.1810,9.4228]...  
,[90.372,98.880,90.110]*degree,'xlla','zllc*','mineral','Talc')
```

N.B. in **Euler frame** 'xlla*', 'zllc' and in the **Tensor frame** 'xlla', 'zllc*'

MTEX knows the orientation of each reference frame and automatically applies all corrections for the orientation difference between **Euler** and **Tensor** frames.

What can we do with elastic constants ?

MTEX commands

- Volume compressibility
- Linear compressibility
- Young's modulus
- Shear modulus
- Poission's ratio
- Christoffel Tensor
- Etc...

```
beta = volumeCompressibility (C)
beta = linearCompressibility (C,x)
E = YoungsModulus (C,x)
G = shearModulus (C,h,u)
nu = PoissonRatio (C,x,y)
T = ChristoffelTensor (C,n)
```

Life was not meant to be easy in elasticity! **But there is MTEX!**

Define Elastic tensor in MTEX

Editor window m-file

```

%%
%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
% Specify Crystal Symmetry (cs)
%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
%
% crystal symmetry - Orthorhombic mmm
% Olivine structure
% (4.7646 10.2296 5.9942 90.00 90.00 90.00) - Orthorhombic
cs_tensor = symmetry('mmm',[4.7646,10.2296,5.9942],...
    [90.00,90.00,90.00]*degree,'xIIa','zIIc','mineral','Olivine');
%
%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
% Import 4th rank tensor as 6 by 6 matrix
%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
%
% Olivine elastic stiffness (Cij) Tensor in GPa
% Abramson E.H., Brown J.M., Slutsky L.J., and Zaugg J.(1997)
% The elastic constants of San Carlos olivine to 17 GPa.
% Journal of Geophysical Research 102: 12253-12263.
%
%   320.5 68.15 71.6    0    0    0
%   68.15 196.5 76.8    0    0    0
%   71.6  76.8 233.5    0    0    0
%   0      0   0      64   0    0
%   0      0   0      0   77   0
%   0      0   0      0   0   78.7
%
% Enter tensor as 6 by 6 matrix,M line by line.
%
M = [[320.5 68.15 71.6    0    0    0];...
    [ 68.15 196.5 76.8    0    0    0];...
    [ 71.6  76.8 233.5    0    0    0];...
    [  0      0   0      64   0    0];...
    [  0      0   0      0   77   0];...
    [  0      0   0      0   0   78.7]];
%
%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
% Define tenor object in MTEX
% Cij -> Cijkl - elastic stiffness tensor
%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
%
C = tensor(M,cs_tensor);
%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
% Define density (g/cm3)
%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
rho=3.355;
%

```

Command window

```

cs_tensor = crystal symmetry (show methods, plot)
mineral : Olivine
symmetry: mmm (mmm)
a, b, c : 4.8, 10, 6

C = tensor (show methods, plot)
rank    : 4 (3 x 3 x 3 x 3)
mineral: Olivine (mmm)

tensor in Voigt matrix representation:
320.5  68.2  71.6    0    0    0
68.2  196.5  76.8    0    0    0
71.6  76.8  233.5    0    0    0
0      0    0      64   0    0
0      0    0      0   77   0
0      0    0      0   0   78.7

rho =
3.3550

```

MTEX: calculate Young's modulus in a specific crystallographic direction

```
%  
% Calculate Young's Modulus in directions [100],[010],[001],[111] and (111)  
%  
% Define crystallographic directions in reference frame of "cs_tensor"  
uvw_100 = Miller(1,0,0,cs_tensor,'uvw');  
uvw_010 = Miller(0,1,0,cs_tensor,'uvw');  
uvw_001 = Miller(0,0,1,cs_tensor,'uvw');  
uvw_111 = Miller(1,1,1,cs_tensor,'uvw');  
hkl_111 = Miller(1,1,1,cs_tensor,'hkl');  
% Calculate Young's modulus (GPa)  
E_100 = YoungsModulus(C,uvw_100);  
E_010 = YoungsModulus(C,uvw_010);  
E_001 = YoungsModulus(C,uvw_001);  
E_111_uvw = YoungsModulus(C,uvw_111);  
E_111_hkl = YoungsModulus(C,hkl_111);  
% Calculate angle between [111] and (111) in degrees  
theta = angle(uvw_111,hkl_111)/degree;
```


MTEX: calculate Young's modulus in a specific crystallographic direction

```
uvw_100 = Miller (show methods, plot)
size: 1 x 1
options: uvw
mineral: Olivine (mmm)
u 1
v 0
w 0
```

```
uvw_010 = Miller (show methods, plot)
size: 1 x 1
options: uvw
mineral: Olivine (mmm)
u 0
v 1
w 0
```

```
uvw_001 = Miller (show methods, plot)
size: 1 x 1
options: uvw
mineral: Olivine (mmm)
u 0
v 0
w 1
```

```
uvw_111 = Miller (show methods, plot)
size: 1 x 1
options: uvw
mineral: Olivine (mmm)
u 1
v 1
w 1
```

```
hkl_111 = Miller (show methods, plot)
size: 1 x 1
mineral: Olivine (mmm)
h 1
k 1
l 1
```

```
E_100 =
286.9284
```

```
E_010 =
164.5770
```

```
E_001 =
196.6730
```

```
E_111_uvw =
170.9797
```

```
E_111_hkl =
200.2764
```

```
theta =
34.6431
```

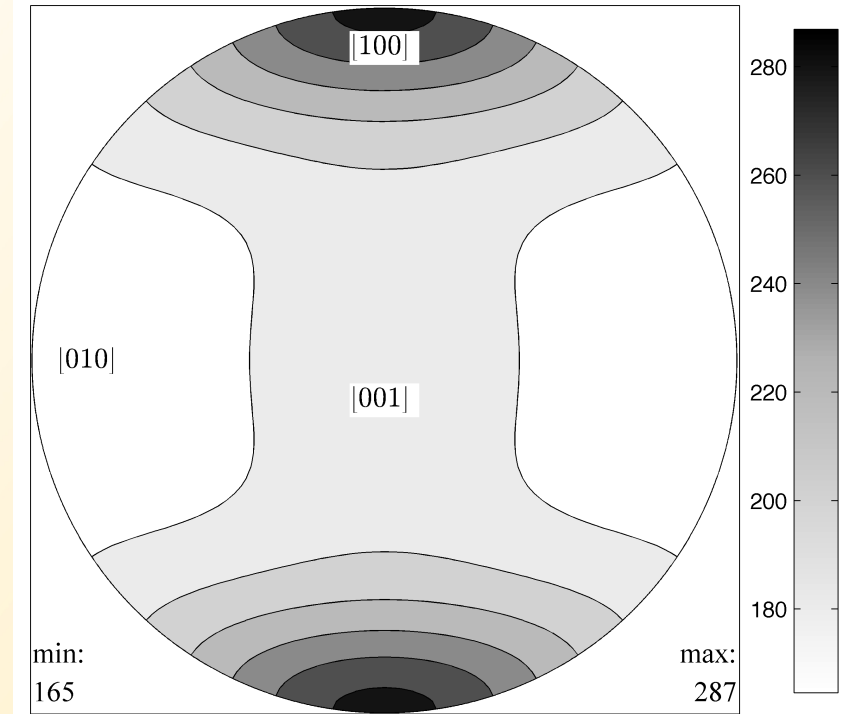
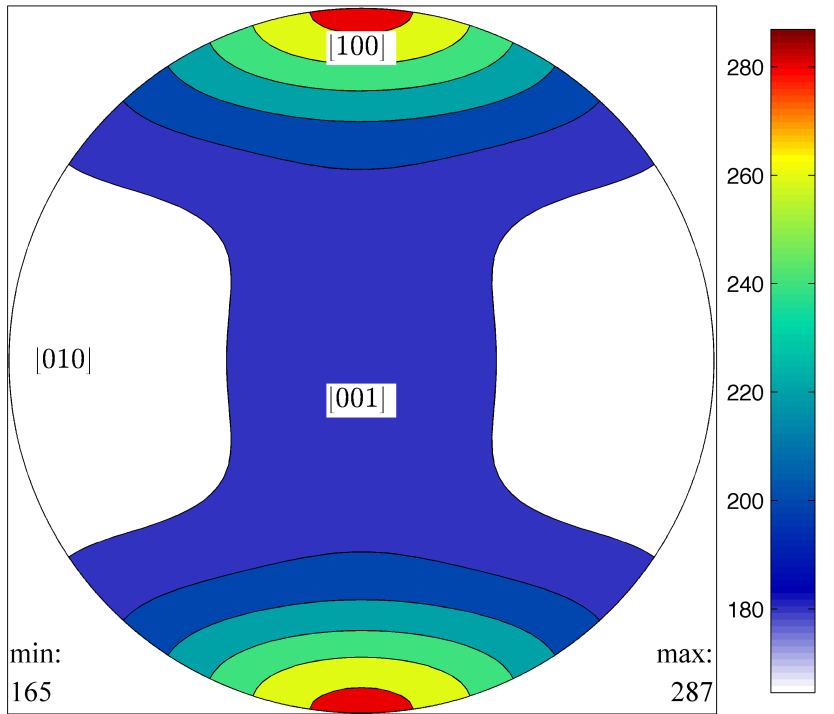

MTEX : Plotting Young's modulus

```

%
%*****
% Plotting section
%*****
%
% plotting convention
% set the default plot direction of the X-axis
%
plotx2north;
%
% plot YoungsModulus
%
setpref('mtex','defaultColorMap',WhiteJetColorMap);
plot(C,'PlotType','YoungsModulus','complete')
colorbar
hold on
plot([xvector,yvector,zvector],'data',{'[100] ','[010] ','[001] '},'backgroundcolor','w');
hold off
% save plot as *.png file
savefigure('/MatLab_Programs/Plot_Olivine_Single_Crystal_Young_colour.png');
% plot YoungsModulus
setpref('mtex','defaultColorMap',grayColorMap);
plot(C,'PlotType','YoungsModulus','complete')
colorbar
hold on
plot([xvector,yvector,zvector],'data',{'[100] ','[010] ','[001] '},'backgroundcolor','w');
hold off
%
% save plot as *.png file
savefigure('/MatLab_Programs/Plot_Olivine_Single_Crystal_Young_gray.png');

```

MTEX : Plotting Young's modulus



```
setMTEXpref('defaultColorMap',WhiteJetColorMap)
```

```
setMTEXpref('defaultColorMap',white2blackColorMap)
```

MTEX – calculate velocities (plotx2north)

```
%%
%*****
% Compute elastic wave velocities and polarizations in direction x = 1 0 0
%*****
%
% compute Vp,Vs1,Vs2,particule motion for P,S1,S2 for direction
% xvector = 1,0,0 unit vector in coordinates of X1,X2, and X3 tensor frame
[vp,vs1,vs2,pp,ps1,ps2] = velocity(C,xvector,rho);
%
```

Velocities in km/s

vp =
9.7739

vs1 =
4.8433

vs2 =
4.7907

pp = [vector3d](#) ([show methods](#), [plot](#))
size: 1 x 1
x y z
1 0 0

ps1 = [vector3d](#) ([show methods](#), [plot](#))
size: 1 x 1
x y z
0 1 0

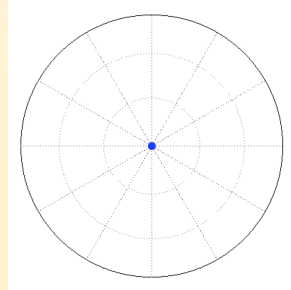
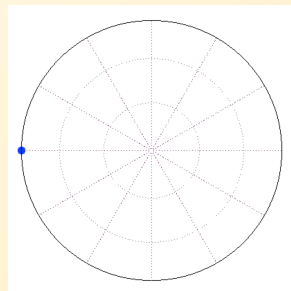
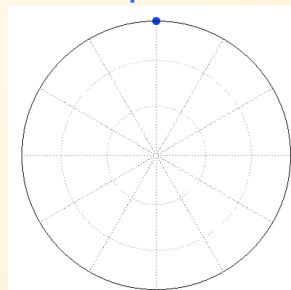
ps2 = [vector3d](#) ([show methods](#), [plot](#))
size: 1 x 1
x y z
0 0 1

pp polarization vp

ps1 polarization vs1

ps2 polarization vs2

[plot](#)



[show methods](#)

Methods for class [vector3d](#):

- | | |
|---|---|
| abs | angle |
| angle outer | char |
| check option | cross |
| ctranspose | delete option |
| display | dot |
| dot outer | double |
| end | eq |
| extract option | find |
| get | getx |
| gety | getz |
| horzcat | kernelDensityEstimation |
| length | mean |
| minus | mtimes |
| ne | norm |
| normalize | numel |
| orth | plot |
| plus | polar |
| project2FundamentalRegion | rdivide |
| repmat | reshape |
| rotate | SchmidTensor |
| set | set option |
| size | subsasgn |
| subsref | sum |
| symmetrise | times |
| transpose | uminus |
| unique | vector3d |
| vertcat | |

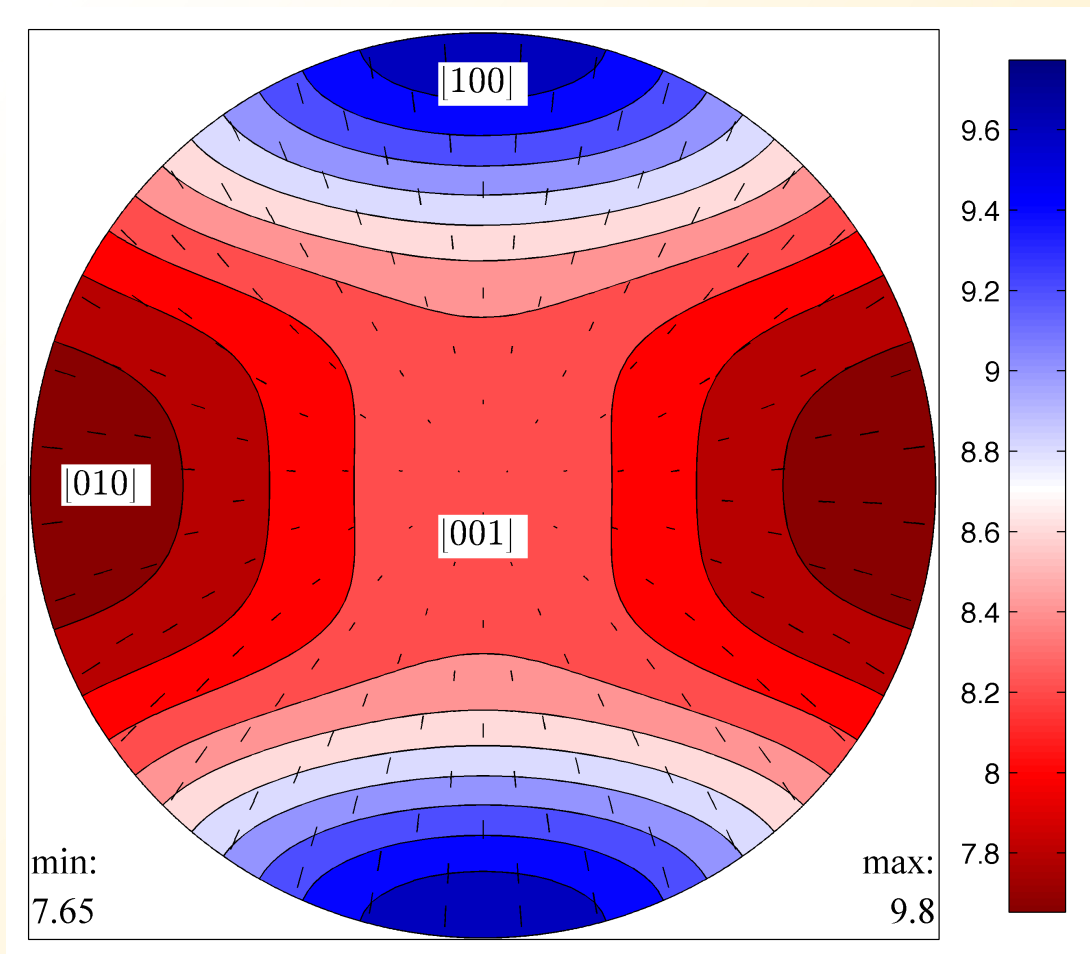
MTEX – plot velocities

```

%%
%*****
% Plotting section
%*****
%
% plotting convention
% set the default plot direction of the X-axis
%
plotx2north;
%
% set colour map to seismicColorMap
setpref('mtex','defaultColorMap',seismicColorMap);
%*****
% Plot P-wave velocity (km/s)
%*****
%
% *** need density rho in plot command for velocity ***
%
%
% Plot P-wave velocity (km/s)
%
plot(C,'density',rho,'PlotType','velocity','vp','complete','contourf')
colorbar;
%
% Add to the plot the P-wave polarization directions
%
hold on
plot(C,'density',rho,'PlotType','velocity','pp','complete');
% add tensor orthogonal axes X1,X2 and X3
hold on
plot([xvector,yvector,zvector],'data',{'[100] ','[010] ','[001] '},'backgroundcolor','w');
hold off
%
% save plot as *.png file
savefigure('/MatLab_Programs/Plot_Olivine_Single_Crystal_Vp_PP.png');

```


MTEX : Plot seismic velocities



```
setMTEXpref('defaultColorMap',blue2redColorMap)
```

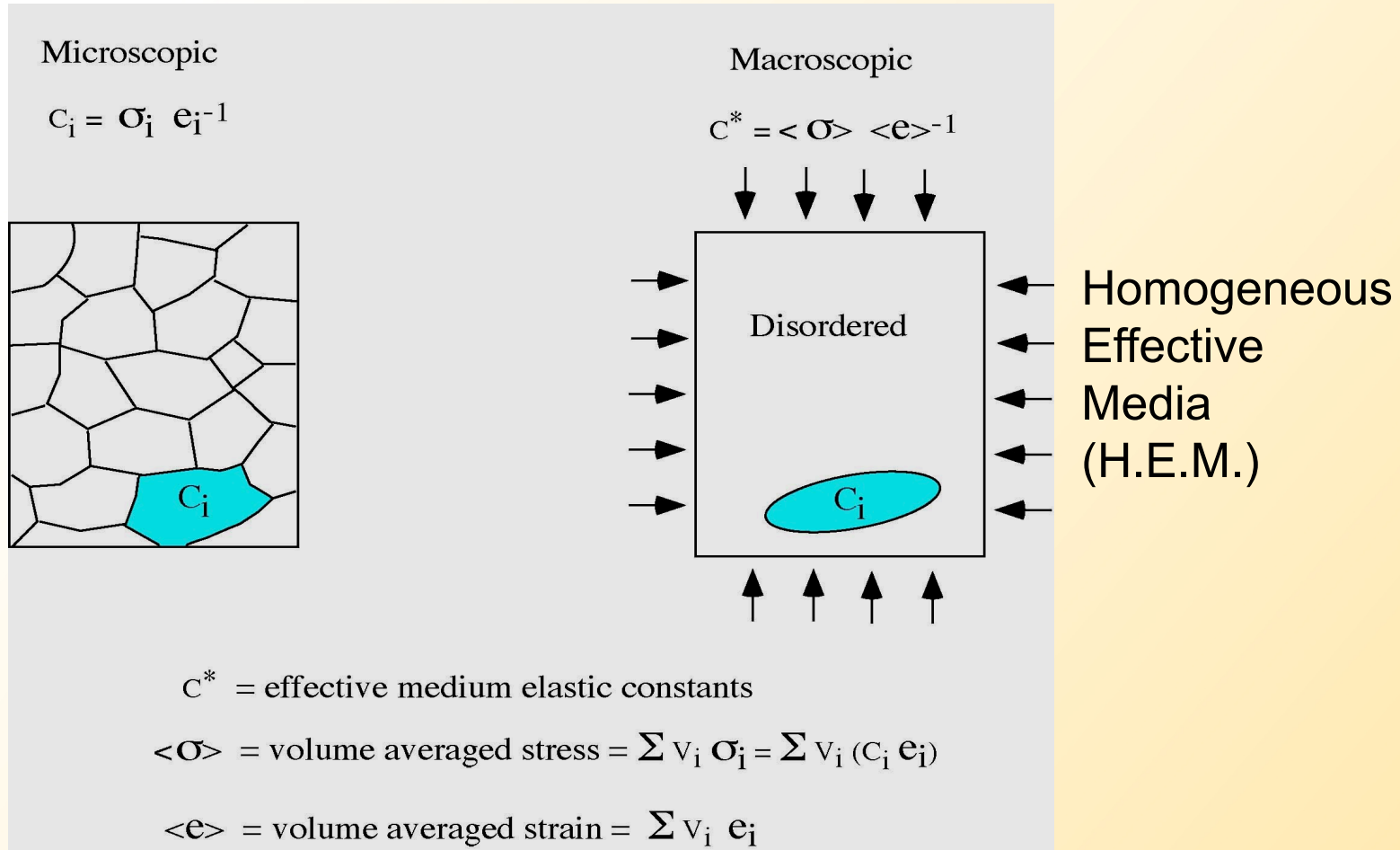
References – Tensors and Crystal Physics

- Curie, P., Oeuvres , pp. 118, Paris, Société Français de Physique (1908)
- **Voigt, W. 1928. Lehrbuch der Kristallphysik. Teubner-Verlag, Leipzig.**
- Cady, W. G. Piezoelectricity: An Introduction to the Theory and Applications of Electromechanical Phenomena in Crystals, New rev. ed., 2 vols. New York: Dover, 1964.
- Bhagavantam, S., Crystal Symmetry and Physical Properties , London, New York, Academic Press (1966)
- D.McKie and C.McKie Crystalline Solids Nelson, 1974 (Chapter 11)
- *Nye, J. F., Physical Properties of Crystals , Oxford, Clarendon Press (1957,1985 2nd Edition)*
- *Sirotnin, Yu. I. & Shakolskaya, M. P. 1982. Fundamentals of Crystal Physics. Mir, Moscow, 654.*
- Putnis, A. 1992. Introduction to Mineral Sciences, Cambridge University Press (Chapter 2)
- Wenk, H.-R. and Bulakh,A., 2004. Minerals Their Constitution and Origin, Cambridge University Press. (Chapter 8)
- **Newnham, R.E., 2005. Properties of materials Anisotropy, Symmetry, Structure. Oxford University Press,Oxford.**
- Mainprice, D. (2007) Seismic anisotropy of the deep Earth from a mineral and rock physics perspective. Schubert, G. ‘*Treatise in Geophysics*’ Volume 2 pp437-492. Oxford: Elsevier. **(I have a PDF of this chapter)**

Effective media : average
properties for polycrystalline
aggregates, Voigt, Reuss and
Hill...

2 Germans & 1 Englishman

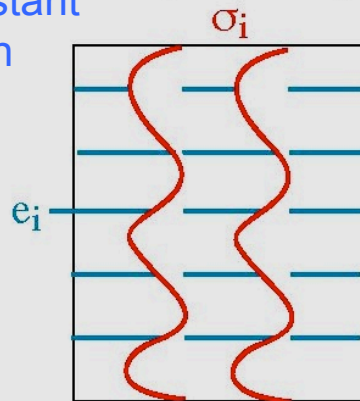
Microscopic and Macroscopic



Simple volume averages

Voigt Average

Constant strain

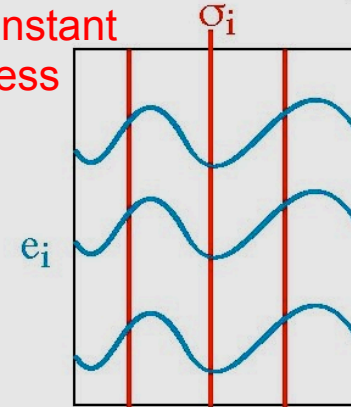


$$e_i = \text{constant: } C^* \langle \sigma \rangle = \sum V_i (C_i e_i)$$

$$C^* \approx C^{\text{Voigt}} = \sum V_i C_i$$

Reuss Average

Constant stress

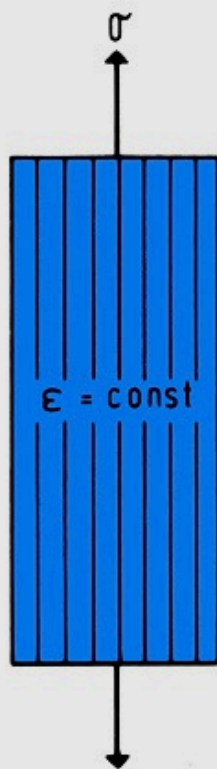


$$\sigma_i = \text{constant: } C^* \langle e \rangle^{-1} = (\sum V_i e_i)^{-1} = (\sum V_i (S_i \sigma_i))^{-1}$$

$$C^* \approx C^{\text{Reuss}} = (\sum V_i S_i)^{-1}$$

- easy to calculate
- widely separated bounds for strongly anisotropic minerals
- cannot introduce microstructure e.g. shape
- very poor bounds for mixtures with very different stiffnesses e.g. solids, liquids and voids

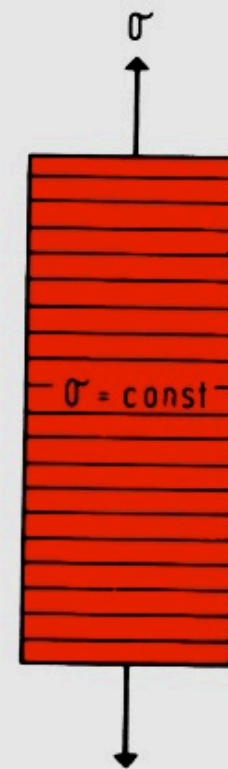
Voigt & Reuss micro-structural models



Voigt (a)

Voigt-Reuss-Hill (VRH) or Hill = $(V+R)/2$

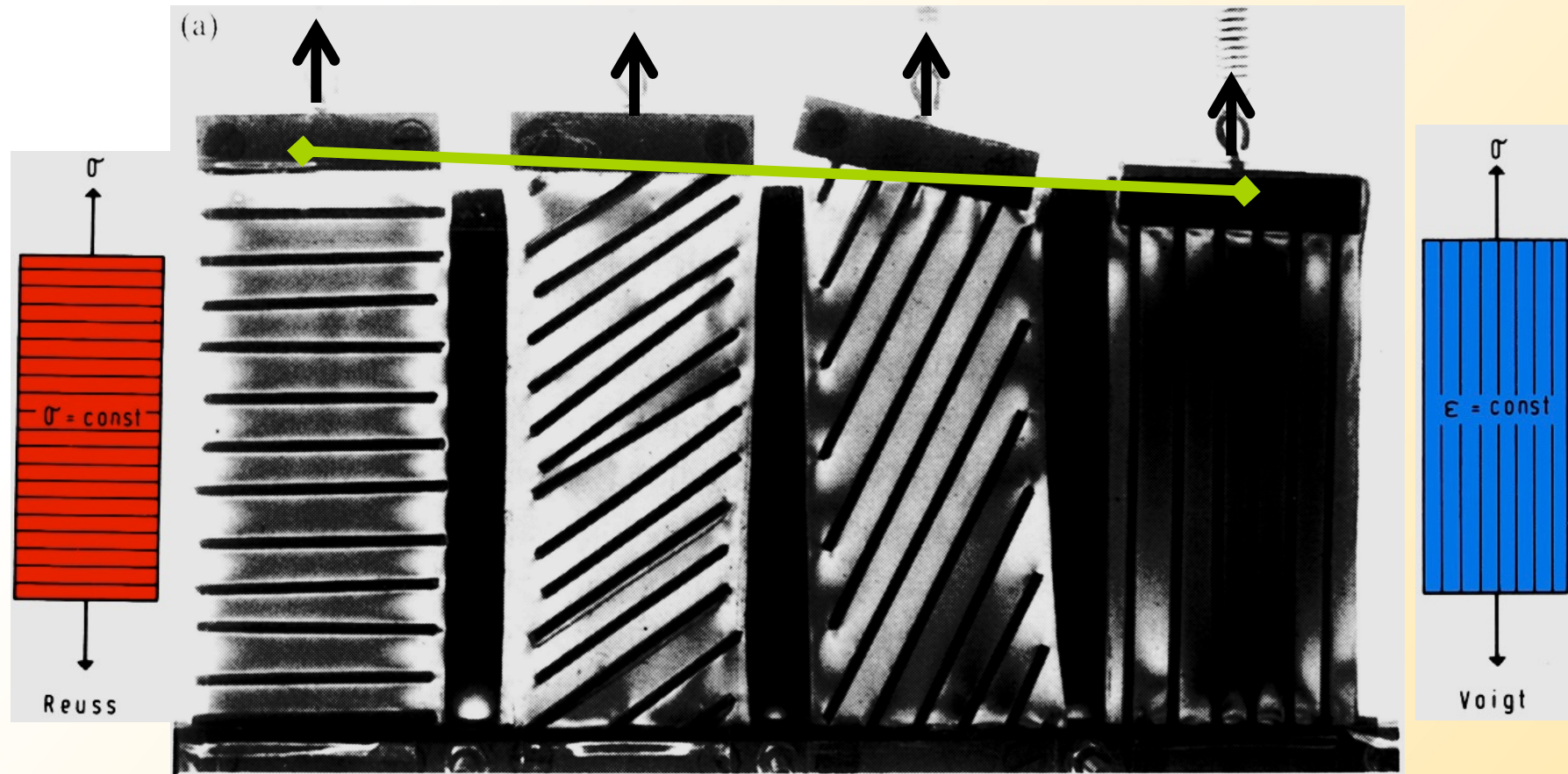
No microstructural or theoretical basis, but in practice it is close to the experimentally measured values



Reuss (b)

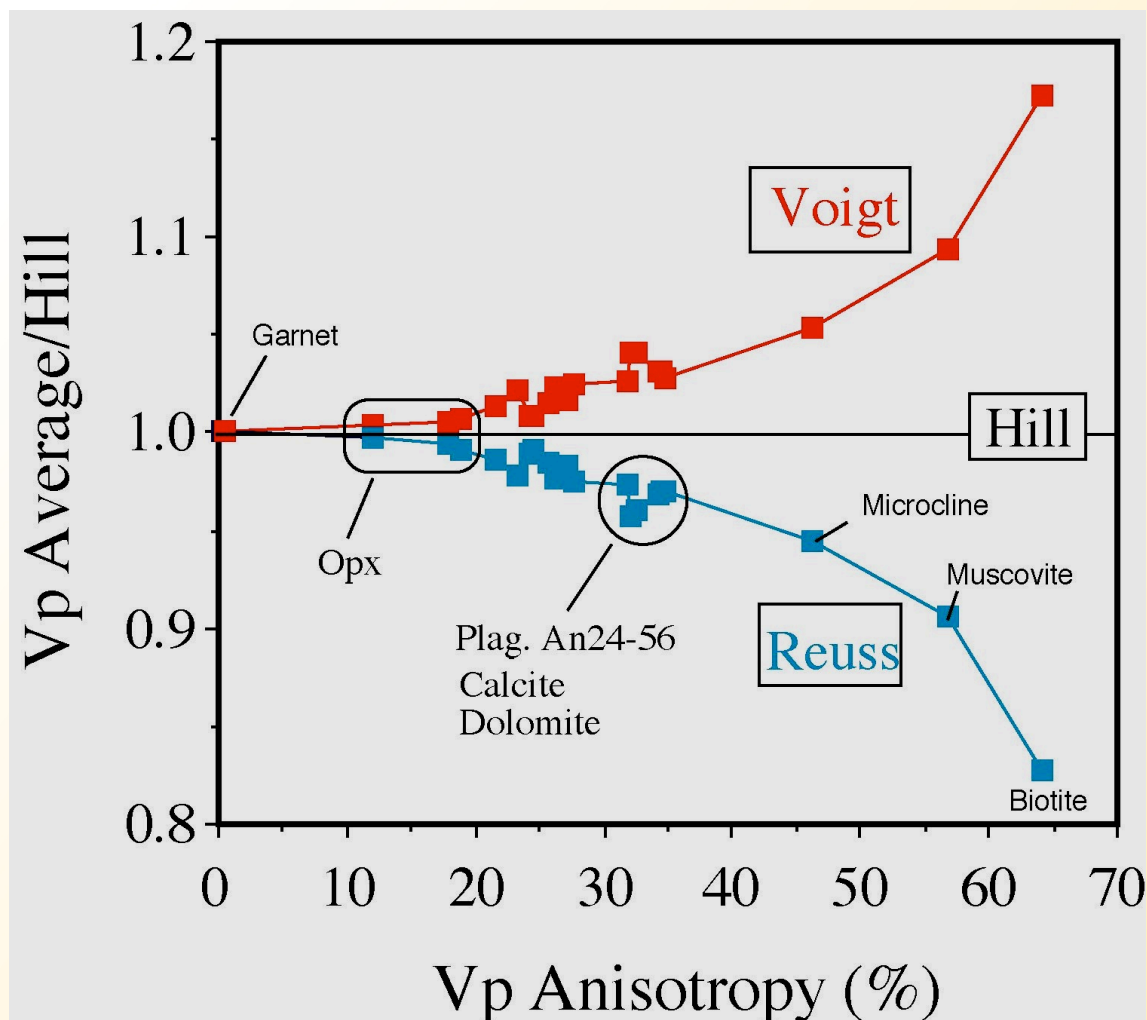
(a) Grain structure of a polycrystal for which Voigt's assumption is valid. (b) Grain structure of a polycrystal for which Reuss's assumption is valid.

Macromodel composite laminates

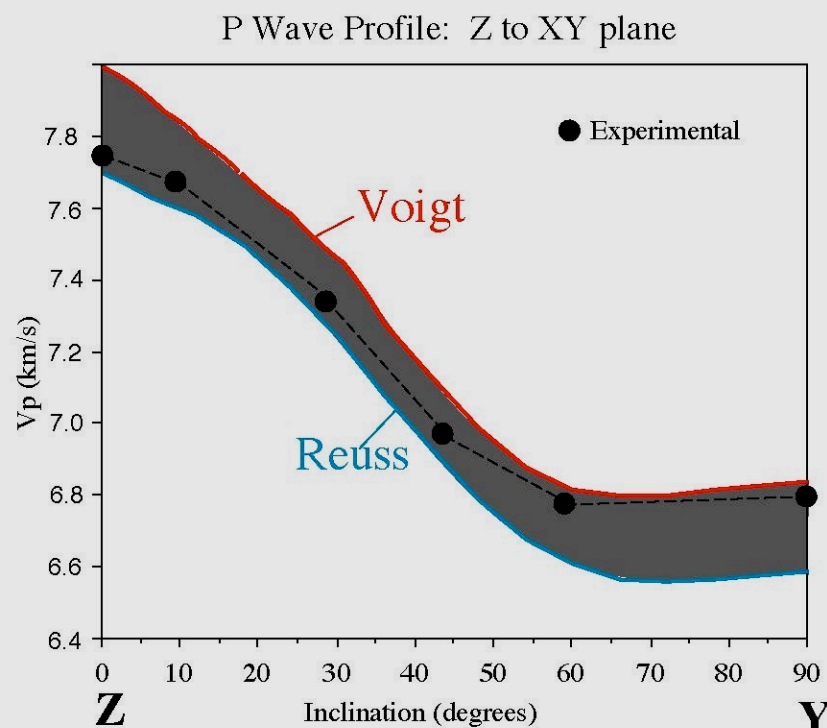
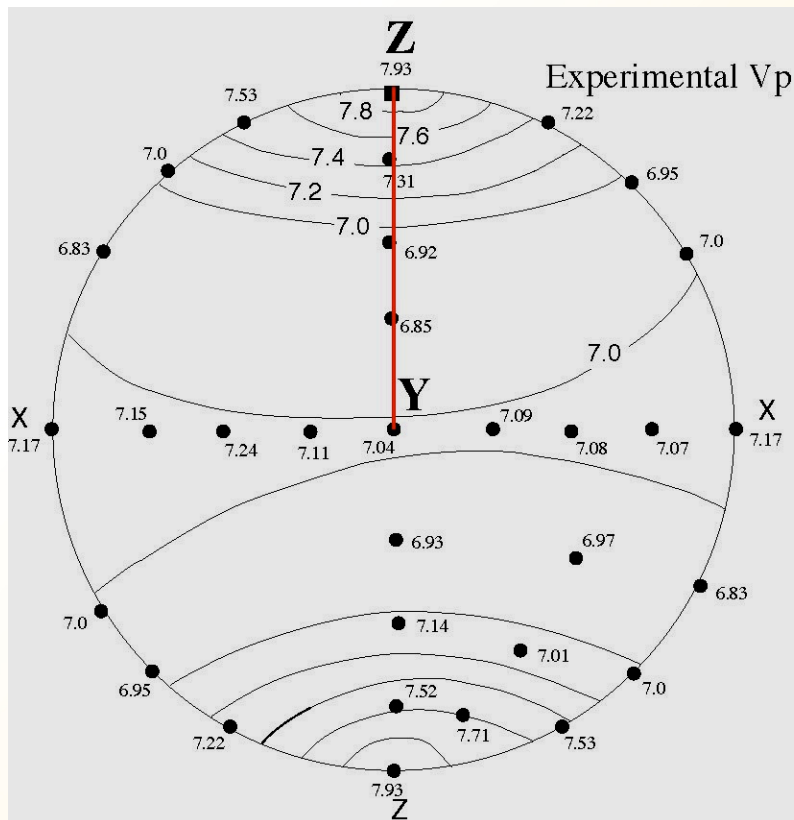


Aligned metal rods(black) in polyurethane matrix grey where lighter greys - larger shear strains

Vp Polycrystal Averages



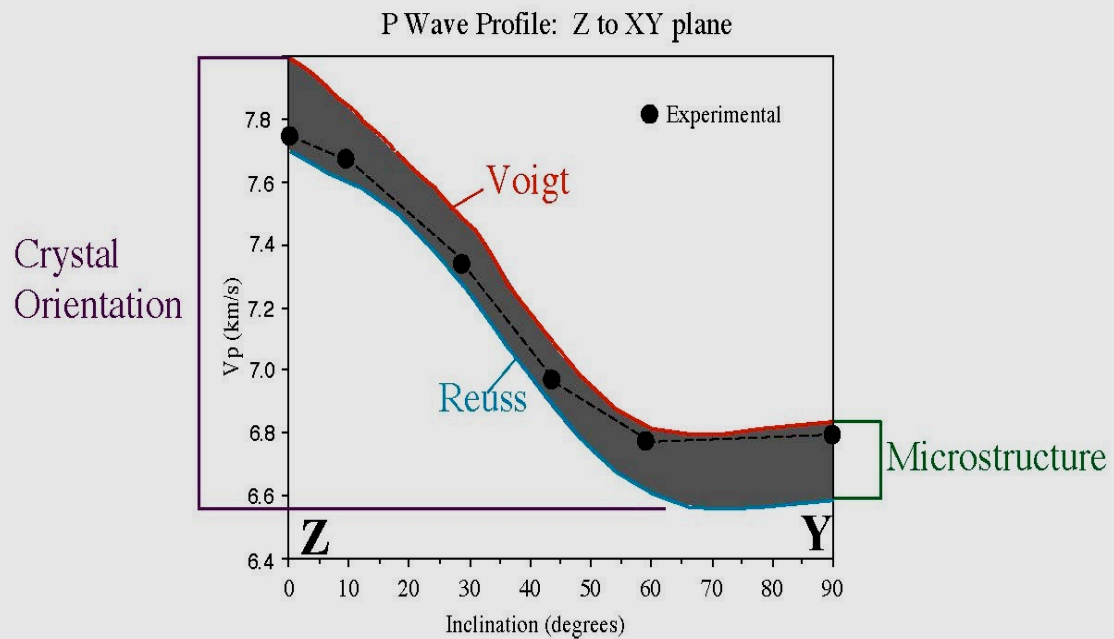
Oklahoma Gabbro Stereo & Vp YZ profile



Confining pressure = 800 MPa

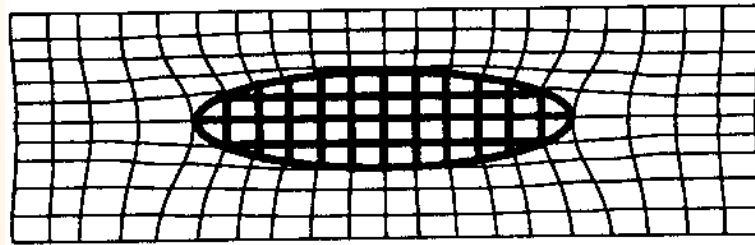
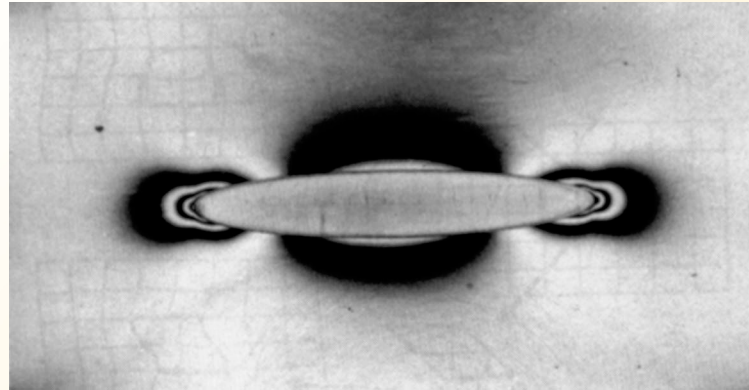
SERONT, B., MAINPRICE, D. AND CHRISTENSEN, N.I. (1993)
J. Geophys. Res. 98: 2209-2221.

Physical Properties & Microstructure - Composite materials approach



Crystal Orientation "signal" = 5 x Microstructure "signal"

Ellipsoidal inclusion - special properties

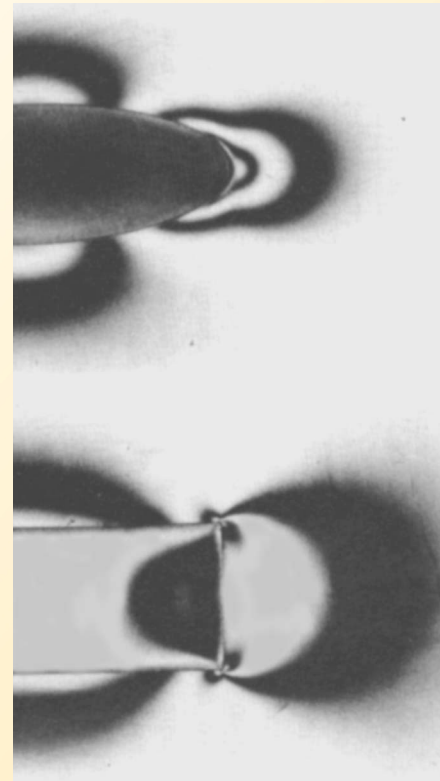


Uniform stress and strain field within the inclusion

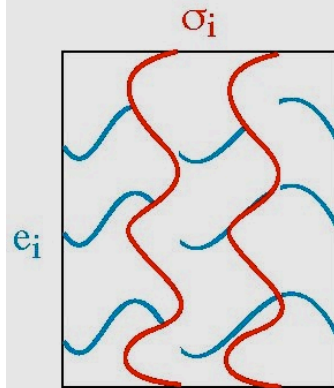
Ellipsoid and Cylinder

homogeneous stress & strain fields

heterogeneous stress & strain fields



Self Consistent (SC) Method



$$e_i \neq \text{constant} \quad \sigma_i \neq \text{constant} \quad : \quad C^* = \langle \sigma \rangle \langle e \rangle^{-1}$$

$$C^* = \langle \sigma \rangle \langle e \rangle^{-1}$$

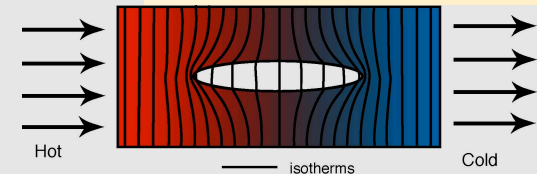
$$\langle \sigma \rangle = \sum v_i (C_i e_i) \quad \langle e \rangle = \sum v_i e_i$$

$$C^* \approx C^{SC} = \langle \sigma \rangle \cdot \langle e \rangle^{-1} = [\sum v_i (C_i e_i)] \cdot [\sum v_i e_i]^{-1}$$

The value of e_i is found using a generalization of Eshelby's inclusion theory to anisotropic inclusions in anisotropic back ground media.

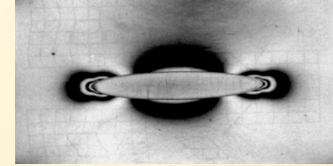
$$e_i = [I + G (C_i - C^*)]^{-1}$$

Where I is the 4th rank identity tensor and G is a Green's tensor involving elliptical intergrals over the inclusion shape.

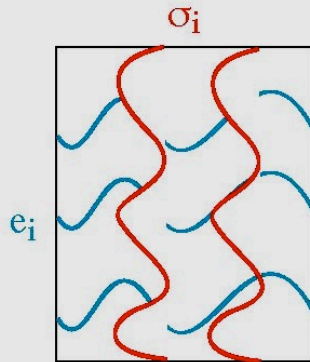


$$\sigma_{ij} = C_{ijkl} \epsilon_{kl} \quad q_i = -K_{ij} \Delta T_j$$

- much more complex to calculate
- "best" bounds for strongly anisotropic minerals
- can introduce microstructure e.g. shape via Green's tensor as ellipsoidal inclusion
- "best" bounds for mixtures with very different stiffnesses e.g. solids, liquids and voids
- treats every object (grain, void, fracture) in an identical manor
- further information about neighbour interaction could be introduced via two-point correlation functions



DEM Method

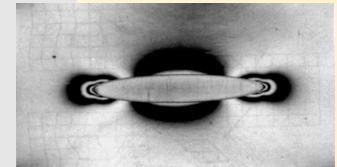


$$e_i \neq \text{constant} \quad \sigma_i \neq \text{constant} \quad : \quad C^* = \langle \sigma \rangle \langle e \rangle^{-1}$$

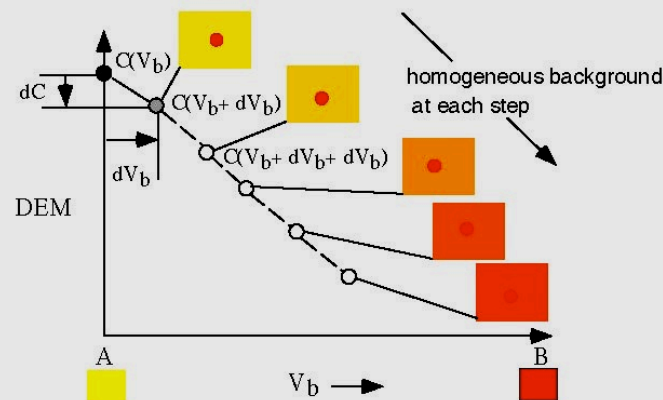
$$C^* = \langle \sigma \rangle \langle e \rangle^{-1}$$

$$\langle \sigma \rangle = \sum v_i (C_i e_i) \quad \langle e \rangle = \sum v_i e_i$$

$$dC^* \approx dC^{\text{DEM}} = \frac{dV_b}{(1-V_b)} (C_i - C^{\text{DEM}}) e_i$$

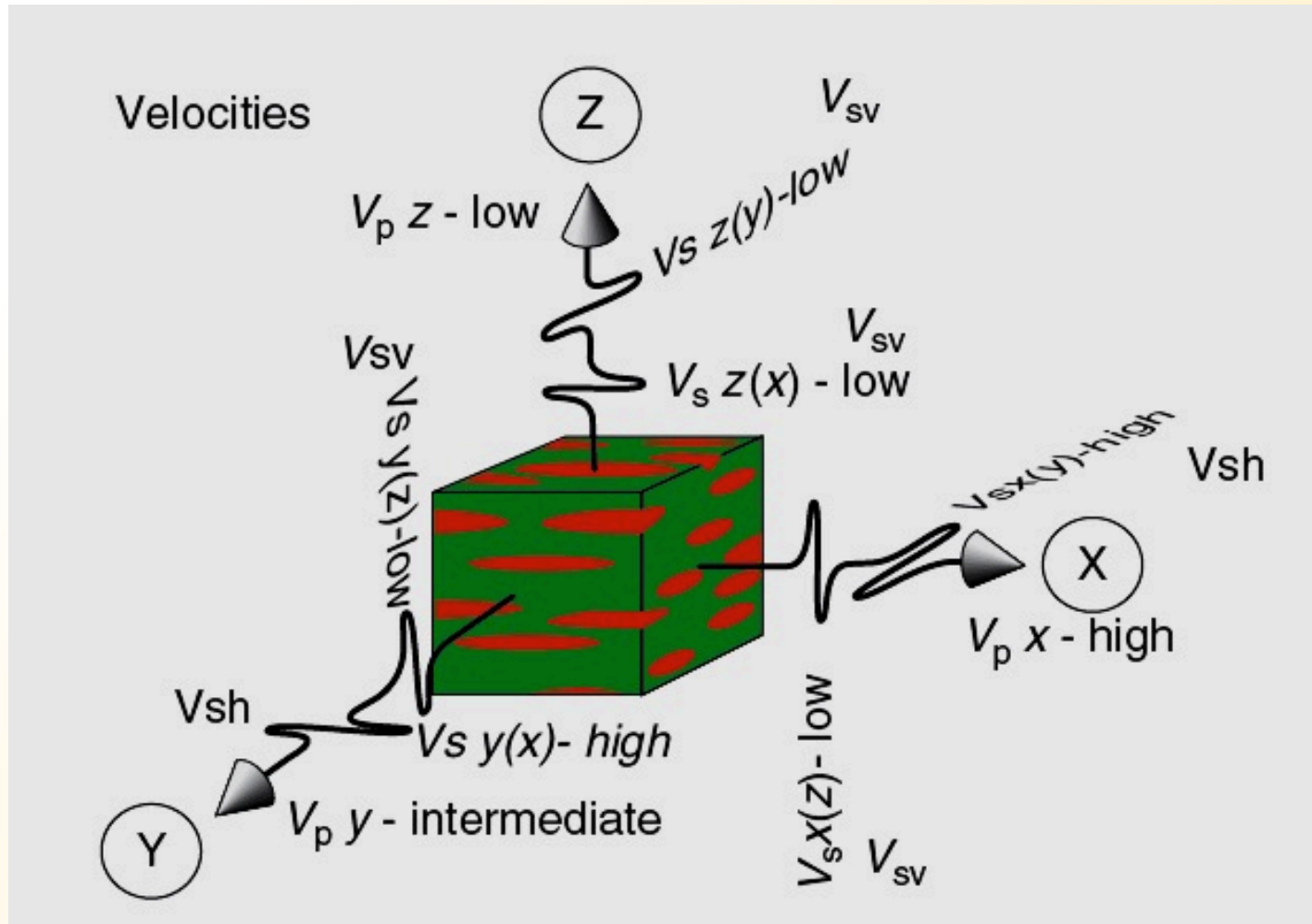


Developed for two phase aggregates type AB (McLaughlin, 1977) where $V_b = 1 - V_a$. The phase B is the included phase and phase A is the host phase. Requires some initial value of C^* at a composition V_b - this may be a pure end member. The incremental nature of the differential approach preserves the percolating (connectivity) properties of the initial estimate. For example if starting from pure A, A will always be connected. If starting from 50:50 A:B then both phases will be connected.

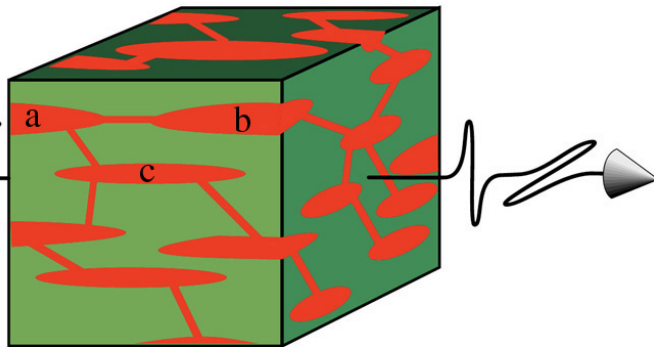


- now phases A and B are microstructurally different (A=Host, B= Inclusion)
- preserves connectivity of initial estimate

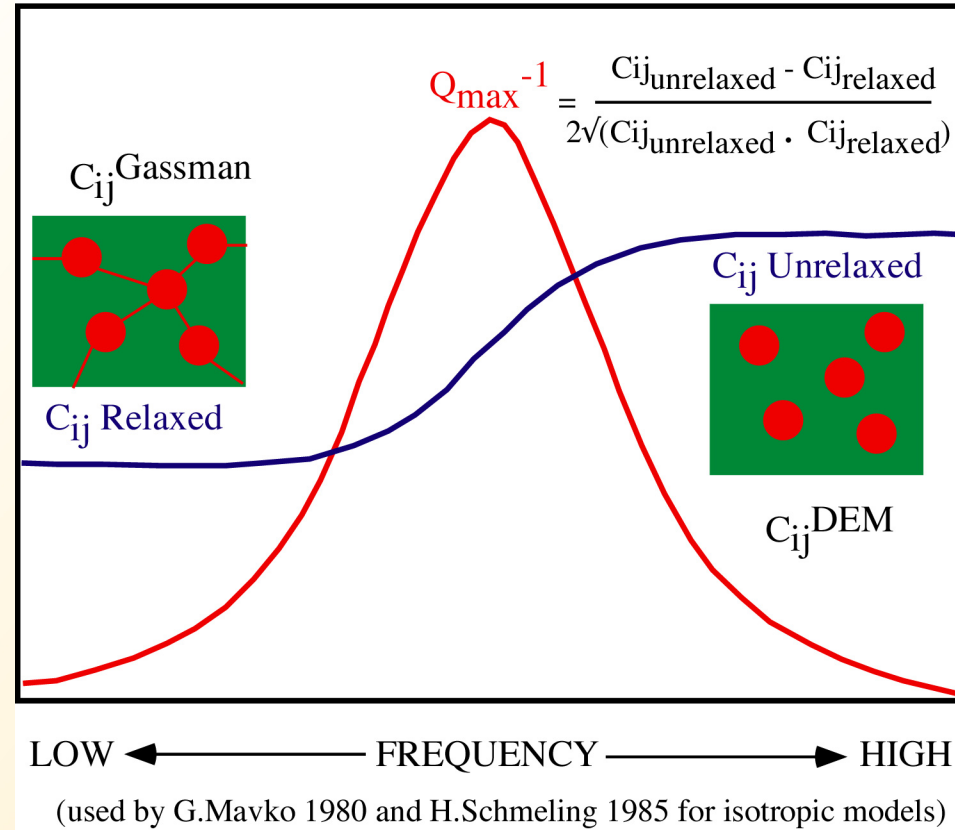
Velocity Cube – with melt inclusions



squirt flow attenuation



Standard Linear Solid (Zener, 1948)

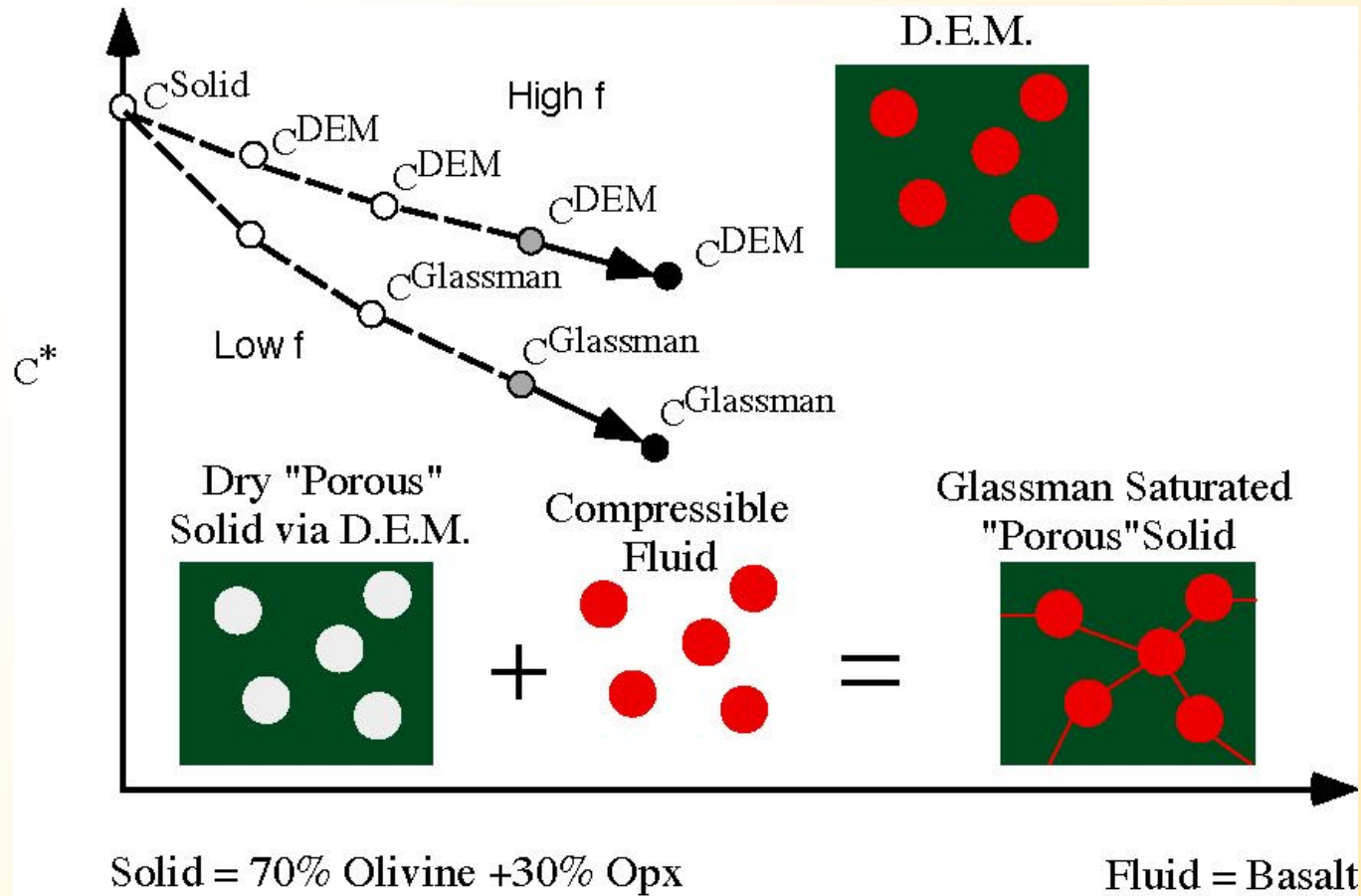


Unrelaxed Modulus - High Frequency : $P_a \neq P_b \neq P_c \dots 0\%$ connectivity
 --> given by Self Consistent (S.C.) or Differential Effective Medium (D.E.M.).

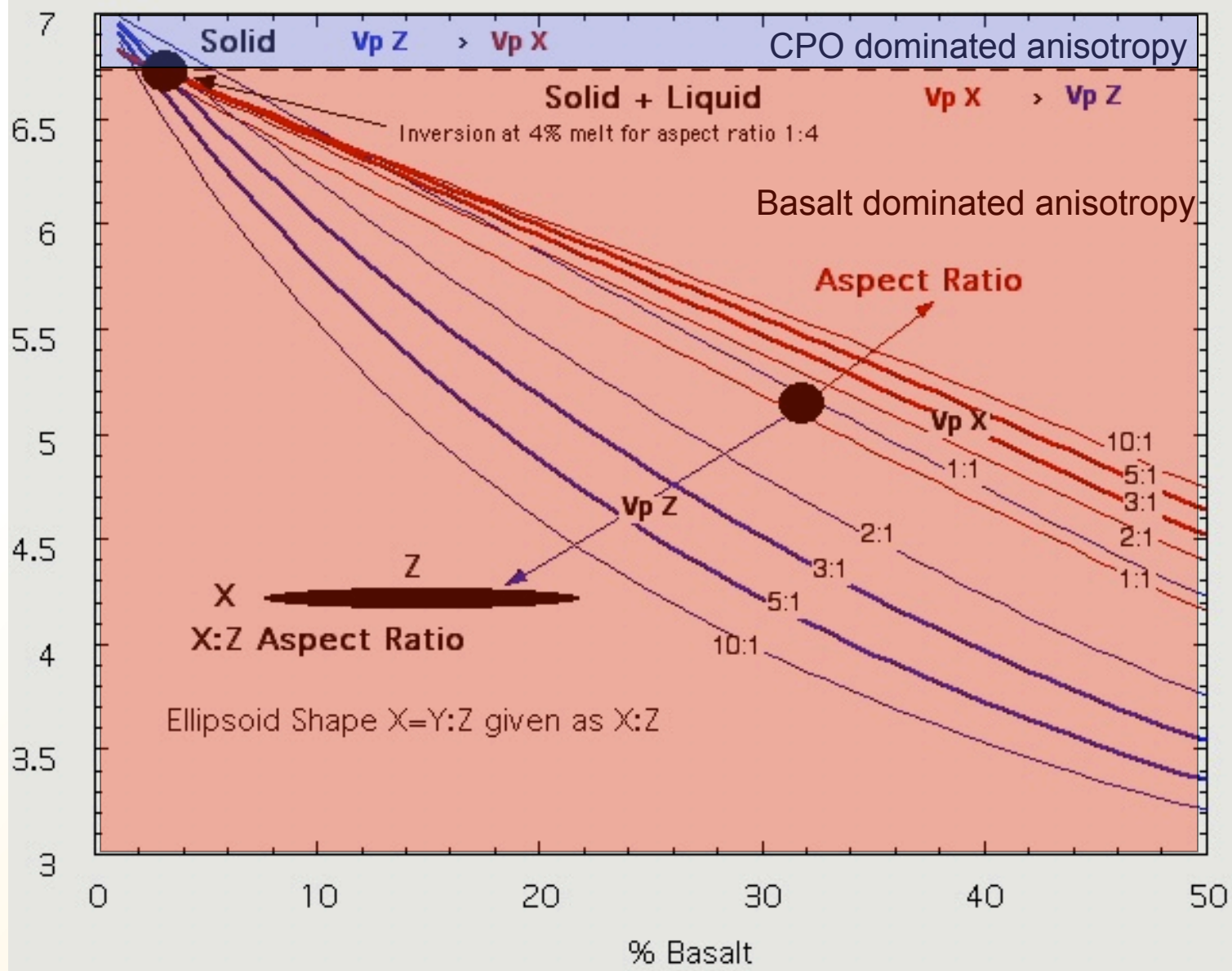
Relaxed Modulus - Low Frequency : $P_a = P_b = P_c \dots 100\%$ connectivity
 --> given by Gassmann (1951) relationship extended to anisotropic media (Brown & Korringa, 1975).



Combining DEM and Poroelasticity theory

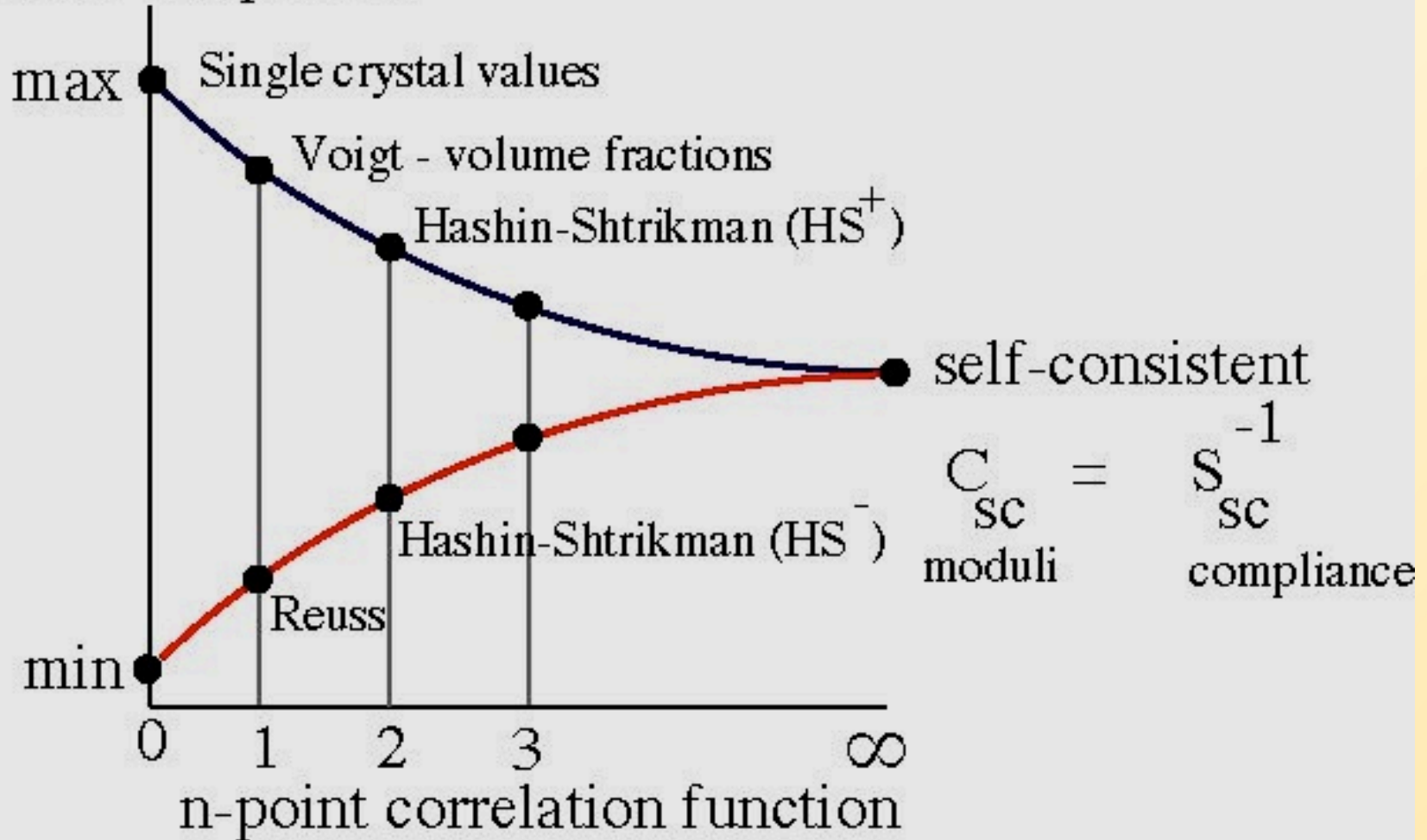


VpX and VpZ Gabbro - Basalt : Effect of Aspect Ratio



MAINPRICE, D. (1997) Tectonophysics 279,161-179.

Elastic Properties



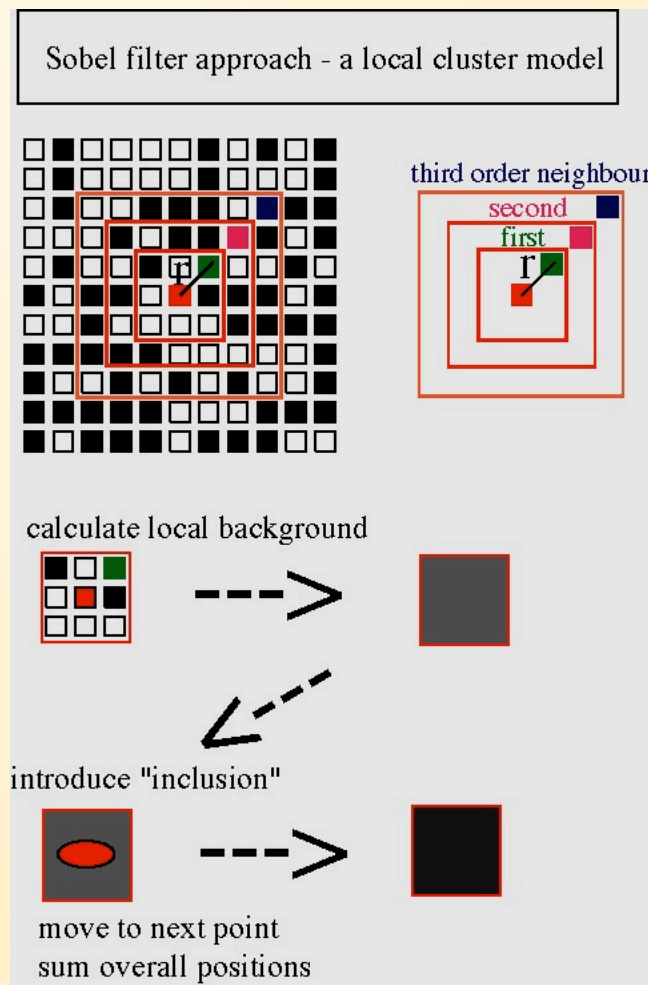
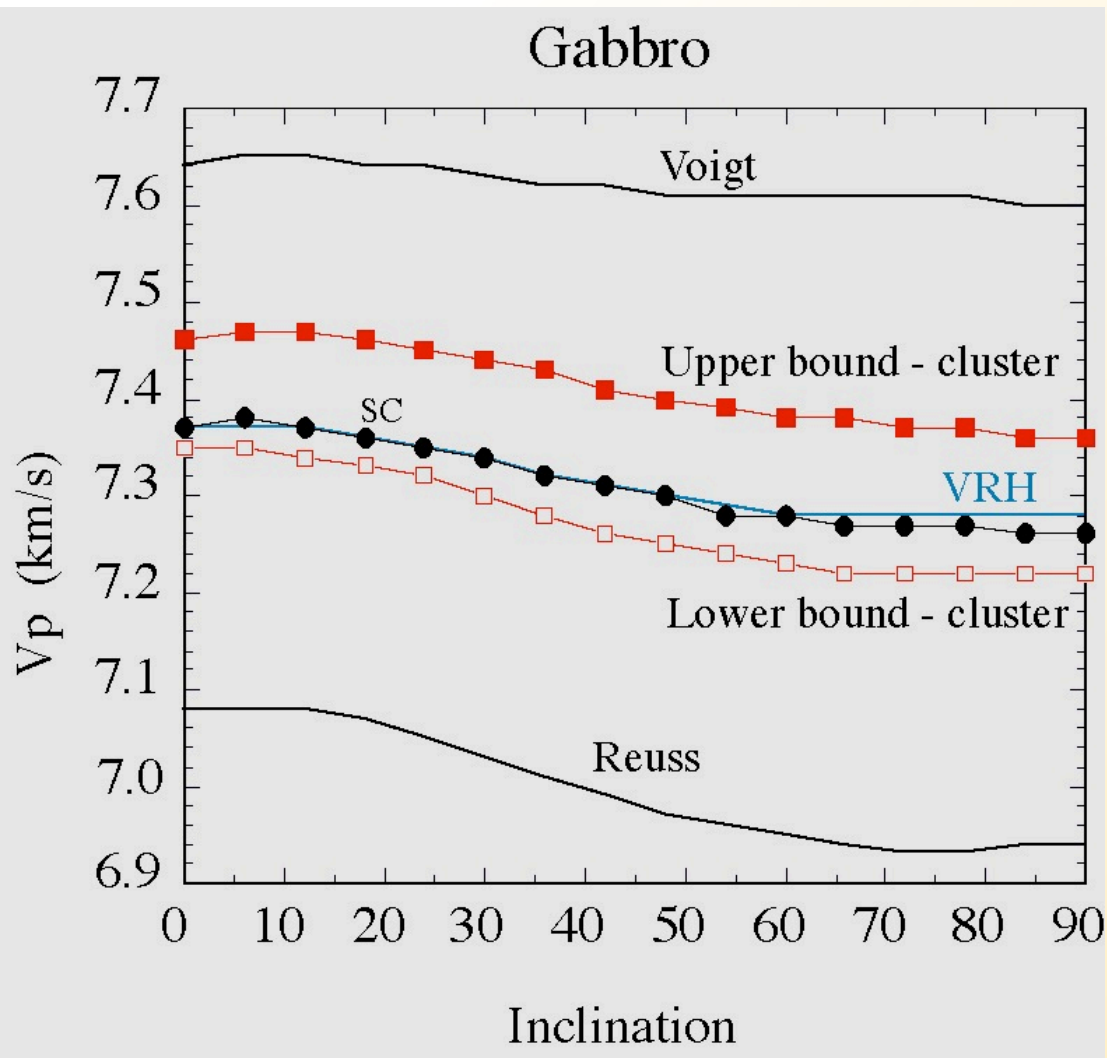
The Hill Series

The ideal of Hill series originally suggested by Adam Morawiec

Using the Hill average it iterates until the $C_{ij} = \text{inv}(S_{ij})$ considered to be of physical importance

I have tested this 3 iterates gives result, it is very close to the Hill average.

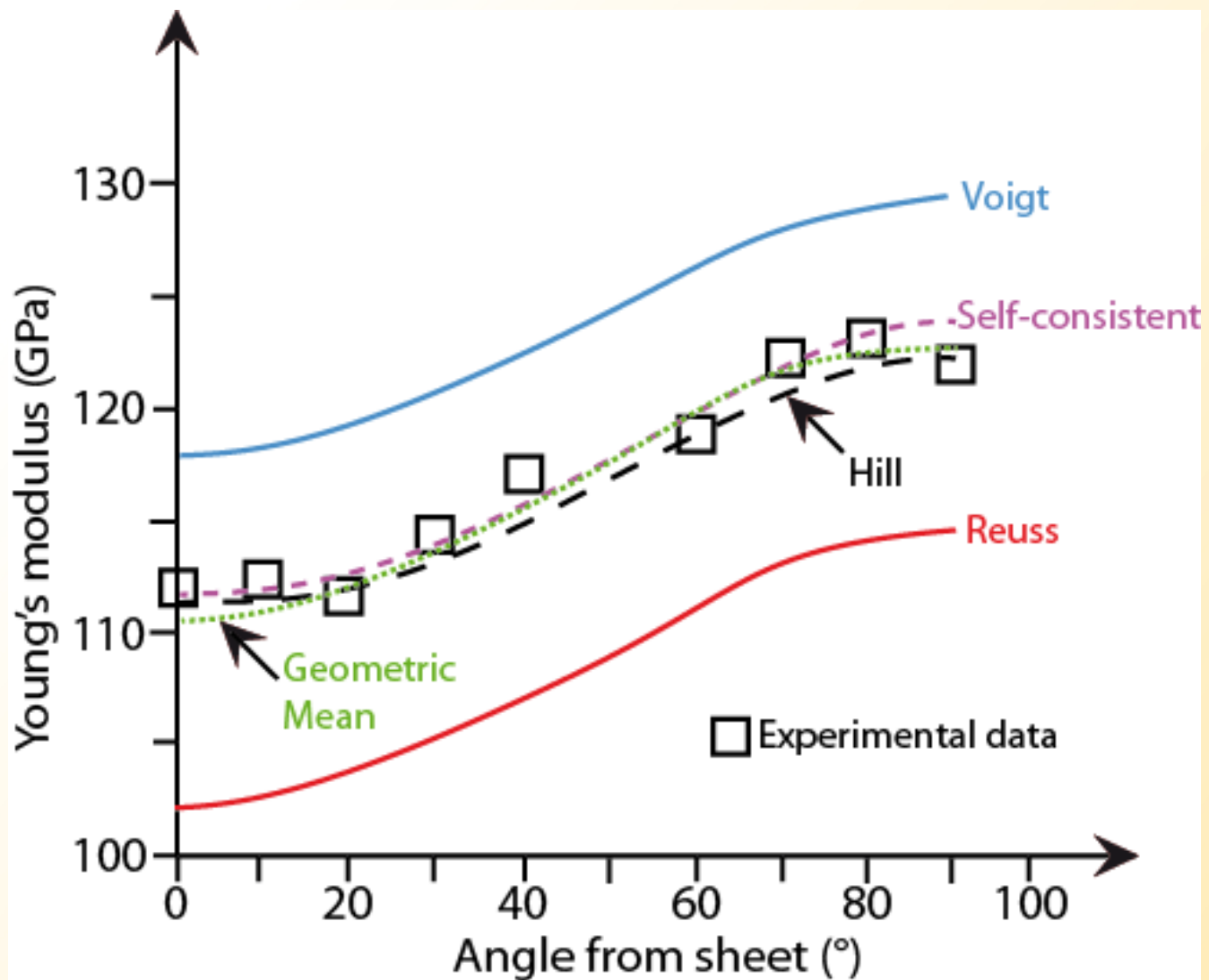
Cluster method



Upper
Voigt

SC(shape) GM

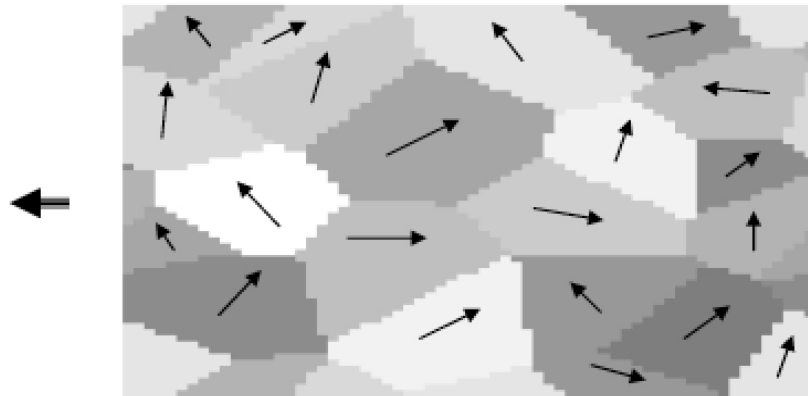
Lower
Reuss



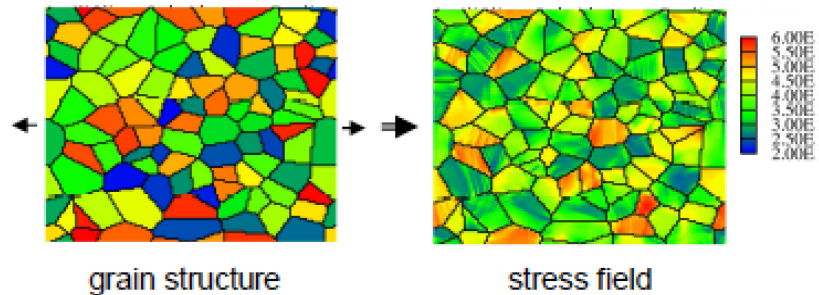
Polycrystal models: full-field vs. mean-field (statistical) approaches

Lebensohn, Tome & Ponte Castaneda, Phil Mag 87, 4287 (2007)

Full-field solution

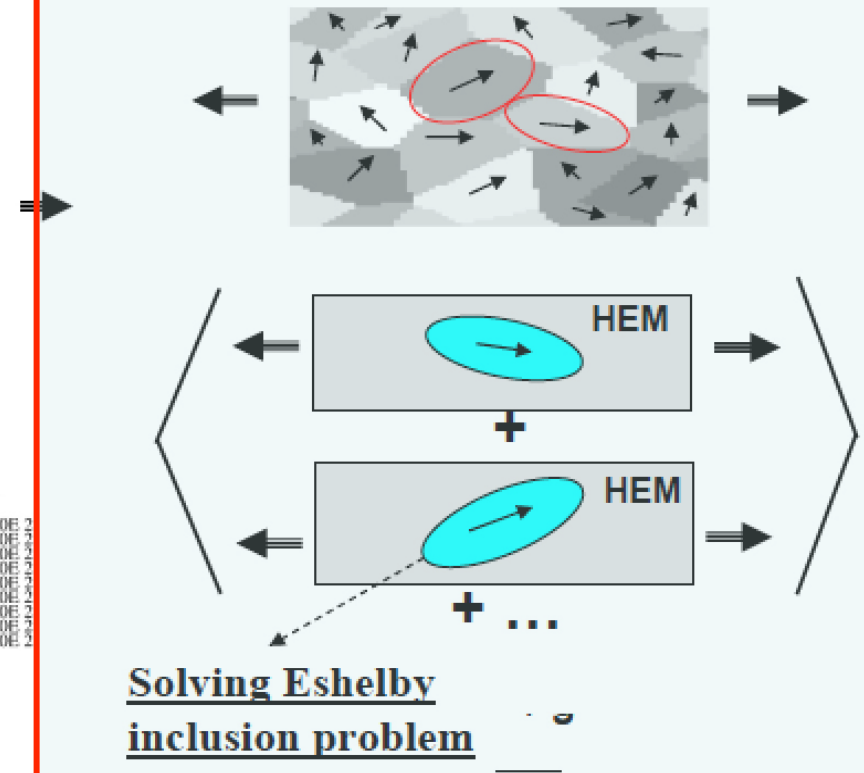


Solving equilibrium + compatibility:



Local behavior inside grains with a particular orientation and a particular neighborhood

Mean-field statistical solution



Average behavior of grains with same orientation and different neighborhoods

Validation of mean-field formulations with full-field models requires ensemble averages of the latter

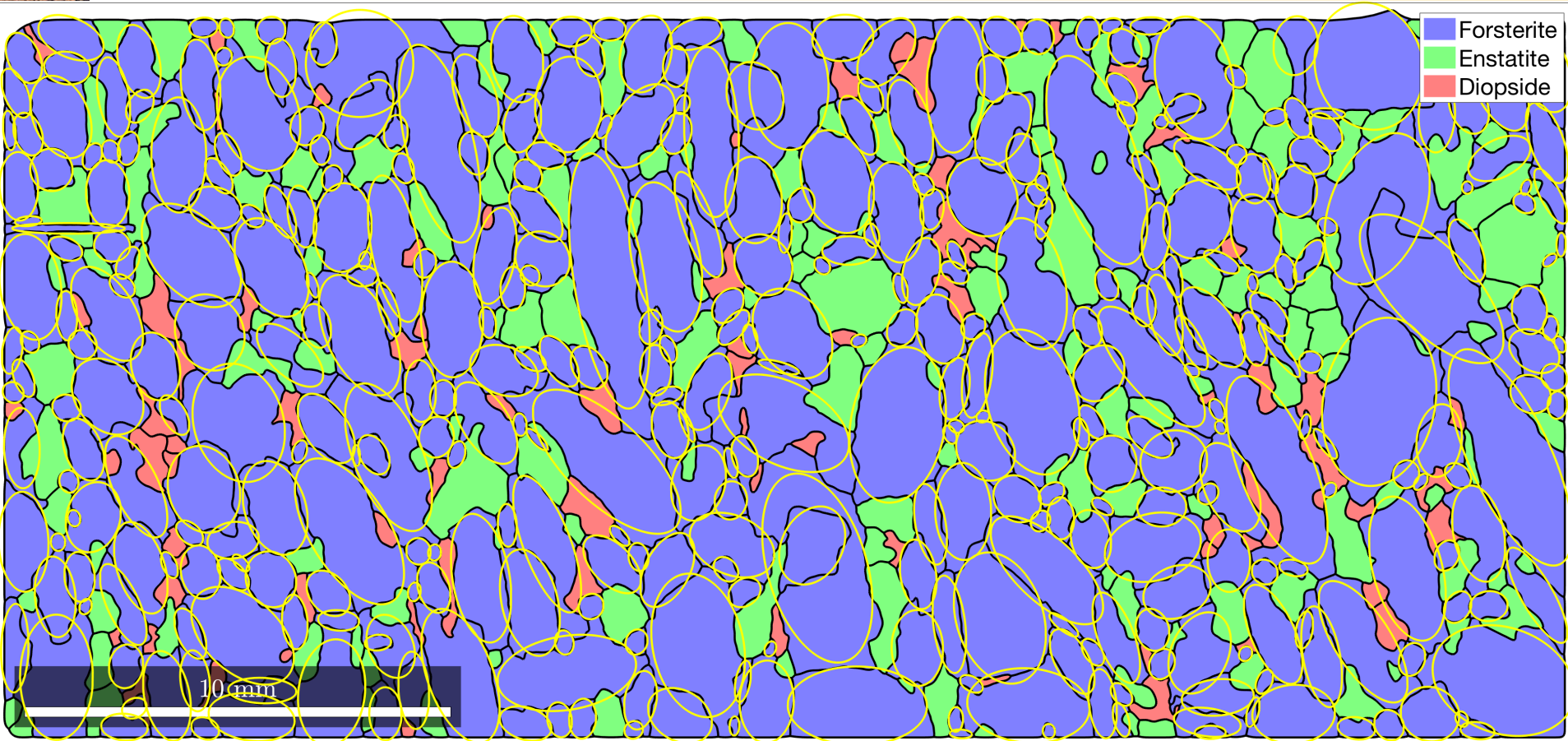
Recent developments; Self-consistent (SC) and Differential Effective Media (DEM), FFT methods

- DEM scheme as translated into MatLab scripted with GUI interface. Published in Computers & Geosciences
- You can download the script and manual from ;
<https://github.com/ekim1419/GassDem>
- The SC scheme has translated into MatLab scripted, we use MTEX function to export EBSD 2D data into the SC scheme. Demo_SC_Data_export.m
- | phi1 | Phi | phi2 | area | omega | a | b |
|-------|---------|---------|-------|---------|---------|---------|
| 39.91 | 131.882 | 337.894 | 85356 | 177.047 | 263.951 | 102.935 |
- **Mean grain orientation area for weighting omega angle of long & short axes**

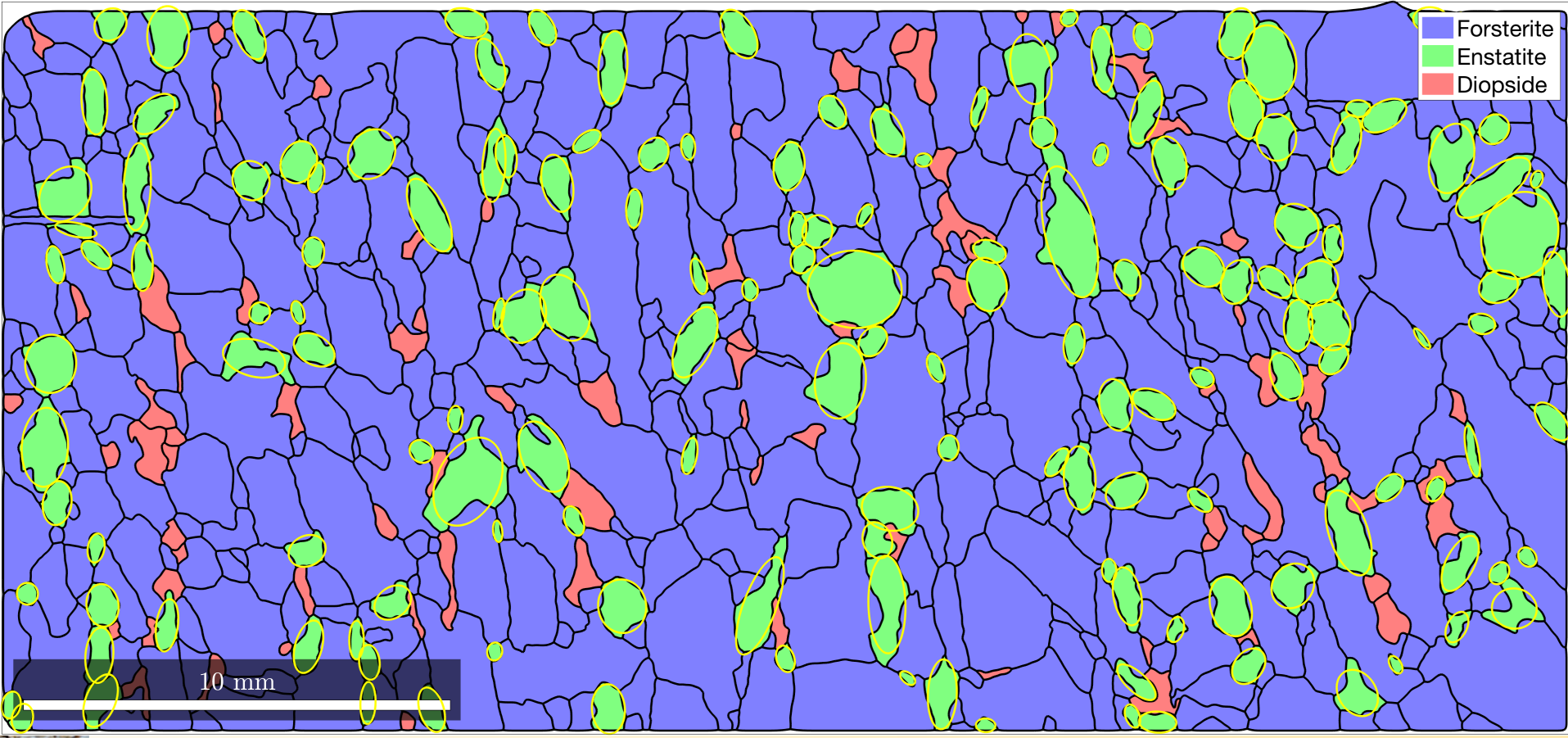
Self-consistent

For each mineral

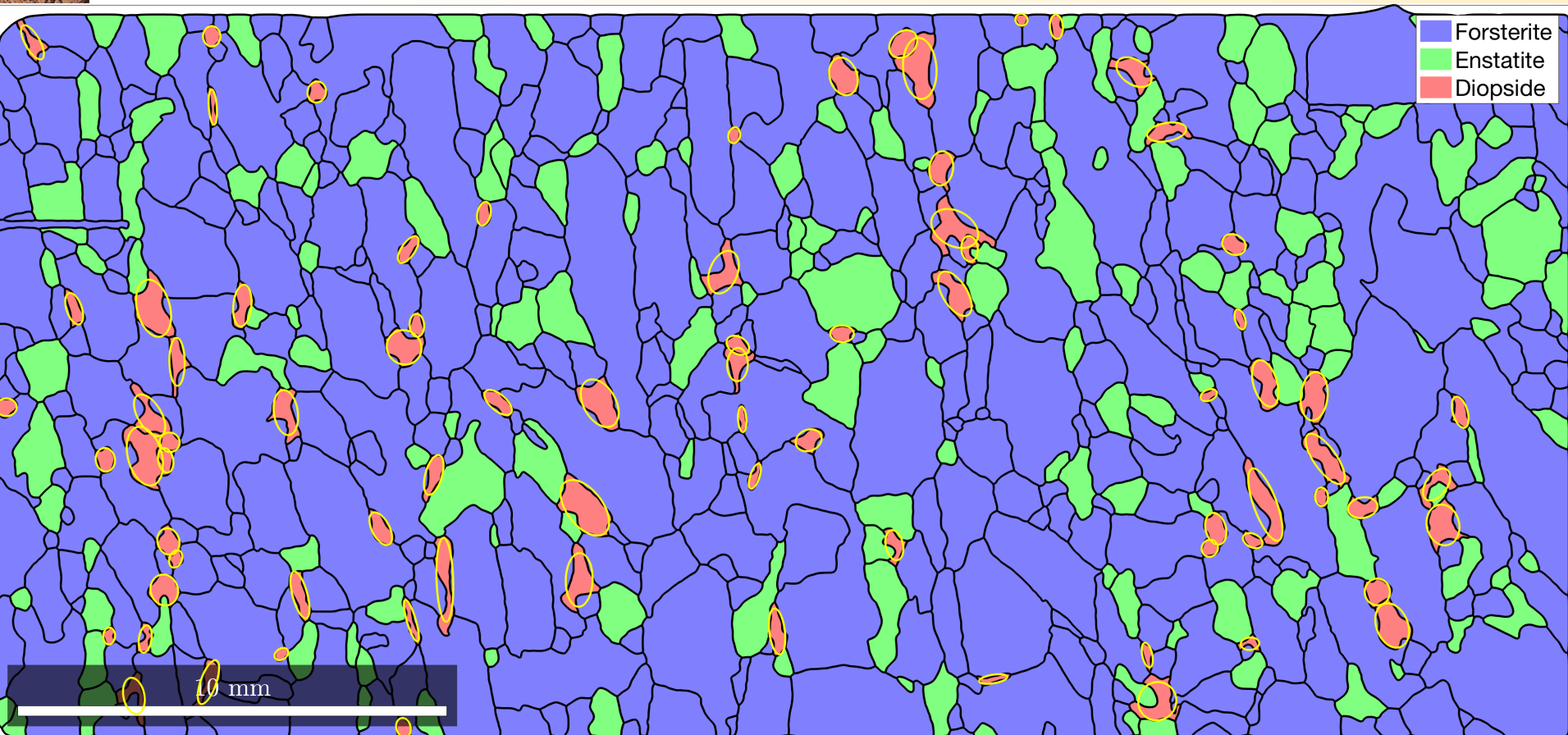
1. Fit ellipses – get long and short axis & orientation (ω)
2. Grain mean orientation, not ebsd pixels !
3. Grain Area for weighting



Enstatite Orthopyroxene green



Diopside Clinopyroxene red



Future developments; Importing single crystal tensors from a database

- Over a hundred files for the elasticity in **M**TEX format for copy and pasting into **M**TEX M-files.
- It was envisaged to use the Materials Open Database “MPOD” initiated by Daniel Chateigner (<http://materialproperties.org>). However no provision for the tensor frame in this database that is supervised by International Union of Crystallography (IUC).
- Need to agree on the **M**TEX file format that includes the the tensor frame.

Anisotropy of for stiffness elastic tensors

- First example for Cubic symmetries Zener(1948)

$$AZ = \frac{2C_{44}}{C_{11} - C_{12}}$$

- Used in geophysics

$$A = \frac{2C_{44} + C_{12}}{C_{11}} - 1$$

- Chung and Buessem (1967) Voigt and Reuss bounds are identical for isotropic case and far apart for anisotropic case.

$$AC = \frac{G^{Voigt} - G^{Reuss}}{G^{Voigt} + G^{Reuss}}$$

- Ranganathan and Ostoja-Starzewski (2008)
- Universal – requires to subtract to get zero

$$AU = \frac{K^{Voigt}}{K^{Reuss}} + 5 \frac{G^{Voigt}}{G^{Reuss}} - 6 > 0$$

- Kube (2016) Log-Euclidean naturally has zero for isotropic

$$AL = \sqrt{\left[\ln \left(\frac{K^{Voigt}}{K^{Reuss}} \right) \right]^2 + 5 \left[\ln \left(\frac{G^{Voigt}}{G^{Reuss}} \right) \right]^2}$$

Ranganathan and Ostoja-Starzewski (2008)

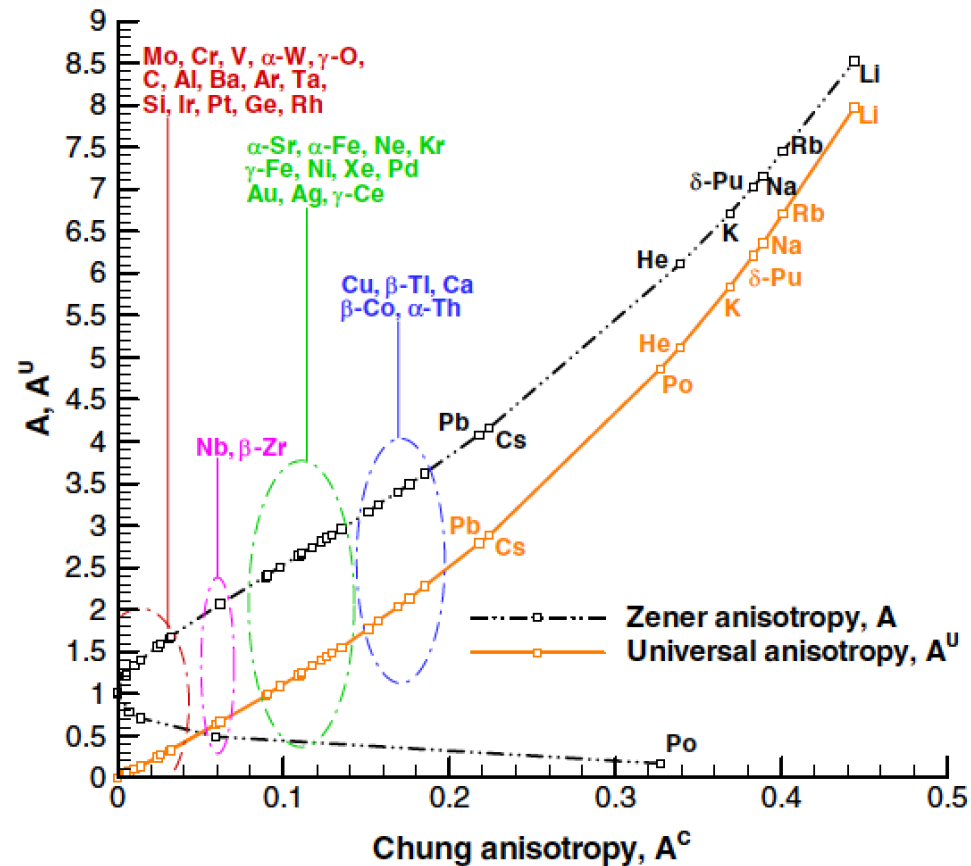
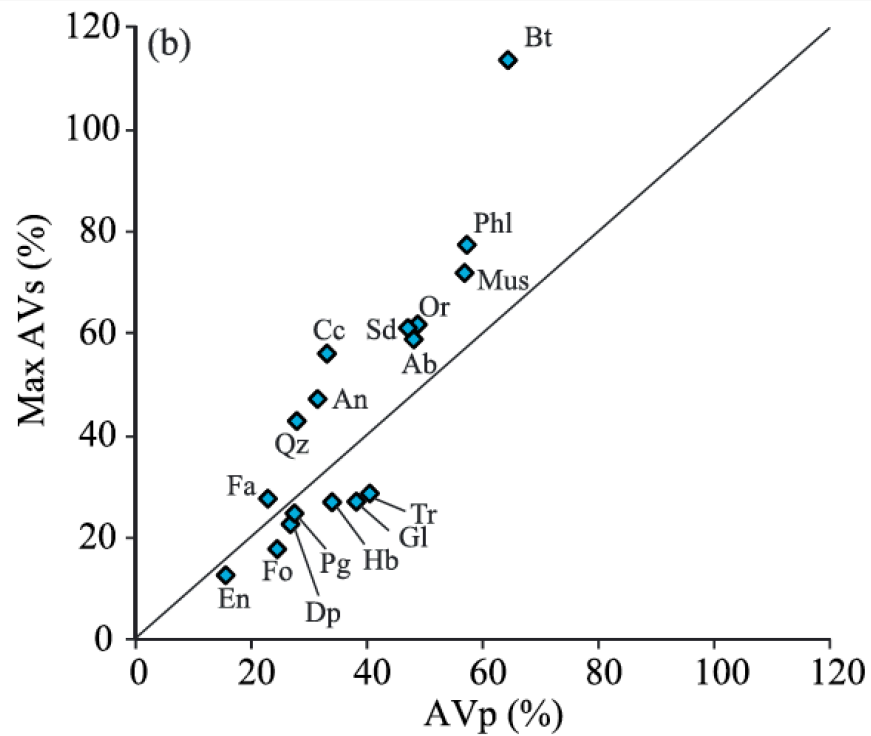
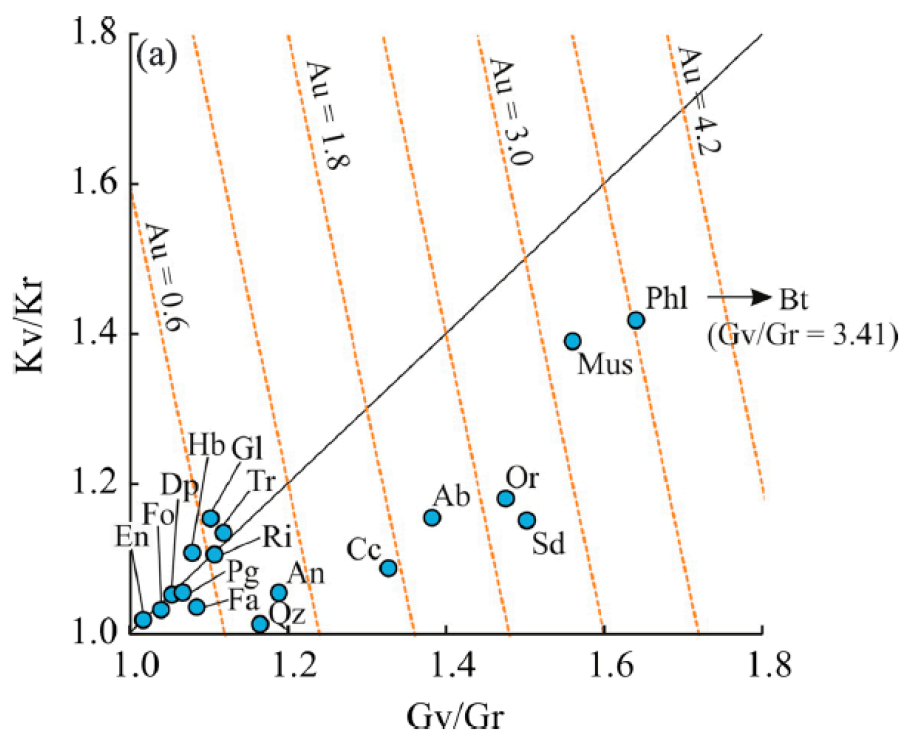
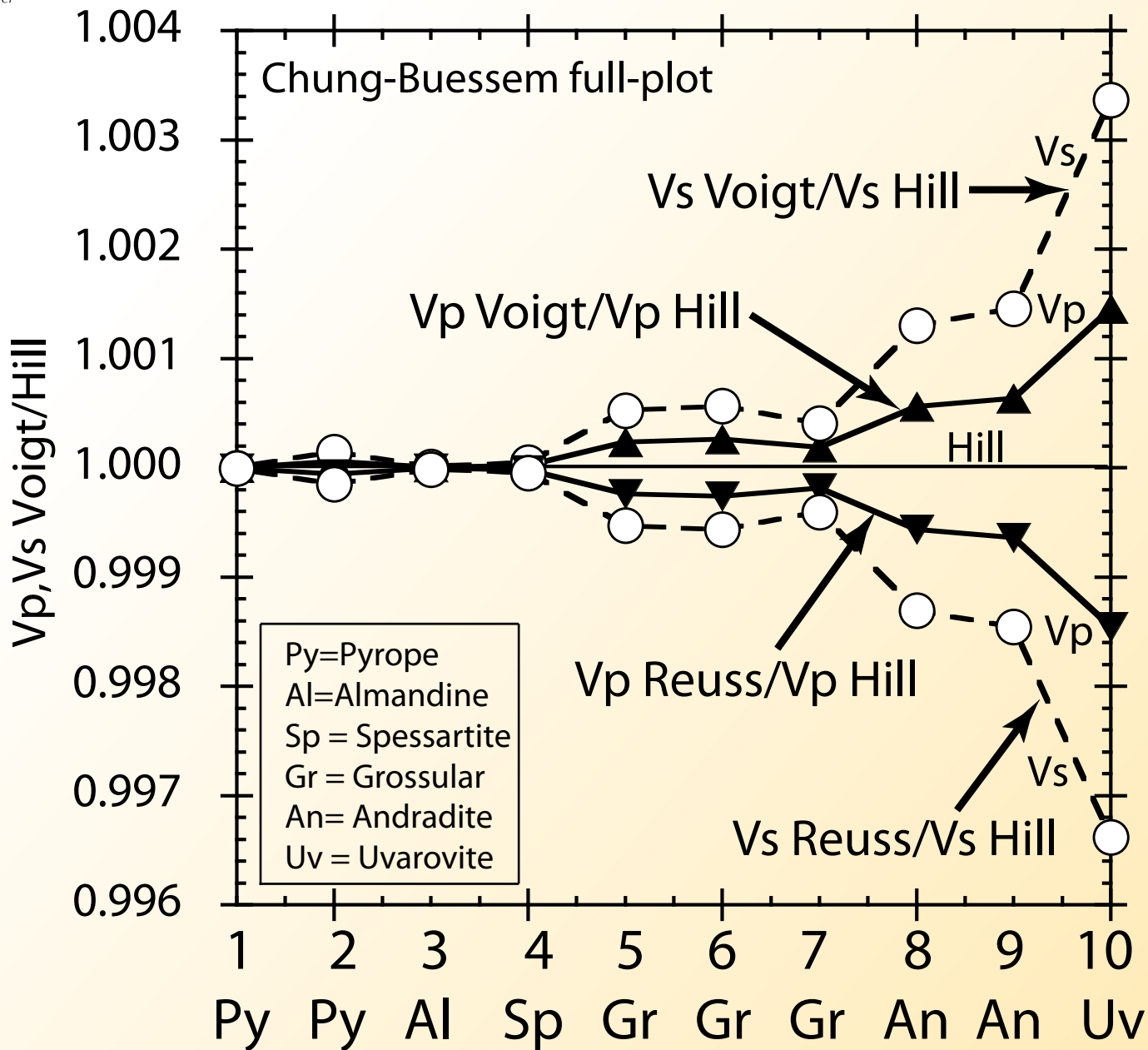


FIG. 1 (color online). Anisotropy measures for cubic crystals (Zener and Ledbetter-Migliori anisotropy measures coincide for cubic crystals). Single crystal data obtained from [3,11].

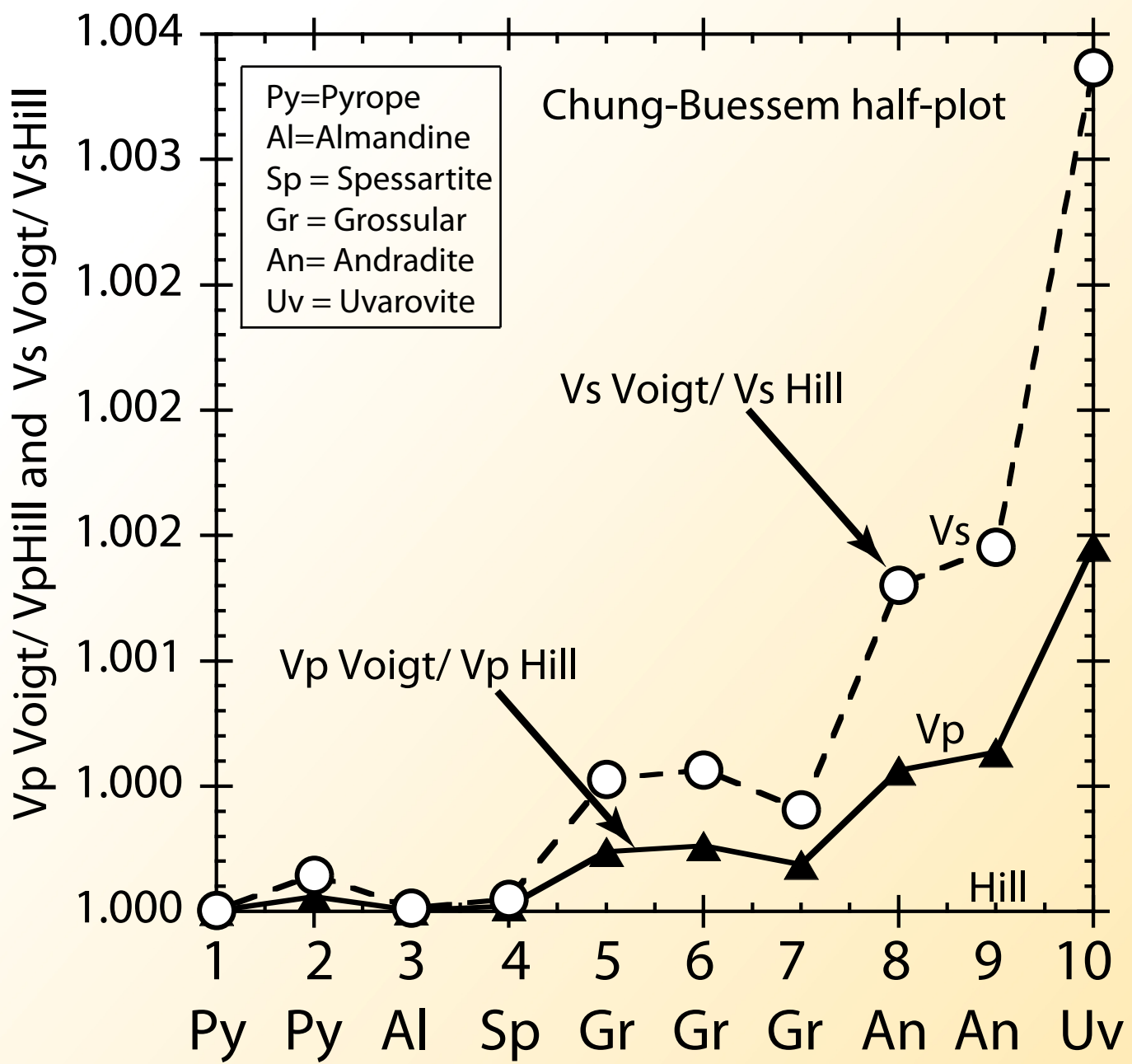
Layered minerals always largest anisotropy



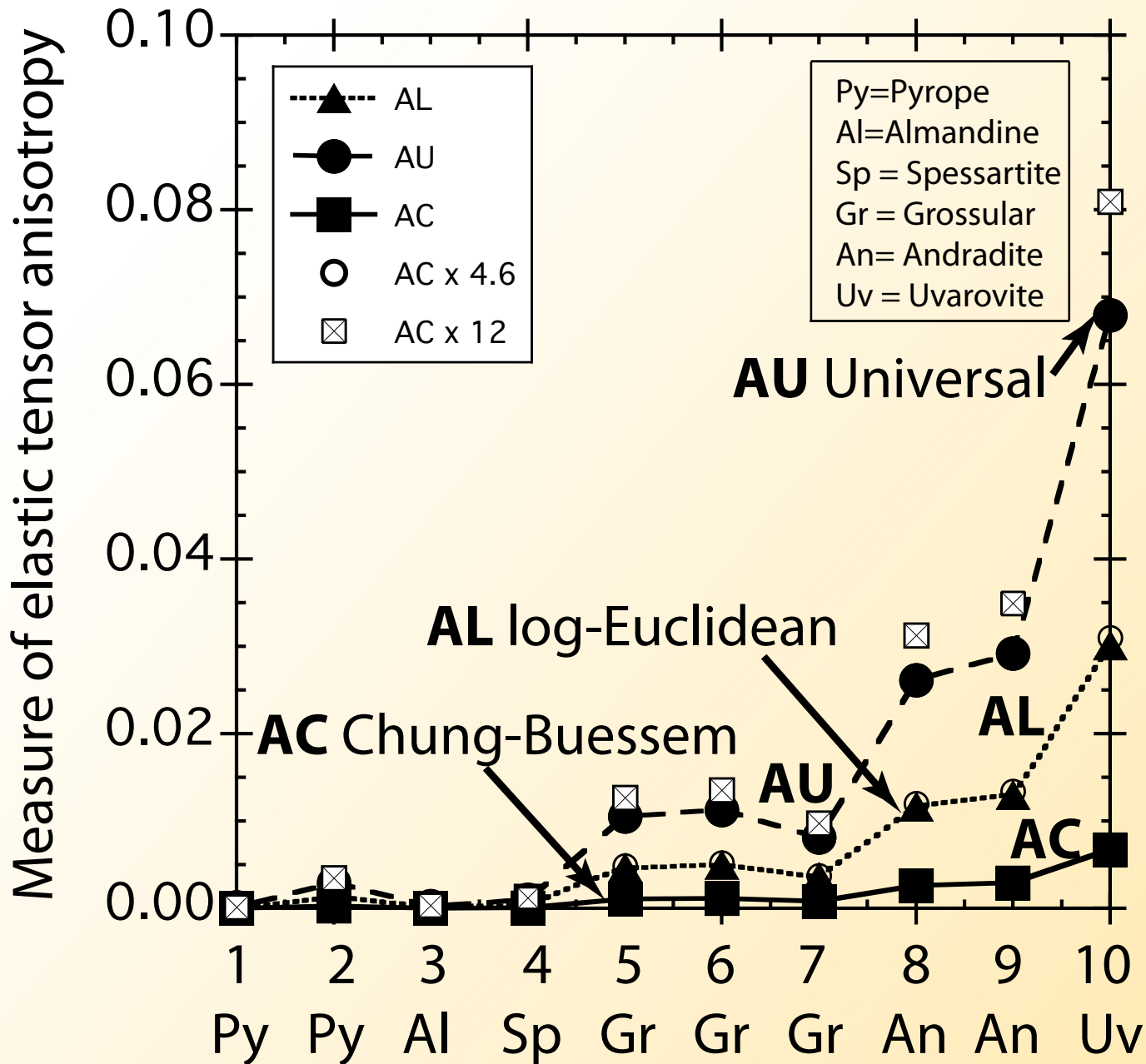
Garnet series anisotropy



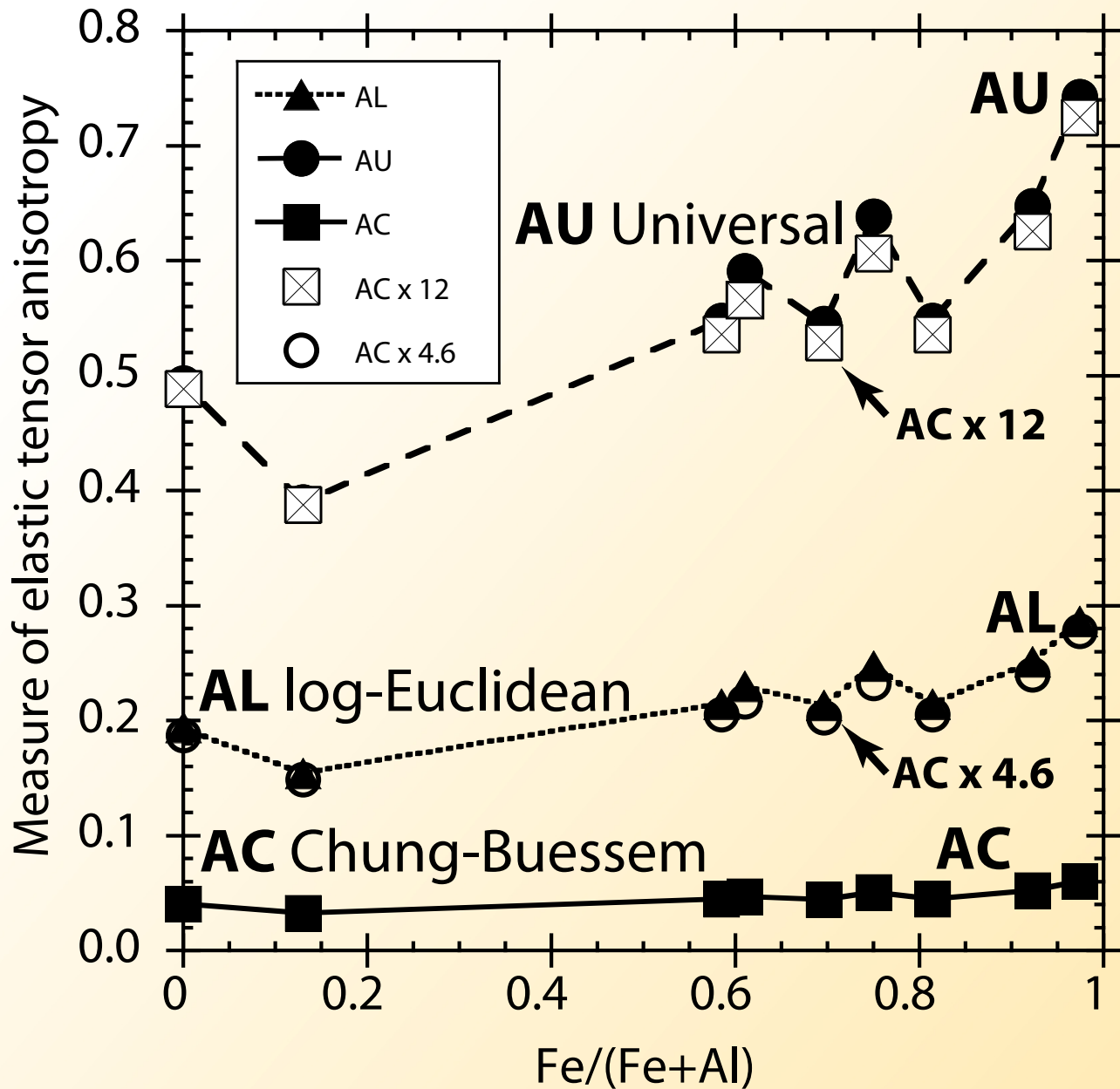
Garnet series anisotropy



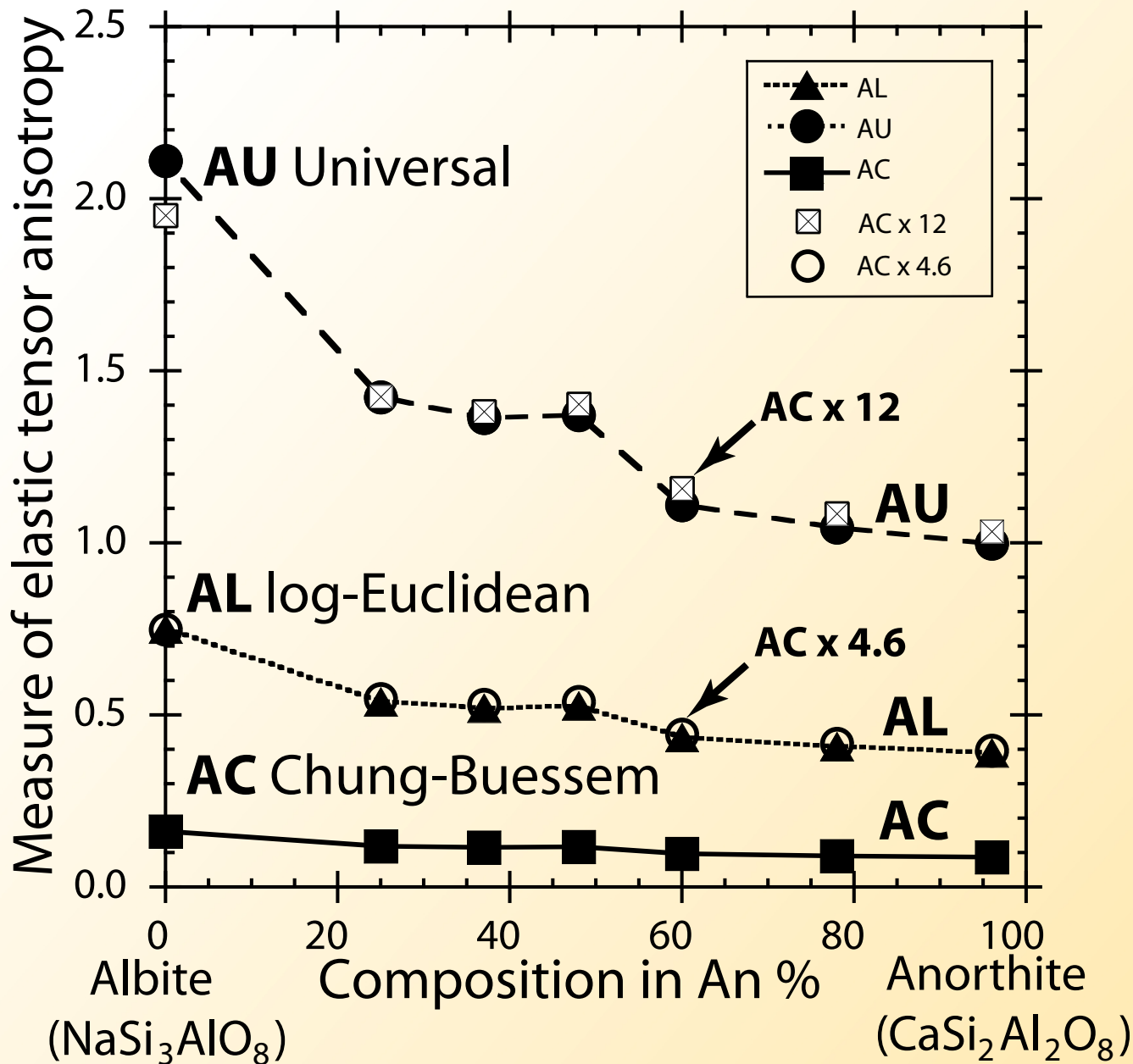
Garnet series anisotropy

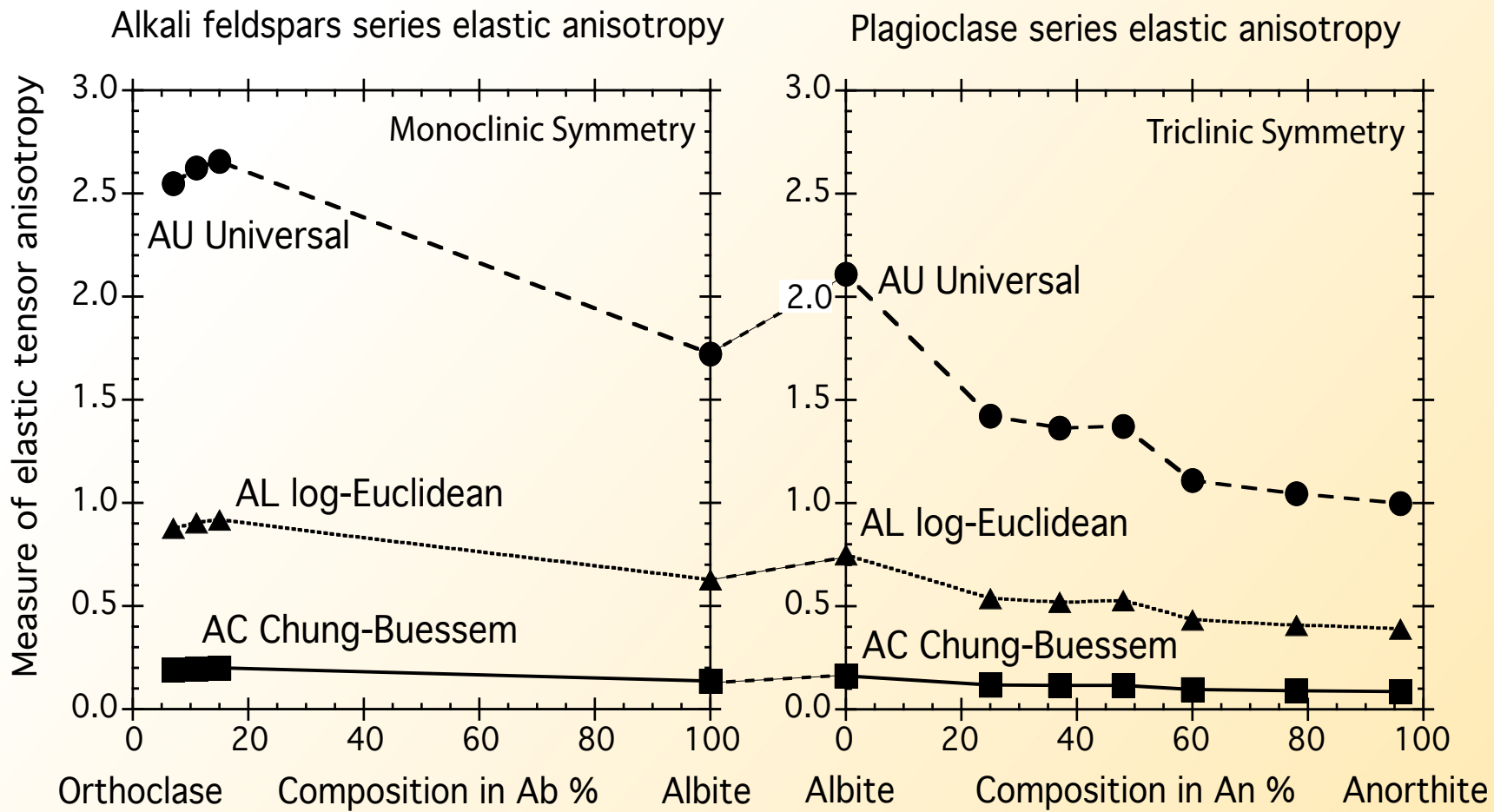


Amphibole series anisotropy



Plagioclase series elastic anisotropy





Albite – monoclinic 13 **Albite – triclinic 21**

Differential effective medium modeling with Matlab

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Seoul National University

Contents

1. GassDem (Gassmann Differential effective medium): A Matlab program for modeling the anisotropic seismic properties of porous medium
 2. Modeling Schemes: Differential effective medium theory and Gassmann's poroelastic relationship
 3. Prediction: Effect of inclusions on seismic properties of rocks
 4. Examples: How to run GassDem?
-

1. GassDem

- A MATLAB program for modeling the anisotropic seismic properties of porous medium using differential effective medium theory and Gassmann's poroelastic relationship.

GassDem

Effective Medium Inclusion Shape and Orientation DEM analysis Help

Two-component system of A and B for DEM calculation

Phase	Name	Elastic constants file (*.txt) (Note: Voigt elastic tensor in GPa)	Density (g/cm3)
A: Host medium	ModelSC3	Morales2018_ModelSC3_GPa.txt	2.729
B: Inclusion (e.g. fluid)	Basalt	Basalt1200C_GPa.txt	2.7

Input files directory: /Volumes/EKim/Research/DEM_Mai

GassDem directory: /Volumes/EKim/Research/DEM_Mai

Fluid properties

Compressibility Bf 1/GPa e.g. Water = 0.4878 1/GPa
Bf = 1/Kf Basalt = 0.0546 1/GPa at 1233 C
Kf: Bulk modulus (GPa)

Viscosity Pa.s e.g. Water = 1.0E-03 Pa.s
Basalt = 10 Pa.s

Vp km/s e.g. Water = 1.480 km/s at 20 C
Basalt = 2.610 km/s at 1233 C

Density g/cm3 e.g. Water = 1.000 g/cm3
Basalt = 2.680 g/cm3 at 1233 C

Max. % error for Green's Tensor *(e.g. 0.1)*

1. GassDem

- A MATLAB program for modeling the anisotropic seismic properties of porous medium using differential effective medium theory and Gassmann's poroelastic relationship.

GassDem

Effective Medium **Inclusion Shape and Orientation** DEM analysis Help

Ellipsoid semi-axes lengths A1:A2:A3

Sphere 1:1:1 Ellipsoid 5 : 5 : 1

Orientation of shape ellipsoid semi-axes

Parallel to elastic tensor axes Parallel to user-defined orientation (SPO)

Ellipsoid axes A // Elastic tensor axes X (for fluid inclusion)

A1 // X1 (e.g. olivine = [100]: North)
A2 // X2 (e.g. olivine = [010]: East)
A3 // X3 (e.g. olivine = [001]: Up)

X1 = 100 in olivine, X1 = a in calcite etc.

	Azimuth (degree)	Inclination (degree)
A1	<input type="text" value="90"/>	<input type="text" value="0"/>
A3	<input type="text" value="0"/>	<input type="text" value="0"/>

A1 should be 90 degrees to A3

Orientaton is OK?

Check 1: Angle A1 to A3 =

Check 2: Angle A1 to A3 =

Check 3: Det|R| should be +1

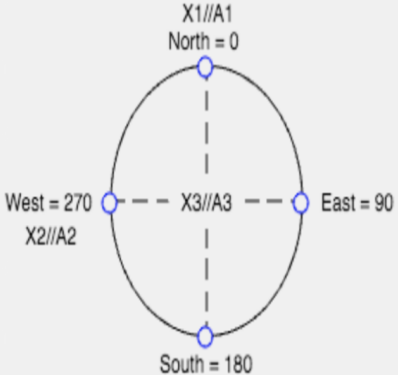
Det|R| =

Shape ellipsoid rotation matrix

<input type="text" value="0"/>	<input type="text" value="0"/>	<input type="text" value="1"/>
<input type="text" value="-1"/>	<input type="text" value="0"/>	<input type="text" value="0"/>
<input type="text" value="0"/>	<input type="text" value="-1"/>	<input type="text" value="0"/>

Fixed orientation of shape ellipsoid axes

A1: AZ =	<input type="text" value="90"/>	INC =	<input type="text" value="0"/>
A2: AZ =	<input type="text" value="360"/>	INC =	<input type="text" value="90"/>
A3: AZ =	<input type="text" value="360"/>	INC =	<input type="text" value="0"/>



1. GassDem

- A MATLAB program for modeling the anisotropic seismic properties of porous medium using differential effective medium theory and Gassmann's poroelastic relationship.

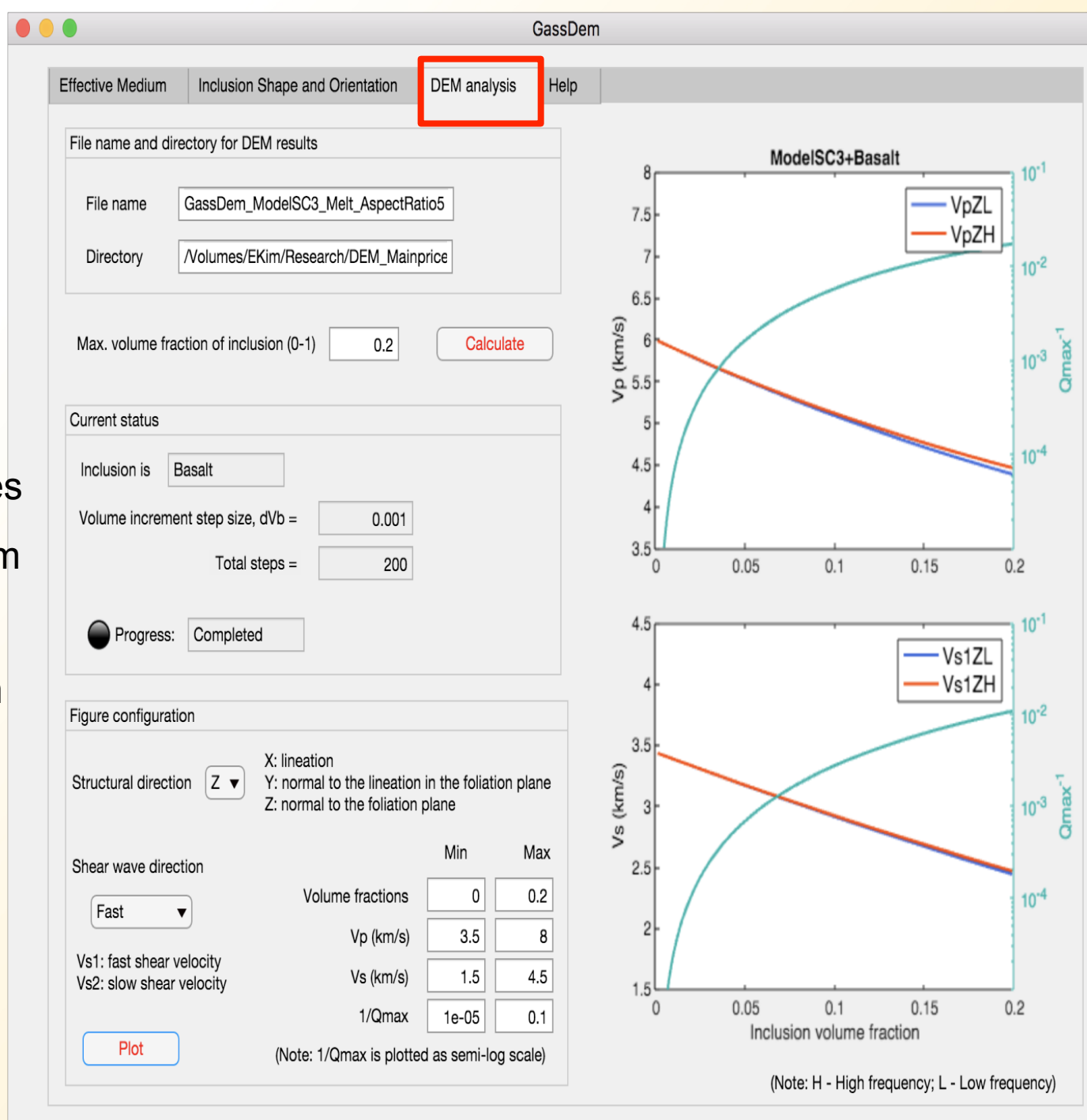


Table of elastic stiffness constants for cubic minerals.

Mineral	Structure	Density (g/cm ³)	C ₁₁ (GPa)	C ₁₂ (GPa)	C ₄₄ (GPa)
Pyrope	Garnet	3.839	301.4	110.0	94.3
Ringwoodite	Spinel	3.559	327.0	114.0	131.0
Ferropériclase	Halite	4.323	499.0	137.0	160.0

H

Table of results cubic symmetry ($m\bar{3}m$) pure modes for propagation along [100],[110] and [111] with polarizations, equations and velocities (km/s) for each mineral.

Propagation	Polarization	Equation	Pyrope	Ringwood	Periclase
[100]	[100]	$V_p = \sqrt{\frac{C_{11}}{\rho}}$	8.86	9.58	10.74
[100]	[010]	$V_s = \sqrt{\frac{C_{44}}{\rho}}$	4.96	6.07	6.08
[110]	[110]	$V_p = \sqrt{\frac{C_{11}+C_{12}+2C_{44}}{2\rho}}$	8.84	9.94	10.52
[110]	$[\bar{1}10]$	$V_s = \sqrt{\frac{C_{11}-C_{12}}{2\rho}}$	4.99	5.47	6.47
[110]	[010]	$V_s = \sqrt{\frac{C_{44}}{\rho}}$	4.96	6.07	6.08
[111]	[111]	$V_p = \sqrt{\frac{C_{11}+2C_{12}+4C_{44}}{3\rho}}$	8.83	10.05	10.44
[111]	in plane(111)	$V_s = \sqrt{\frac{C_{11}-C_{12}+C_{44}}{3\rho}}$	5.68	5.67	6.34

To determine the elastic constants C_{11} and C_{44} ; $\rho(V_p^{[100]})^2 = C_{11}$ and $\rho(V_s^{[100]})^2 = C_{44}$
From the equations above determine C_{12} .

Hint : generic velocity $V = \sqrt{M/\rho}$ units km/s and ρ units g/cm³ and generic modulus is $M = \rho * V^2$ (units GPa) use either

$$\rho(V_p^{[110]})^2 = \sqrt{\frac{C_{11} + C_{12} + 2C_{44}}{2\rho}}$$

or

$$\rho(V_p^{[111]})^2 = \sqrt{\frac{C_{11} + 2C_{12} + 4C_{44}}{3\rho}}$$

Thank you

- I thank Ralf for inviting me to this workshop.
- I thank you all for listening.
- Most things I talked about (programs, pdf of publications, MTEX examples and link to the MTEX site) can be accessed via my webpage
<http://www.gm.univ-montp2.fr/PERSO/mainprice/>
- Also look at MTEX website
<http://mtex-toolbox.github.io>